



UTAH GEOLOGICAL SURVEY

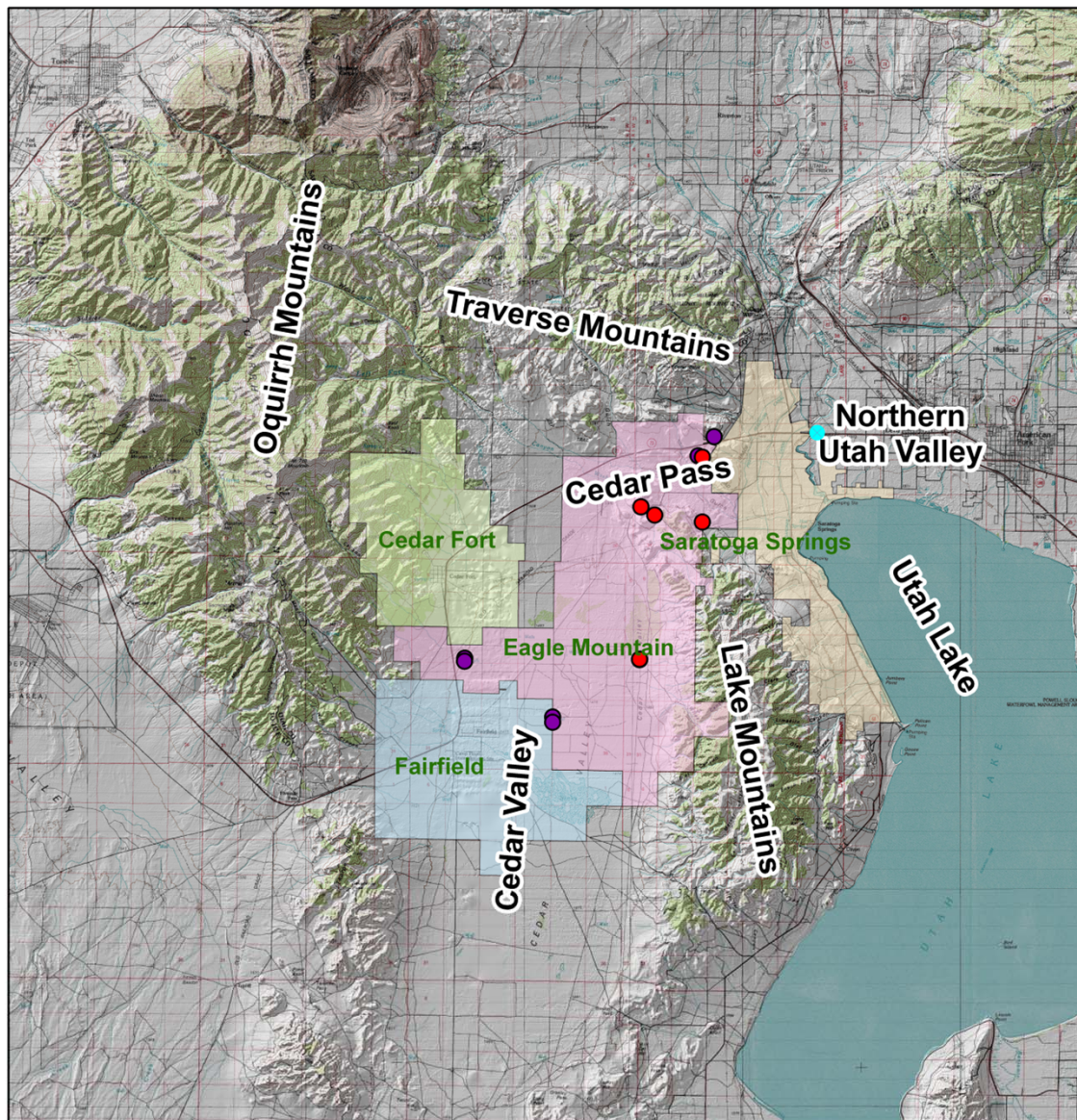
Geologic Factors Controlling the Response of Fractured Bedrock, Basin-fill, and Perched Aquifers to Recent Groundwater Development in the Cedar Pass Area, Eagle Mountain City, North-central Utah

J. Lucy Jordan

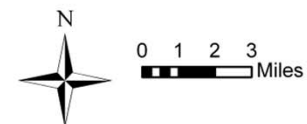
Utah Geological Survey - Groundwater Program

Rocky Mountain/Cordilleran
Geological Society of
America meeting
May 19, 2011





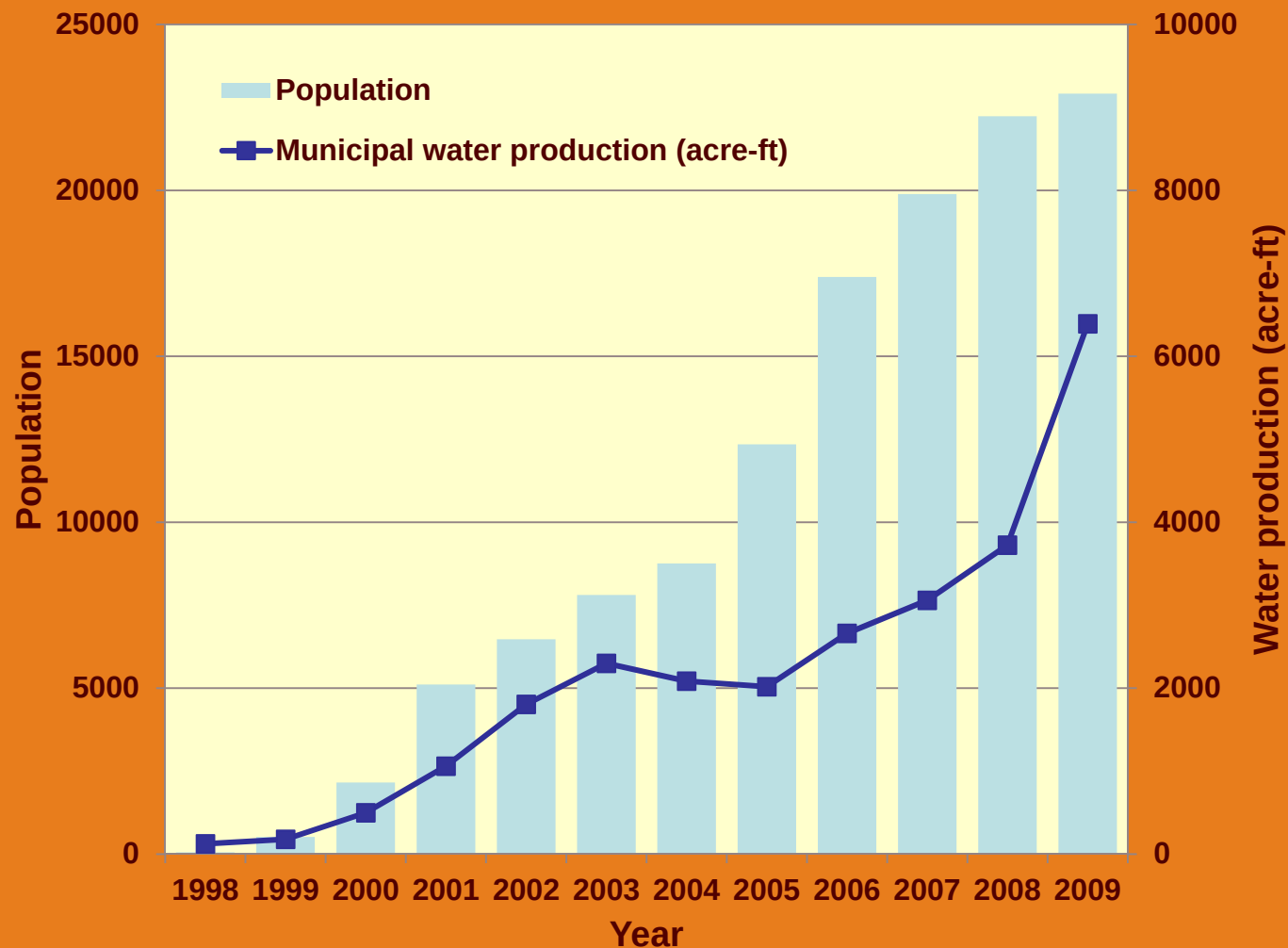
- EM production wells
- Other production wells





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Population Growth and Municipal Groundwater Pumping for Eagle Mountain City





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New residential development at Cedar Pass



Oquirrh Mountains primary recharge area and Eagle Mountain City



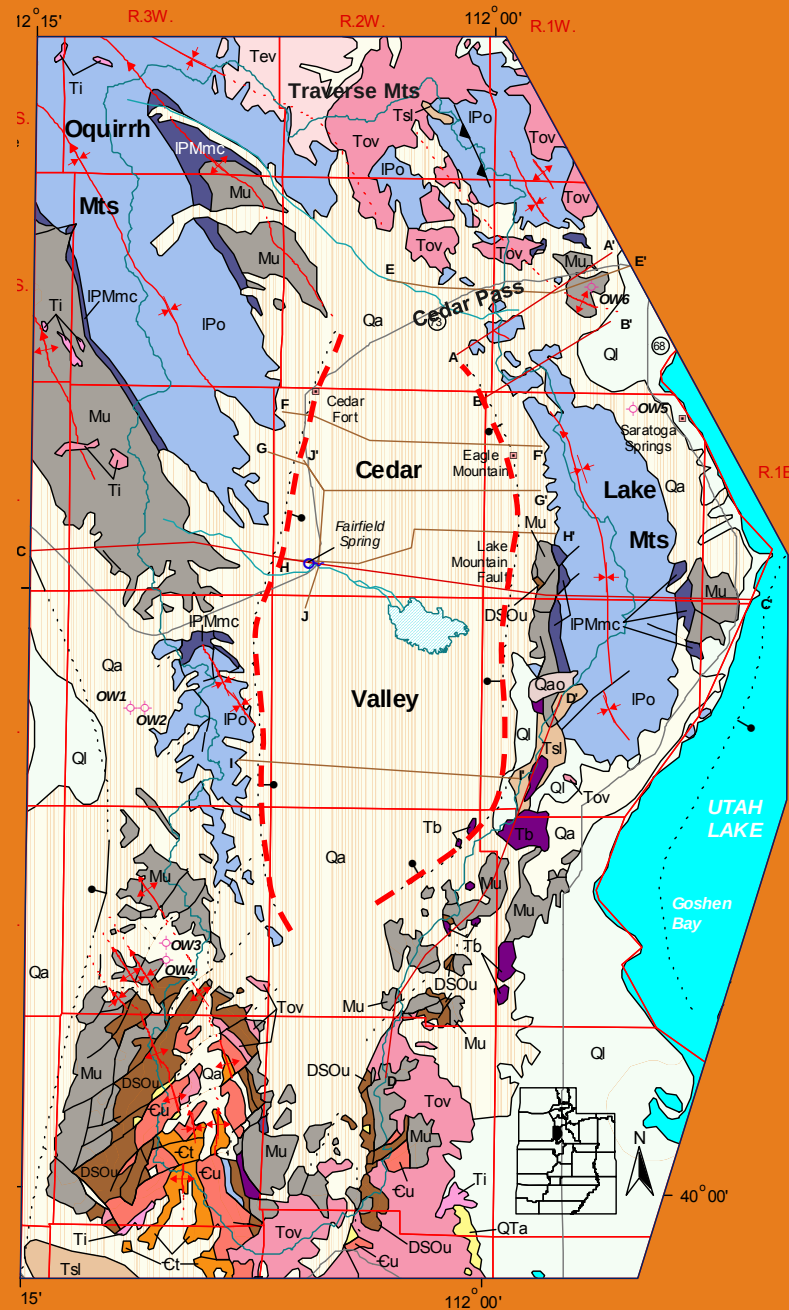
Work Performed

- ✓ Quarterly water-level monitoring 2007-2009 on 18 wells
- ✓ Measured water level in 27 private wells in Cedar Pass area.
- ✓ 3 monitoring wells drilled at 2 locations
- ✓ Long-term aquifer tests on 2 large production wells
- ✓ Water-quality sampling of 12 wells for general chemistry and isotopes

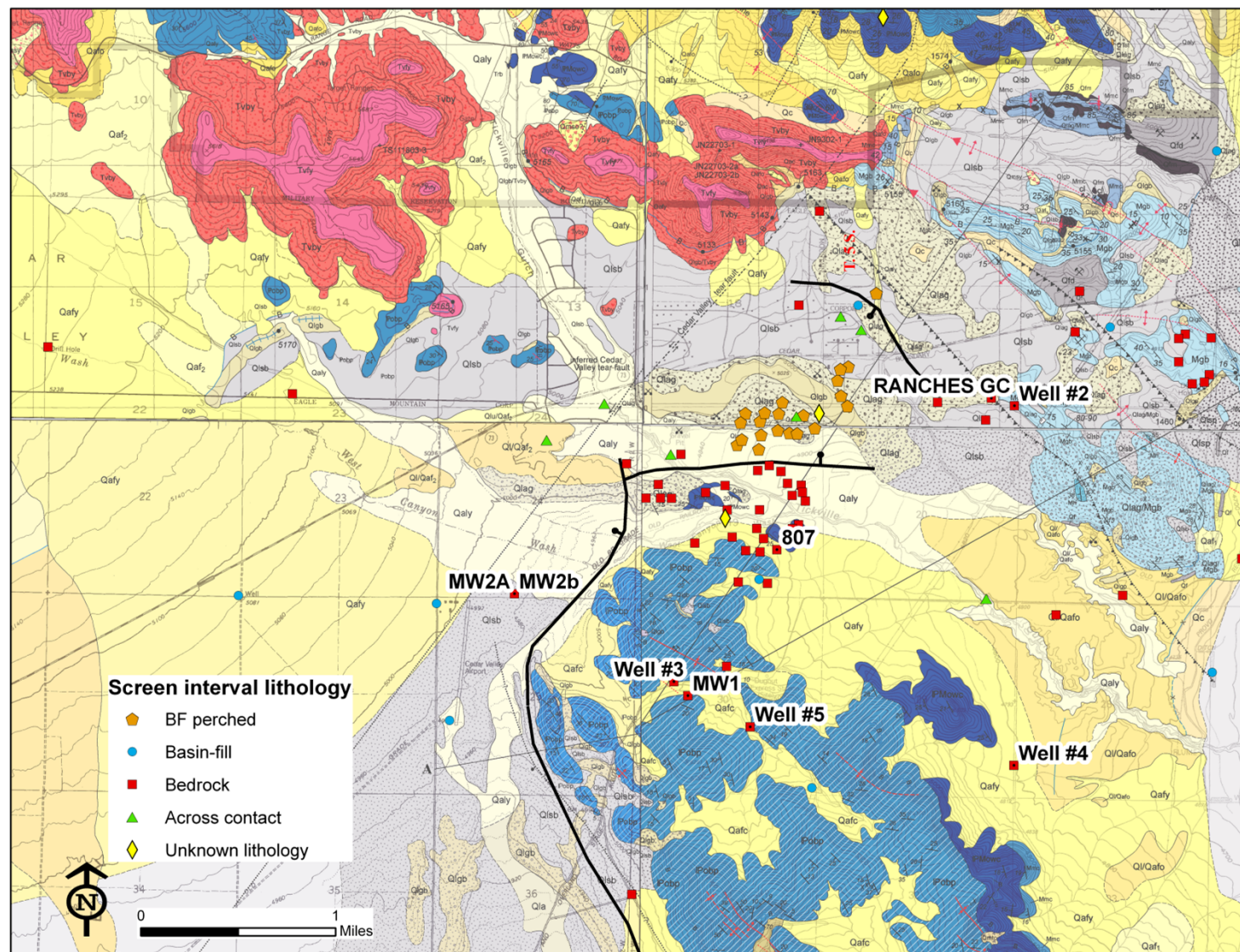


Generalized Geologic Map

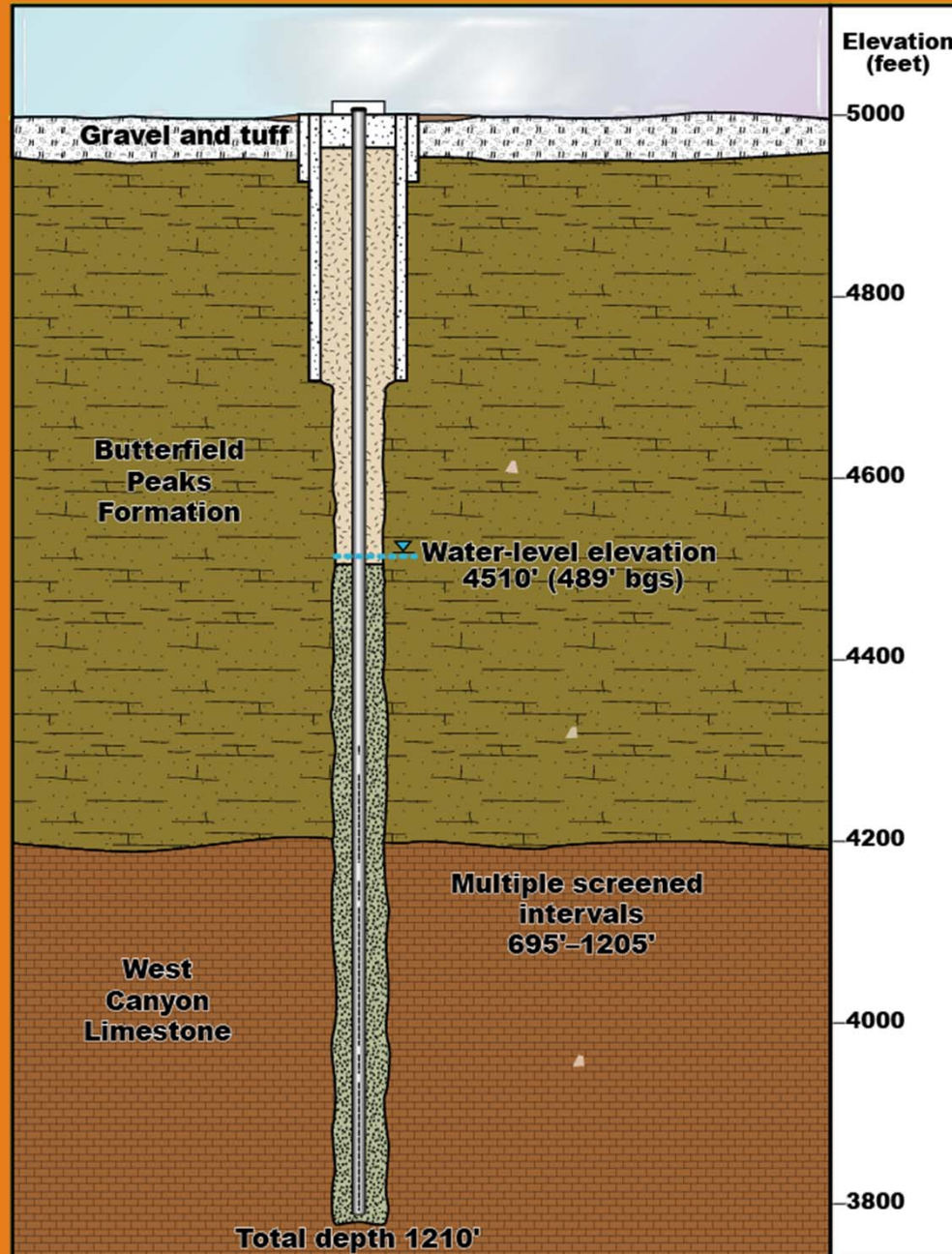
From Hurlow, H.A., 2004,
The geology of Cedar
Valley, Utah County, Utah,
and its relation to ground-
water conditions, Utah
Geological Survey Special
Study, Utah Geological
Survey 109.



Cedar Pass Geology



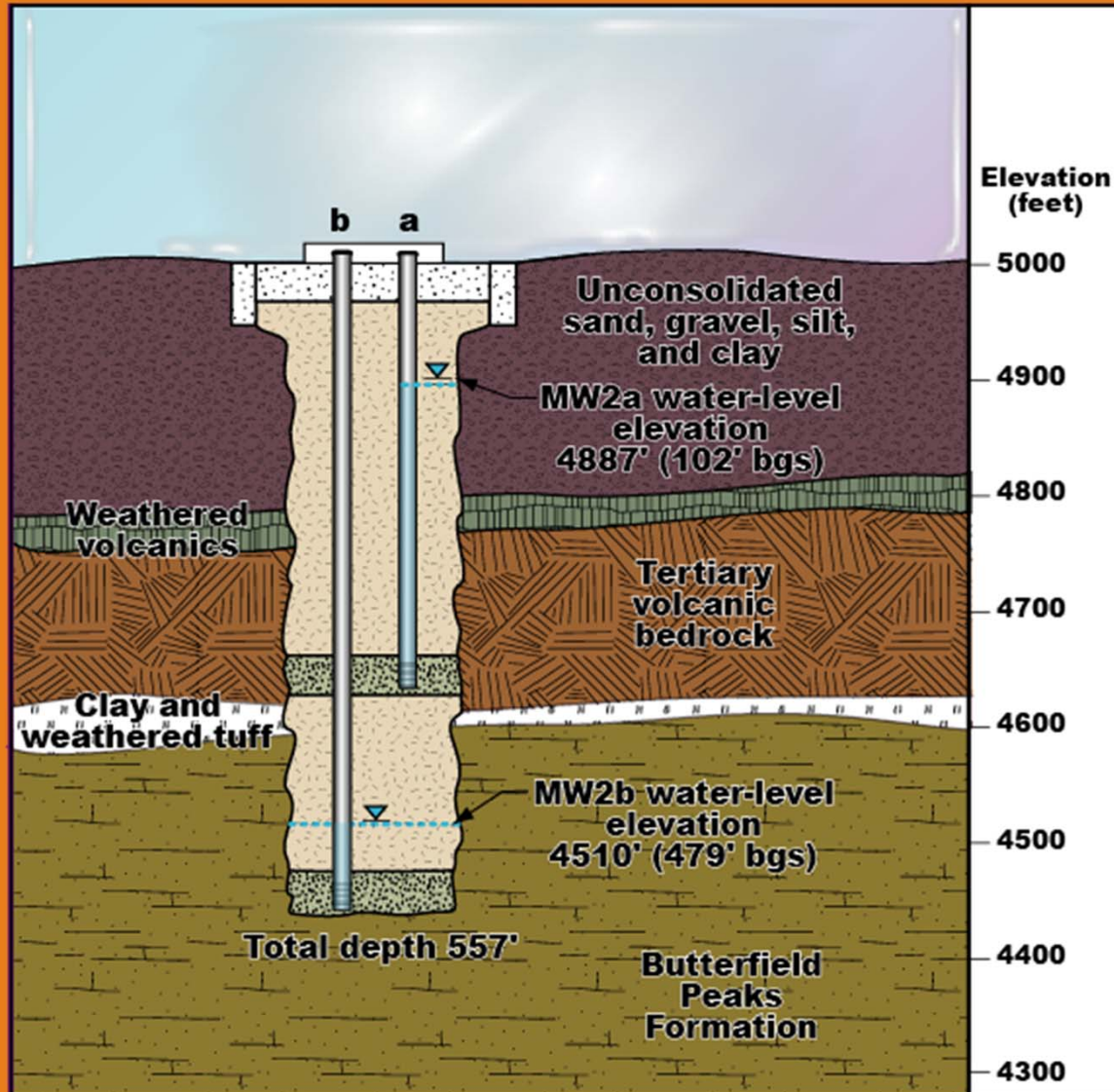
Fault locations refined by this study and Cederberg et al, 2009.



MW1

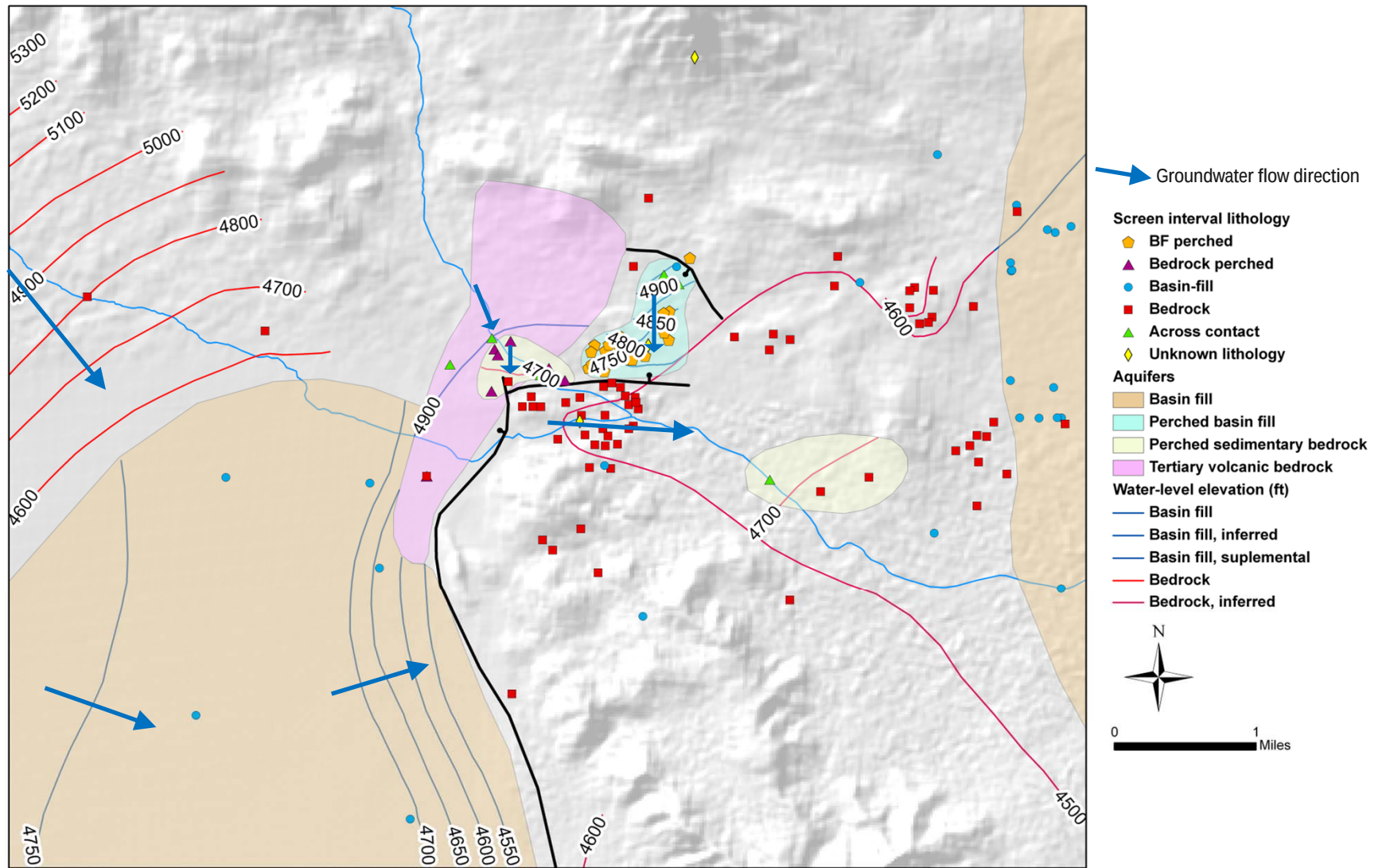
- Designed to mimic completion interval of Eagle Mountain production well #3
- Well #3 located 500 feet NW along strike of small anticline.

MW2ab



- Double completion well designed to have upper well in basin fill and lower well in Paleozoic aquifer
- Volcanics came in higher than saturated basin fill
- Unsaturated zone below volcanics/above unconfined Paleozoic aquifer

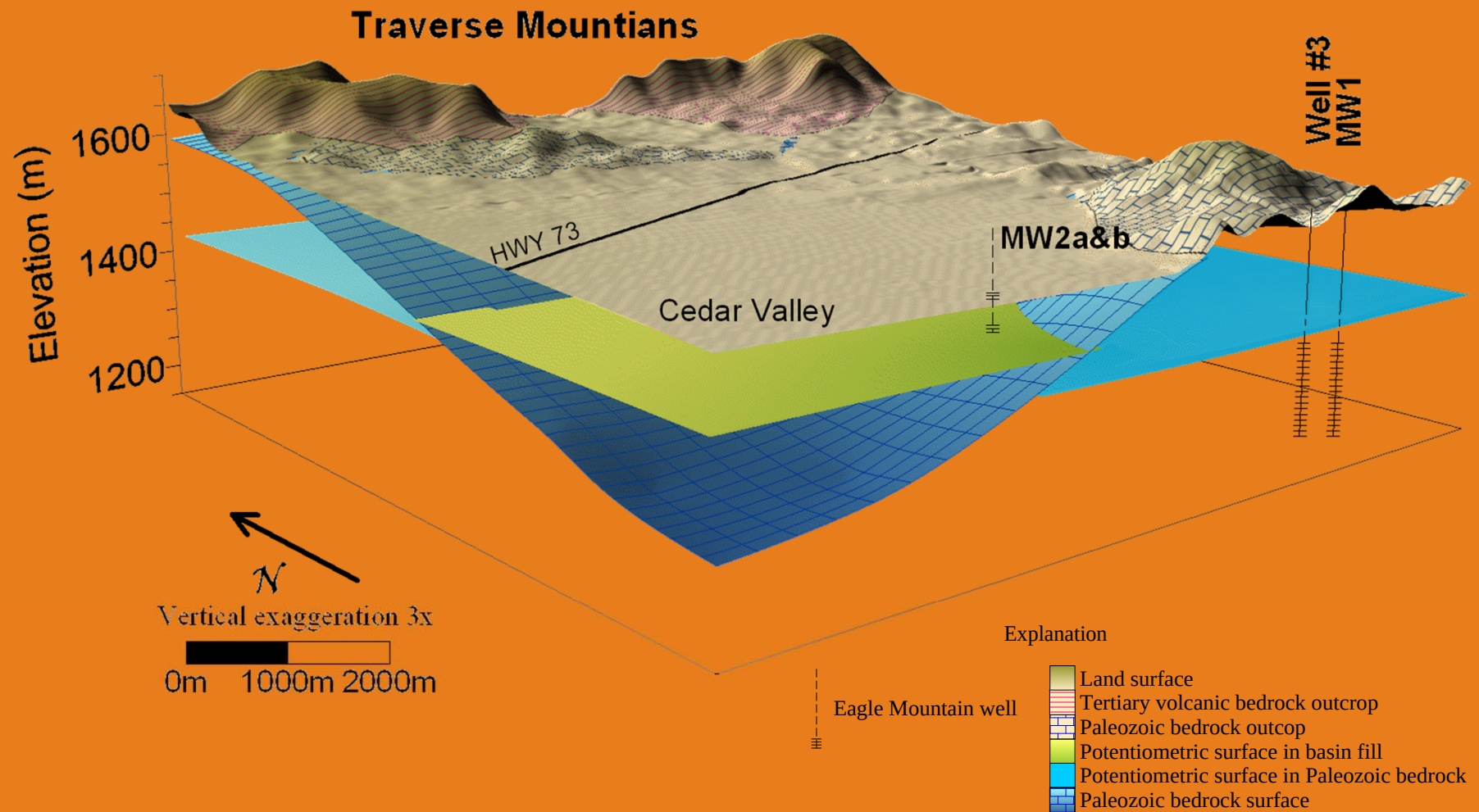
Cedar Pass area aquifers and potentiometric contours





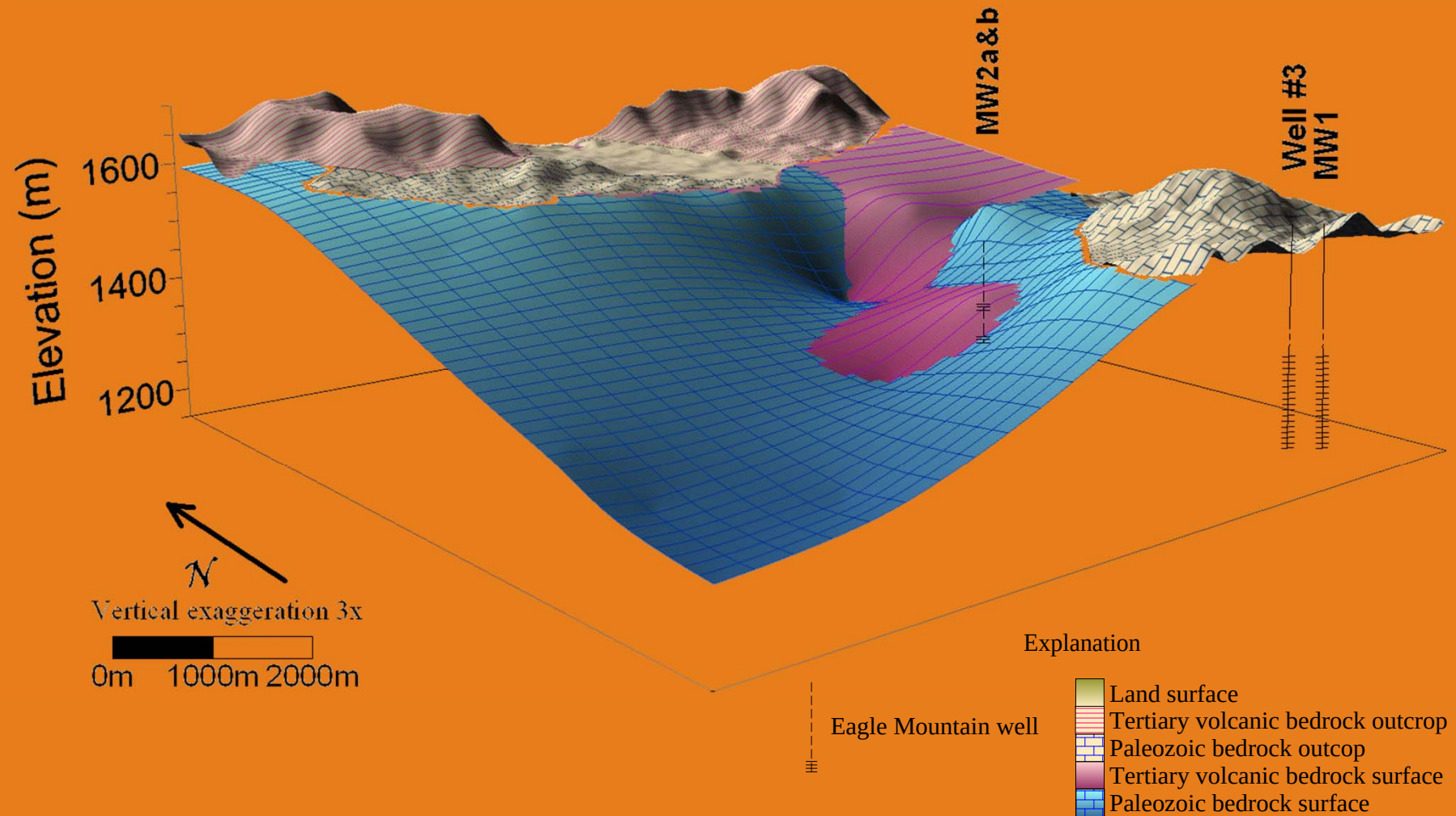
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Hydrostratigraphy of Cedar Pass area





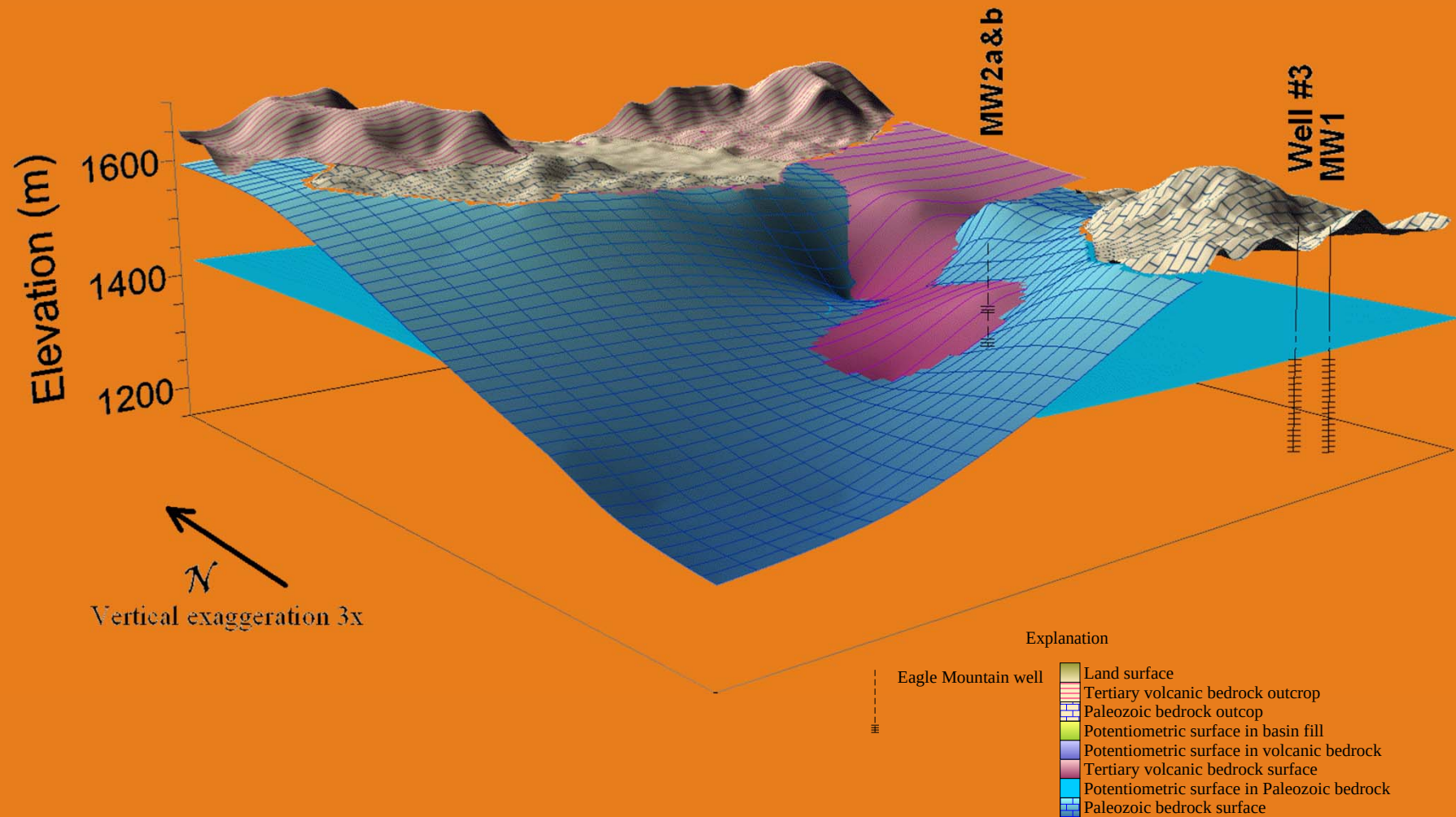
Hydrostratigraphy of Cedar Pass area – geologic units





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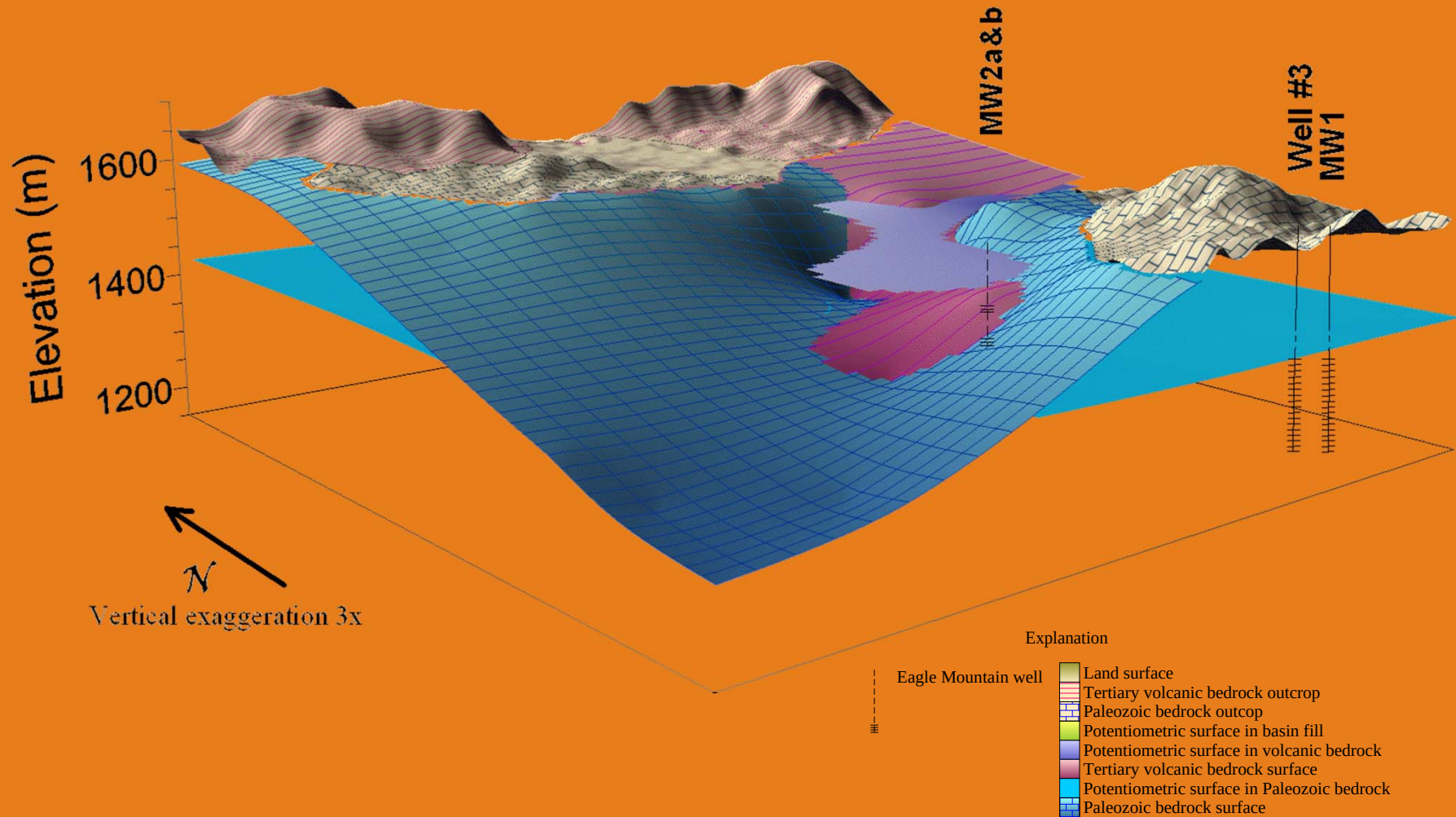
Hydrostratigraphy of Cedar Pass area – geologic units + potentiometric surface in the Paleozoic aquifer





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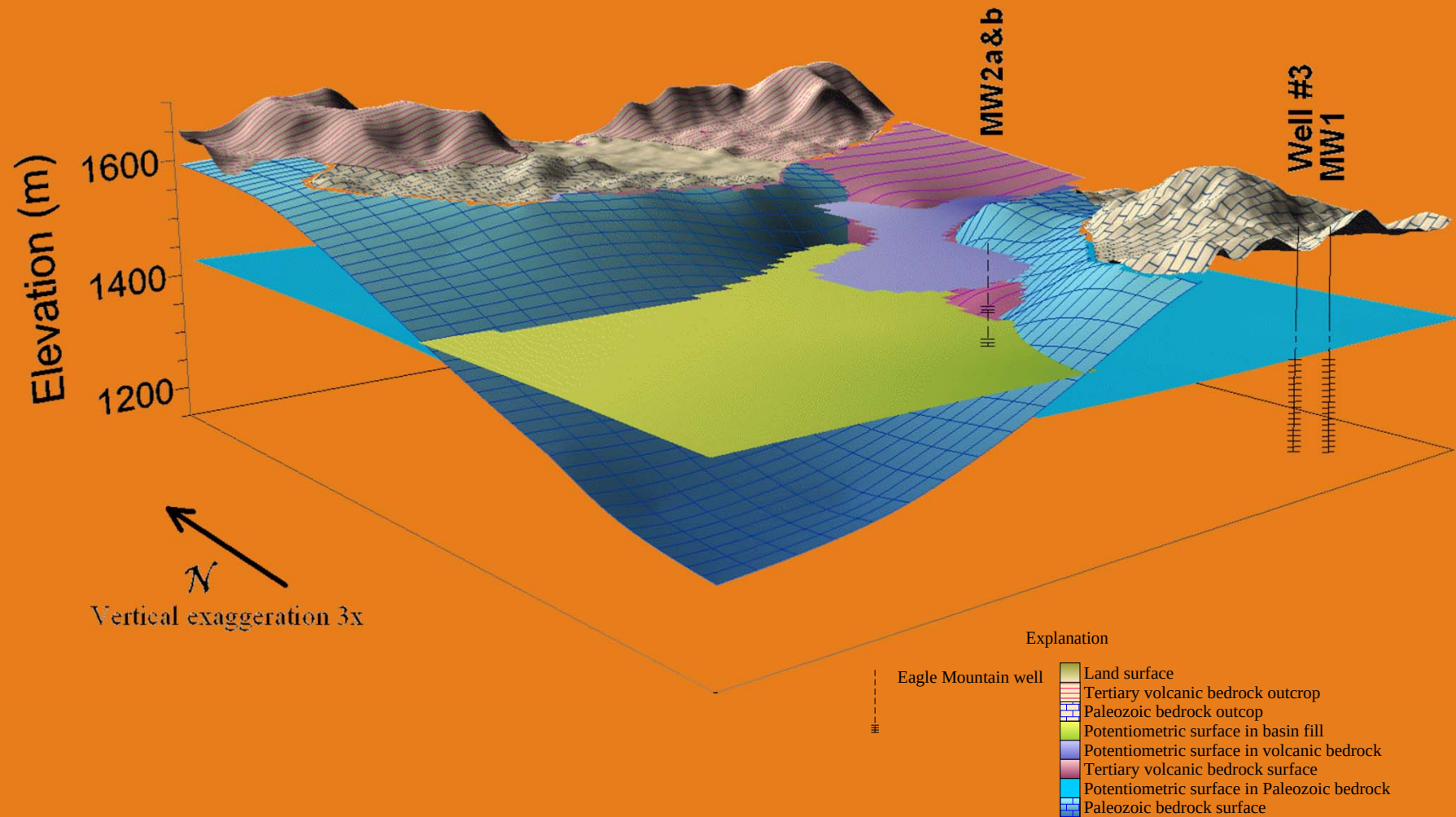
Hydrostratigraphy of Cedar Pass area – geologic units + potentiometric surface in the Paleozoic and volcanic aquifers





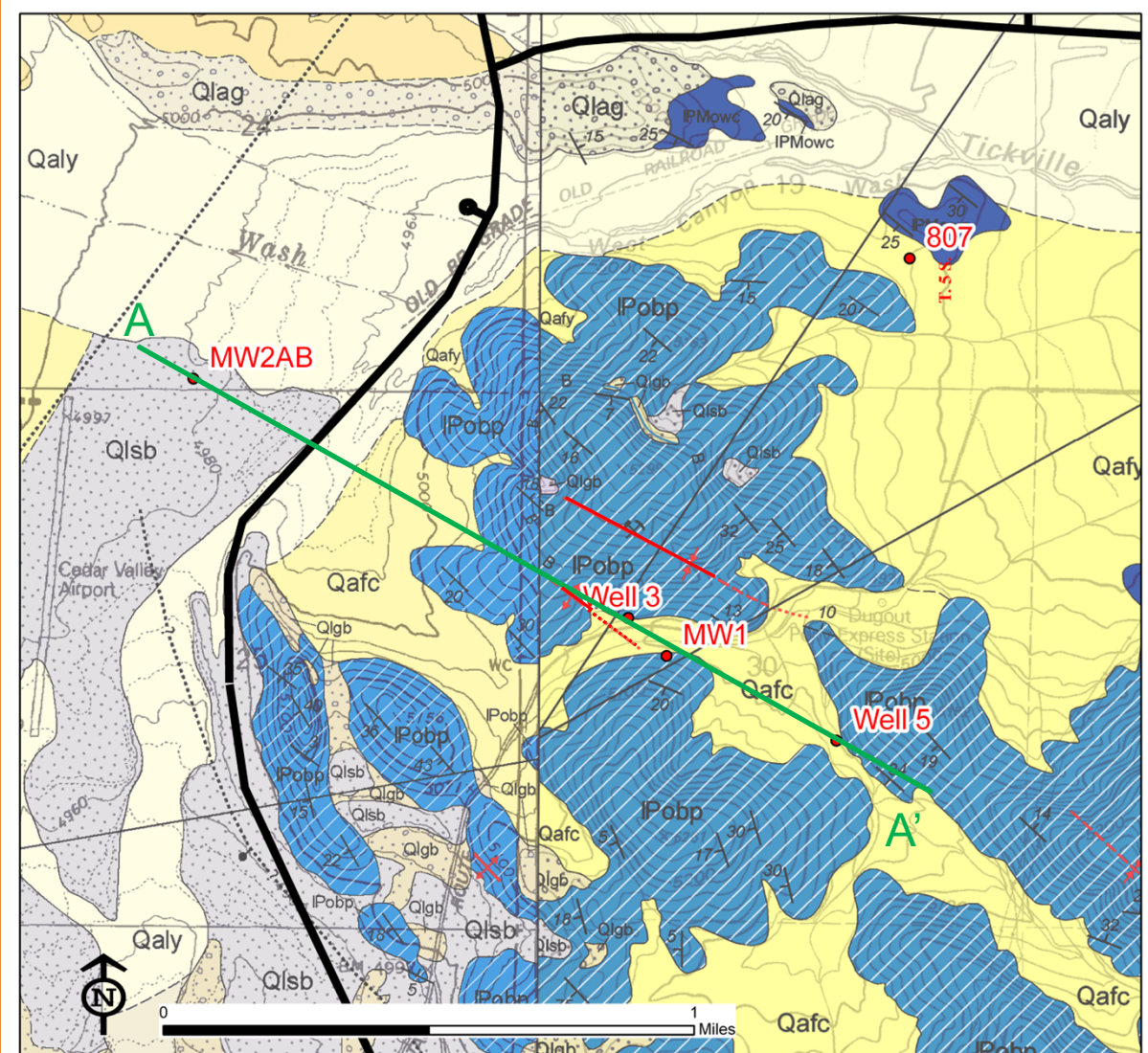
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Hydrostratigraphy of Cedar Pass area – geologic units + potentiometric surfaces



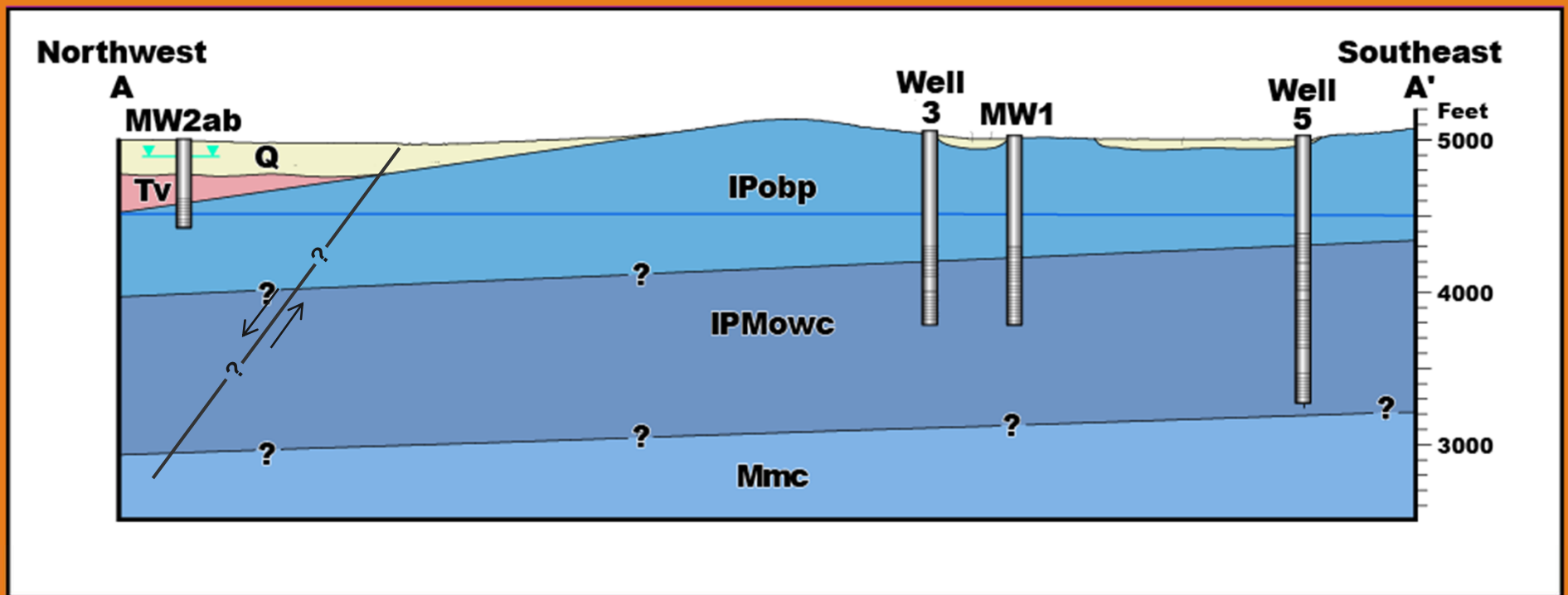
Aquifer test: Eagle Mountain Well #3

- Well #3: 16"-diameter well, screened 700-1200 ft below ground level in Pennsylvanian-Mississippian Oquirrh Group (Butterfield Peaks and West Canyon Limestone Fms)
- 5 months long, 1930 gpm
- MW1 ~500 feet from pumping well and designed to monitor aquifer response over entire thickness of aquifer
- MW2b and 807 ~ 5000 feet distant
- MW1 and MW2b along direction of primary T, 807 perpendicular to primary T



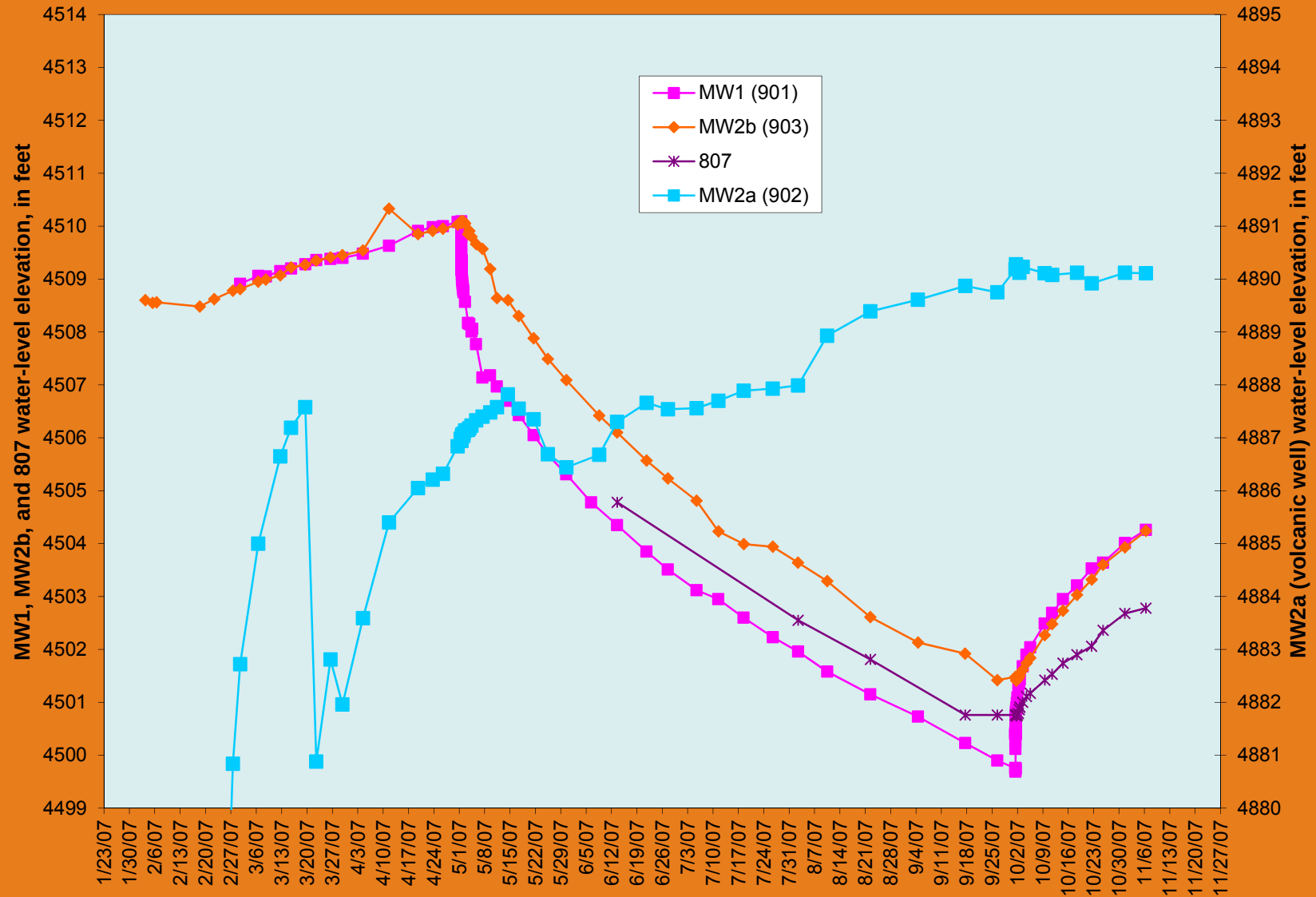


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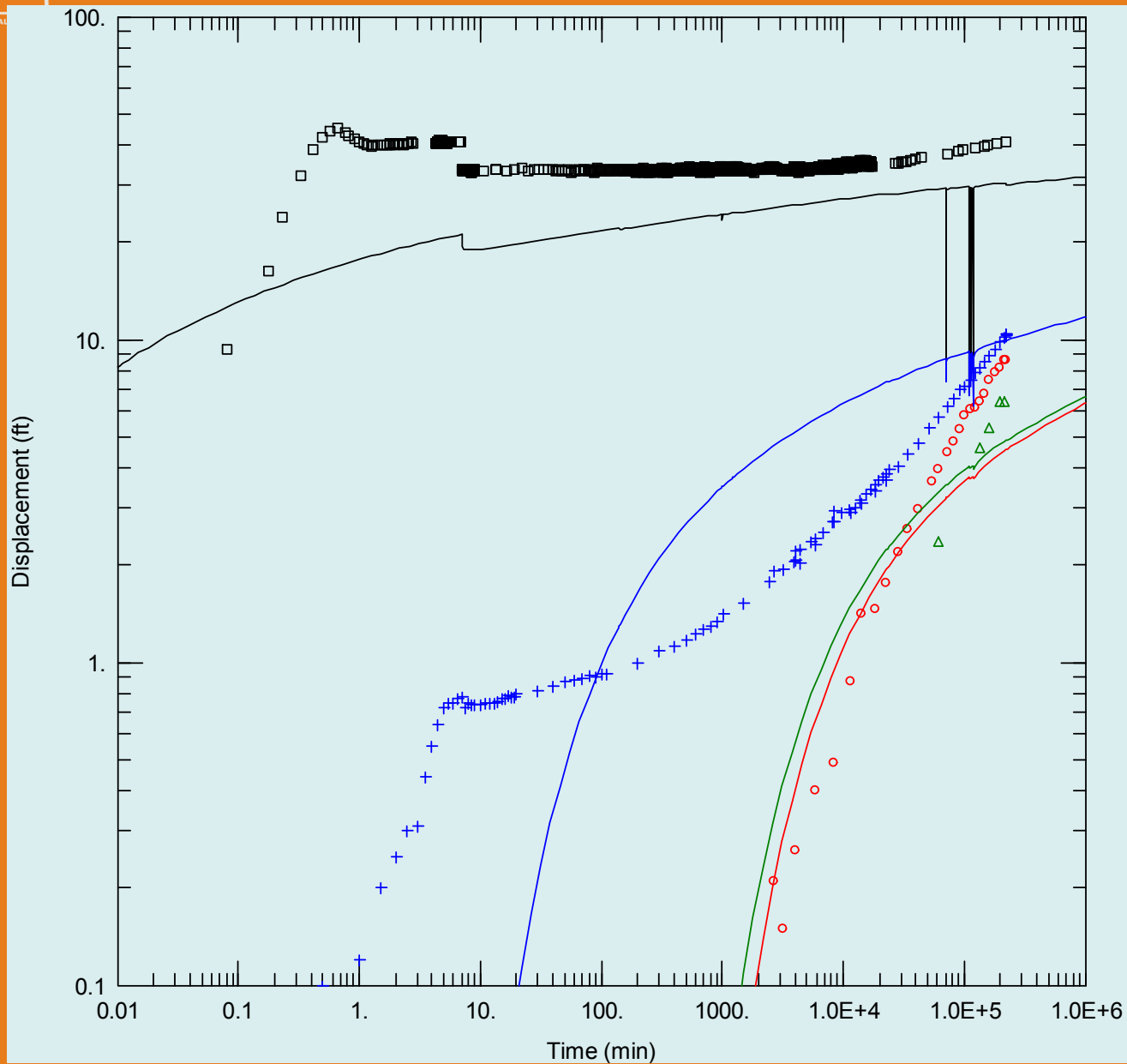


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Obs. Wells

- Well 3 (992)
- + MW1 (901)
- MW2b (903)
- △ 807

Aquifer Model

Confined

Solution

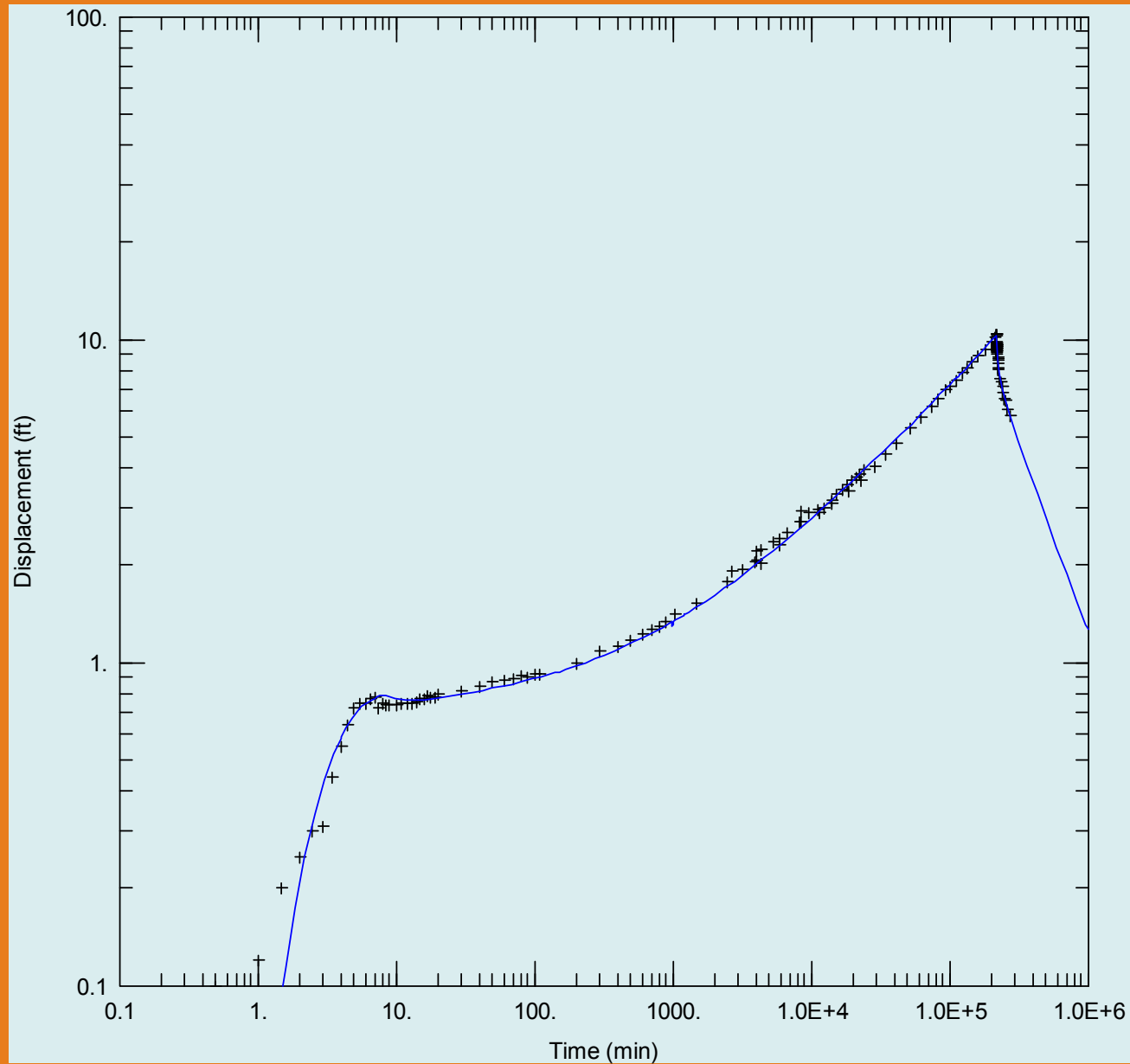
Theis

Parameters

$T = 2.385E+4 \text{ ft}^2/\text{day}$
 $S = 0.008467$
 $Kz/Kr = 1.$
 $b = 530. \text{ ft}$



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Obs. Wells

+ MW1 (901)

Aquifer Model

Fractured

Solution

Moench w/spherical blocks

Parameters

$K = 10.73 \text{ ft/day}$

$S_s = 3.316\text{E-}7 \text{ ft}^{-1}$

$K' = 1.435\text{E-}5 \text{ ft/day}$

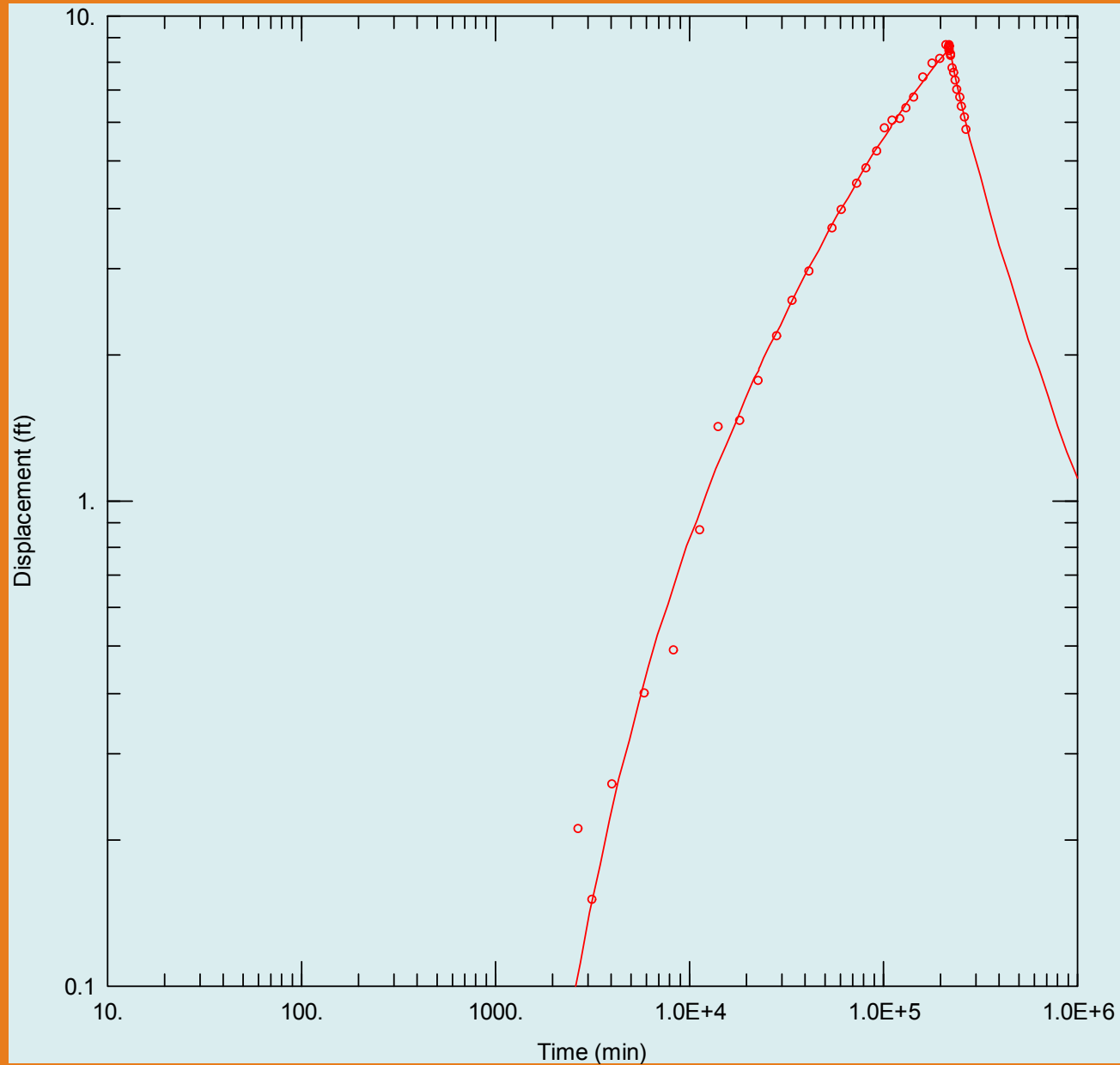
$S_s' = 0.001991 \text{ ft}^{-1}$

$Sw = 0.$

$Sf = 0.08446$



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Obs. Wells

○ MW2b (903)

Aquifer Model

Fractured

Solution

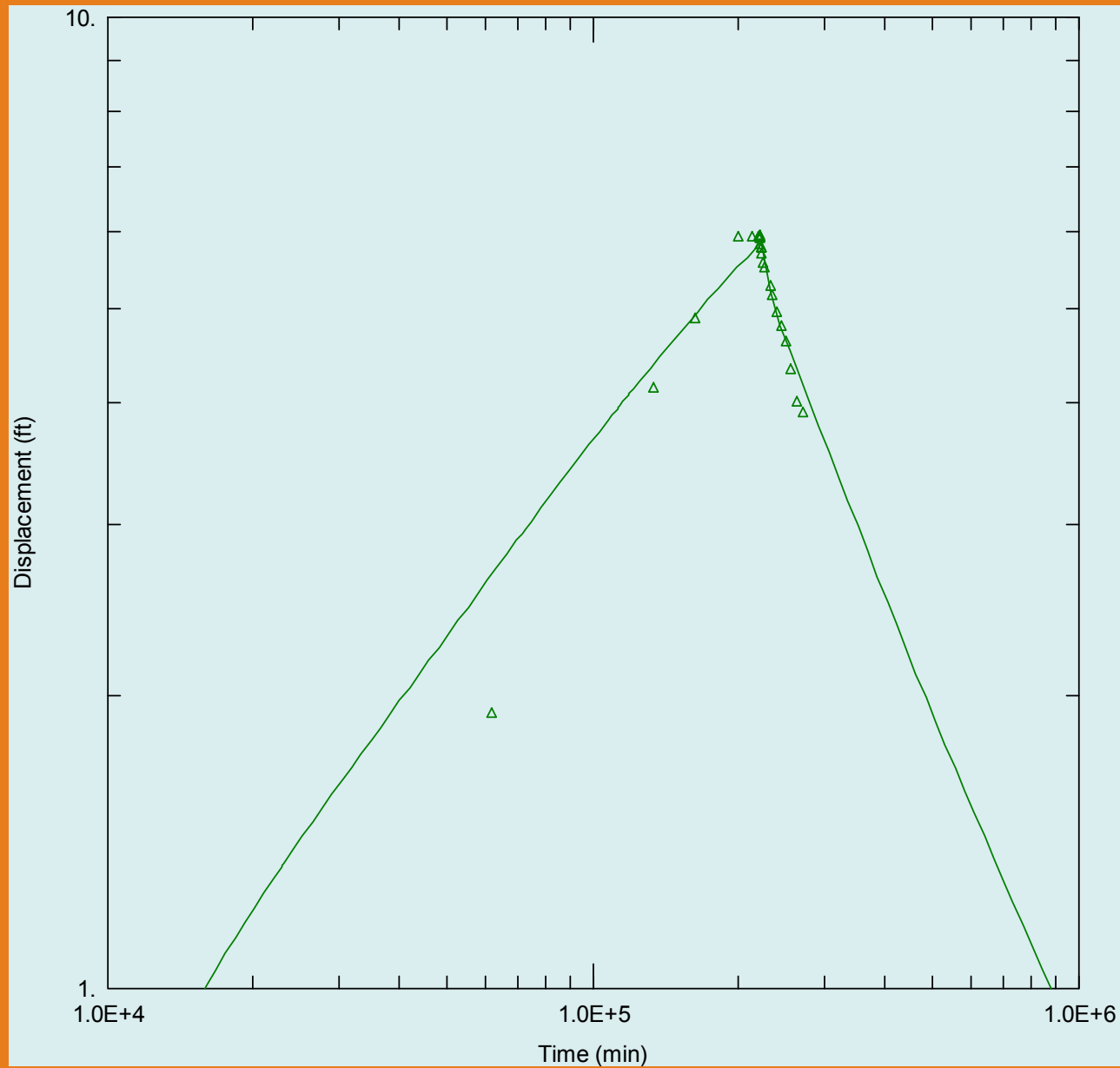
Moench w/spherical blocks

Parameters

$K = 12.18 \text{ ft/day}$
 $S_s = 4.953\text{E-}6 \text{ ft}^{-1}$
 $K' = 4.239\text{E-}7 \text{ ft/day}$
 $S_s' = 2.416\text{E-}5 \text{ ft}^{-1}$
 $Sw = 20.$
 $Sf = 0.$



UTAH GEOLOGICAL SURVEY



Obs. Wells

△ 807

Aquifer Model

Fractured

Solution

Moench w/spherical blocks

Parameters

$K = 15.44 \text{ ft/day}$

$S_s = 1.748\text{E-}5 \text{ ft}^{-1}$

$K' = 1.132\text{E-}6 \text{ ft/day}$

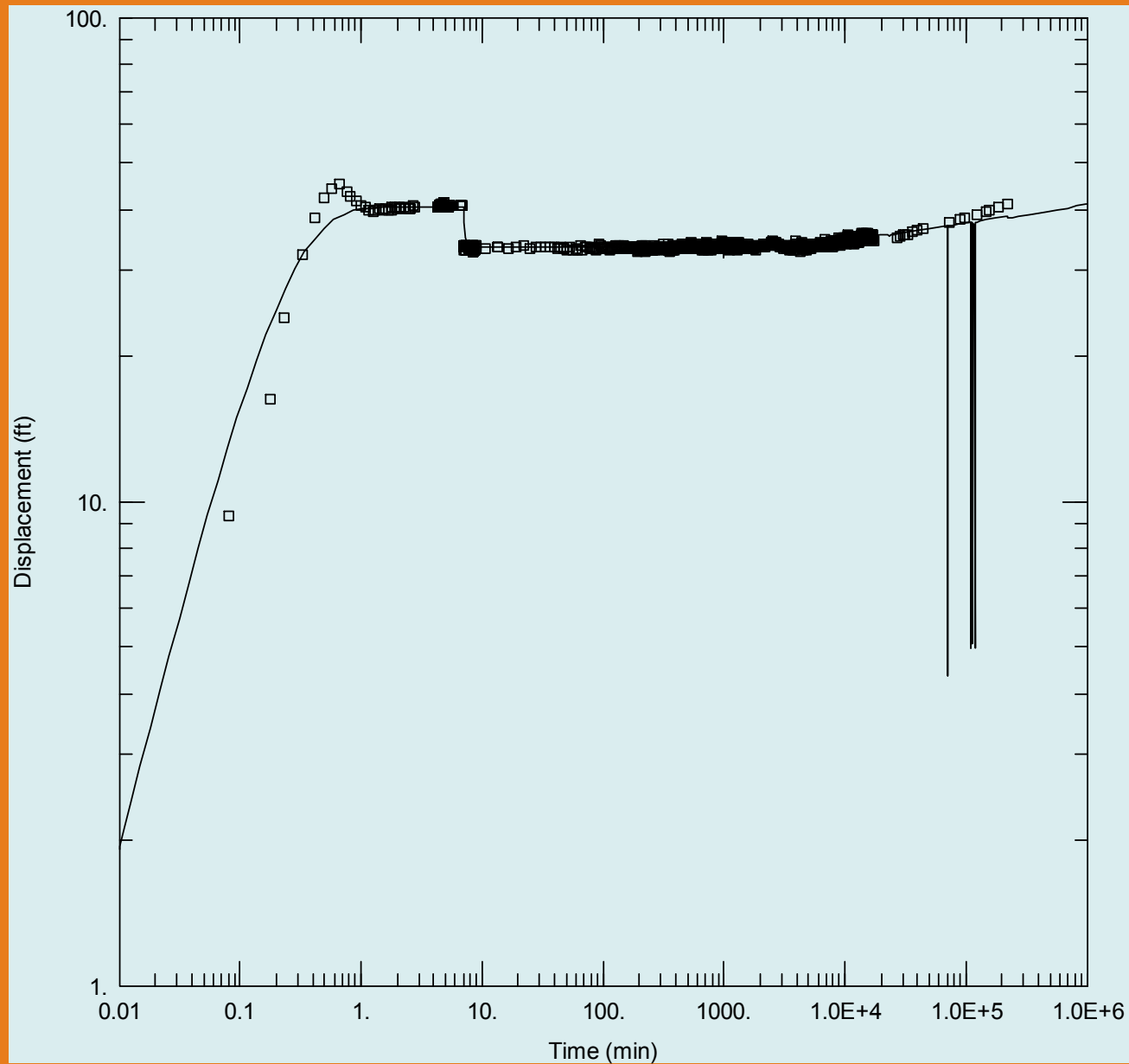
$S_s' = 4.654\text{E-}5 \text{ ft}^{-1}$

$S_w = 0.$

$S_f = 0.3762$



UTAH GEOLOGICAL SURVEY



Obs. Wells

□ Well 3 (992)

Aquifer Model

Fractured

Solution

Moench w/spherical blocks

Parameters

$K = 30.73 \text{ ft/day}$
 $S_s = 6.156\text{E-}11 \text{ ft}^{-1}$
 $K' = 2.879\text{E-}7 \text{ ft/day}$
 $S_s' = 4.619\text{E-}6 \text{ ft}^{-1}$
 $Sw = 1.128$
 $Sf = 0.6061$

Determining transmissivity anisotropy

- Method: Heilweil, V.M., and Hsieh, P.A., 2006, Determining anisotropic transmissivity using a simplified Papadopoulos method: Ground Water, v. 44, no. 5, p. 749-753.

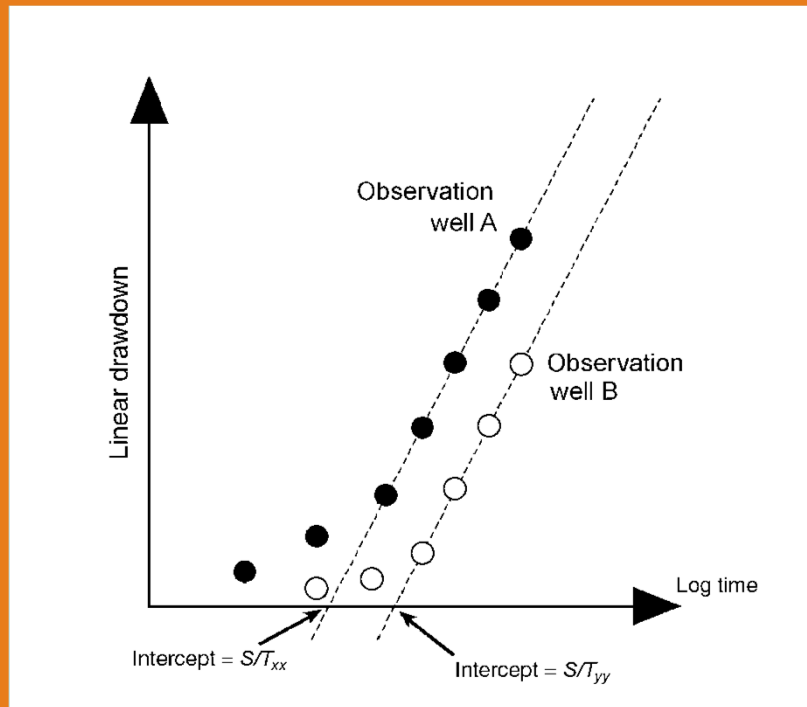
Papadopoulos, 1965

$$s(x, y, t) = \frac{Q}{4\pi\sqrt{T_{xx}T_{yy}}} W(u_{xy})$$

- Simplification of Papadopoulos
- Fractured aquifers in which principal directions of the transmissivity are known
- Assumes that observation wells are drilled along the two principal directions of transmissivity from the pumped well
- Only 2 observation wells needed

$$u_{xy} = \frac{S}{4T} \left(\frac{T_{xx}y^2 + T_{yy}x^2}{T_{xx}T_{yy}} \right)$$

- ✓ Cooper-Jacob method: slope of *late time data* (straight line part) is used to calculate T (Δs over 1 log cycle time into equation) and intercept gives S/T



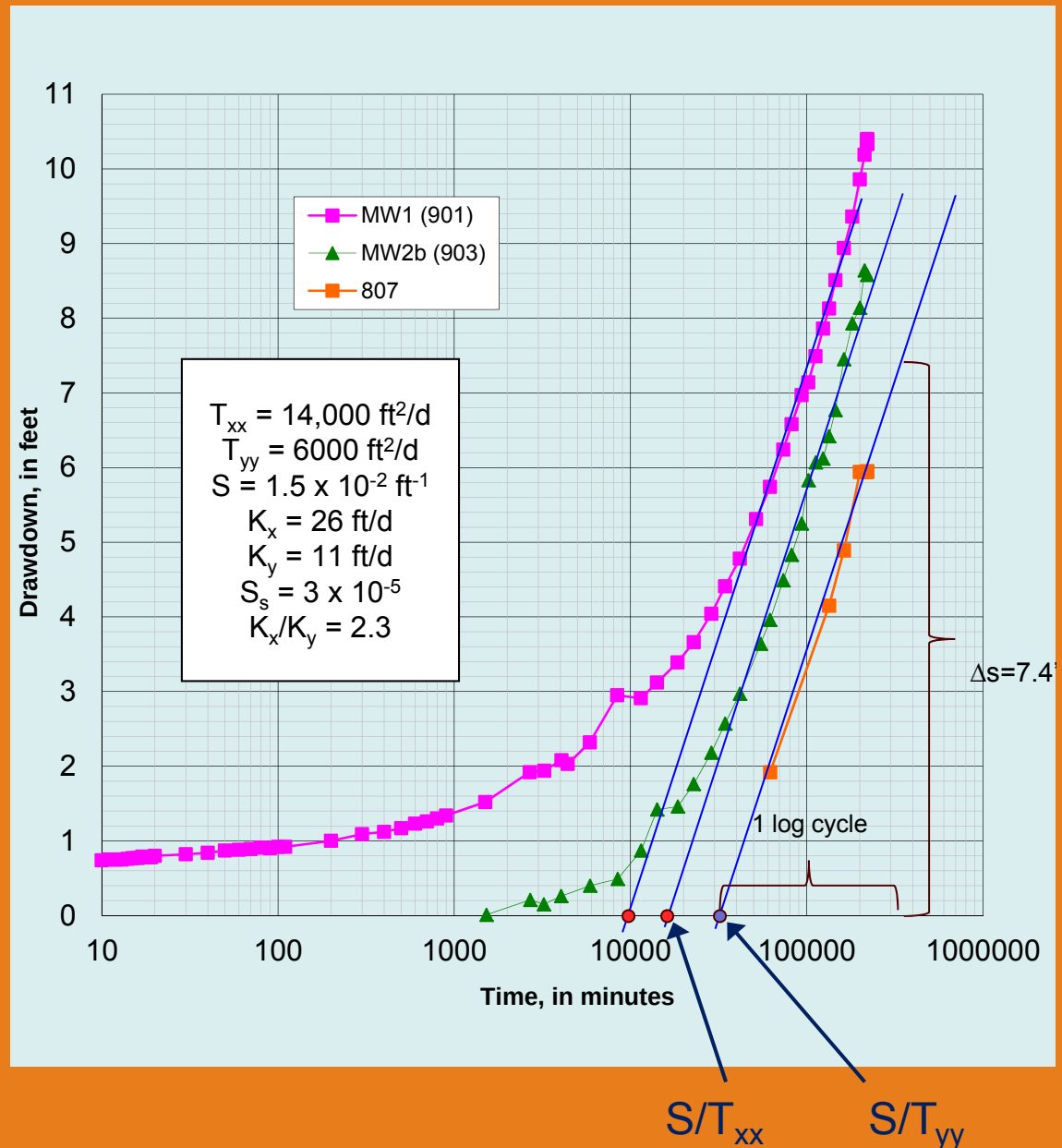
$$\sqrt{T_{xx}T_{yy}} = \frac{2.3Q}{4\pi\Delta s}$$

$$\frac{S}{T_{yy}} = \frac{2.25t_{ob}}{[y_B]^2}$$

- 1) Square $\sqrt{T_{xx}T_{yy}}$ to obtain $T_{xx}T_{yy}$
- 2) Multiply S/T_{xx} and S/T_{yy} to obtain $S^2/(T_{xx}T_{yy})$
- 3) Multiply the result from steps 1 and 2 to get S^2 , then take the square root to arrive at S
- 4) Divide the S obtained from step 3 by S/T_{xx} to get T_{xx}
- 5) Divide the S obtained from step 3 by S/T_{yy} to get T_{yy} .

Determine transmissivity anisotropy

Semi-log plot of drawdown versus time in monitoring wells during Well 3 aquifer test. Parallel lines are fitted to the straightest part of the late-time data. The value of drawdown over one log cycle, 7.4 feet, is used as input for determining anisotropic transmissivity in fractured-rock aquifers.





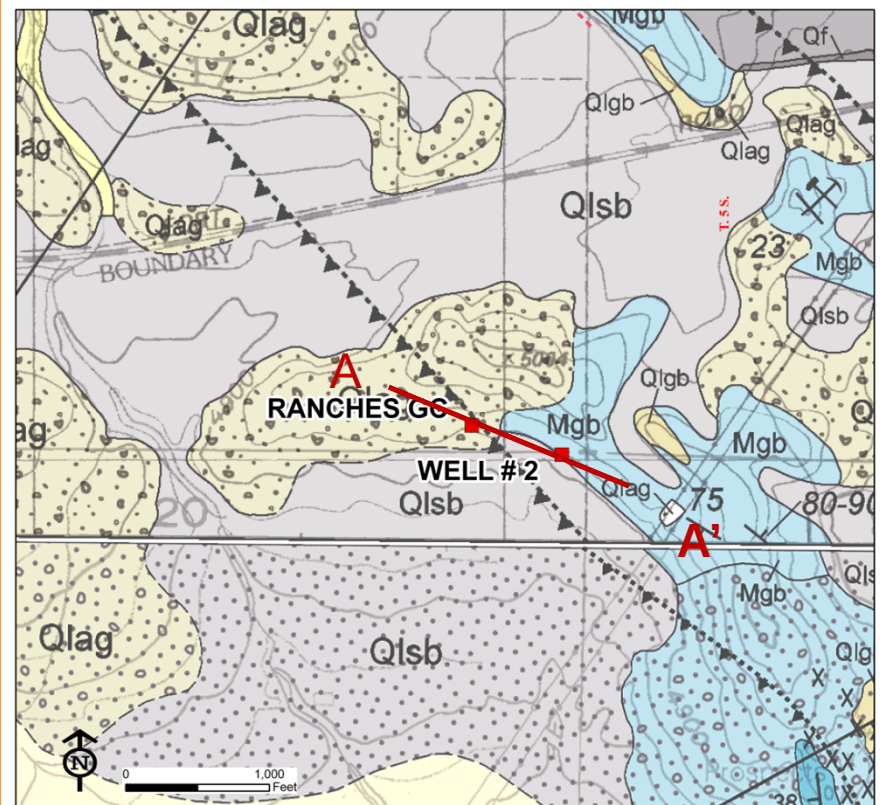
Well #3 test results

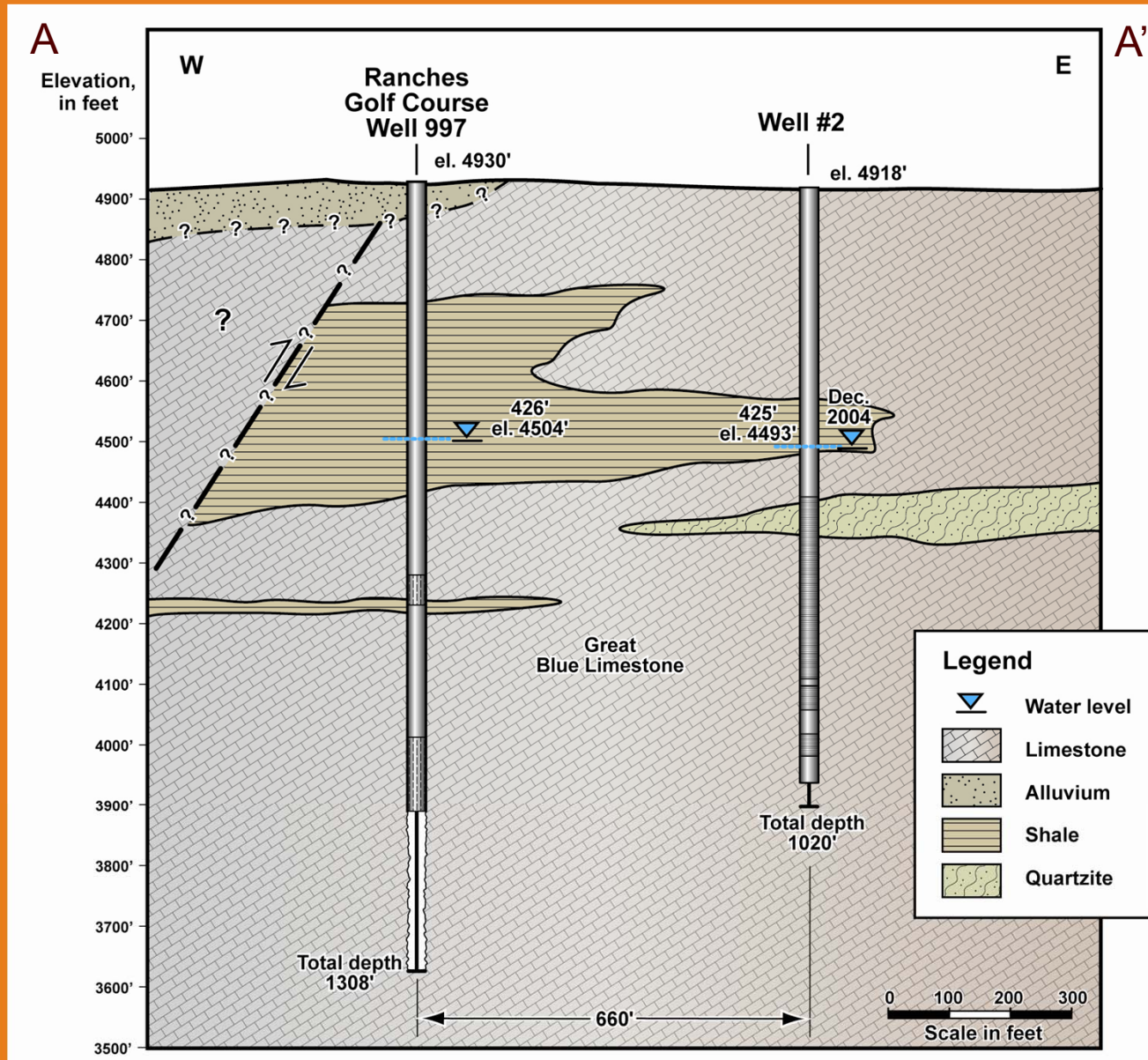
- Dual-porosity system: 3-foot spherical blocks
- Anisotropic with principal direction of T 2+ times greater in the direction parallel to structure

Hydraulic conductivity of fractures (K)	11 to 31 ft/d
Hydraulic conductivity of matrix blocks (K')	$3 \cdot 10^{-7}$ to $1 \cdot 10^{-5}$ ft/d
Specific storage of the fractures (Ss)	$3 \cdot 10^{-7}$ to $3 \cdot 10^{-5}$ ft ⁻¹
Specific storage of the matrix blocks (Ss')	$2 \cdot 10^{-5}$ to $2 \cdot 10^{-3}$ ft ⁻¹
Transmissivity (T)	5800 to 16,000 ft ² /d
Storage capacity or storativity (S)	0.01
Aquifer thickness (b)	530 ft

Aquifer test: Eagle Mountain Well #2

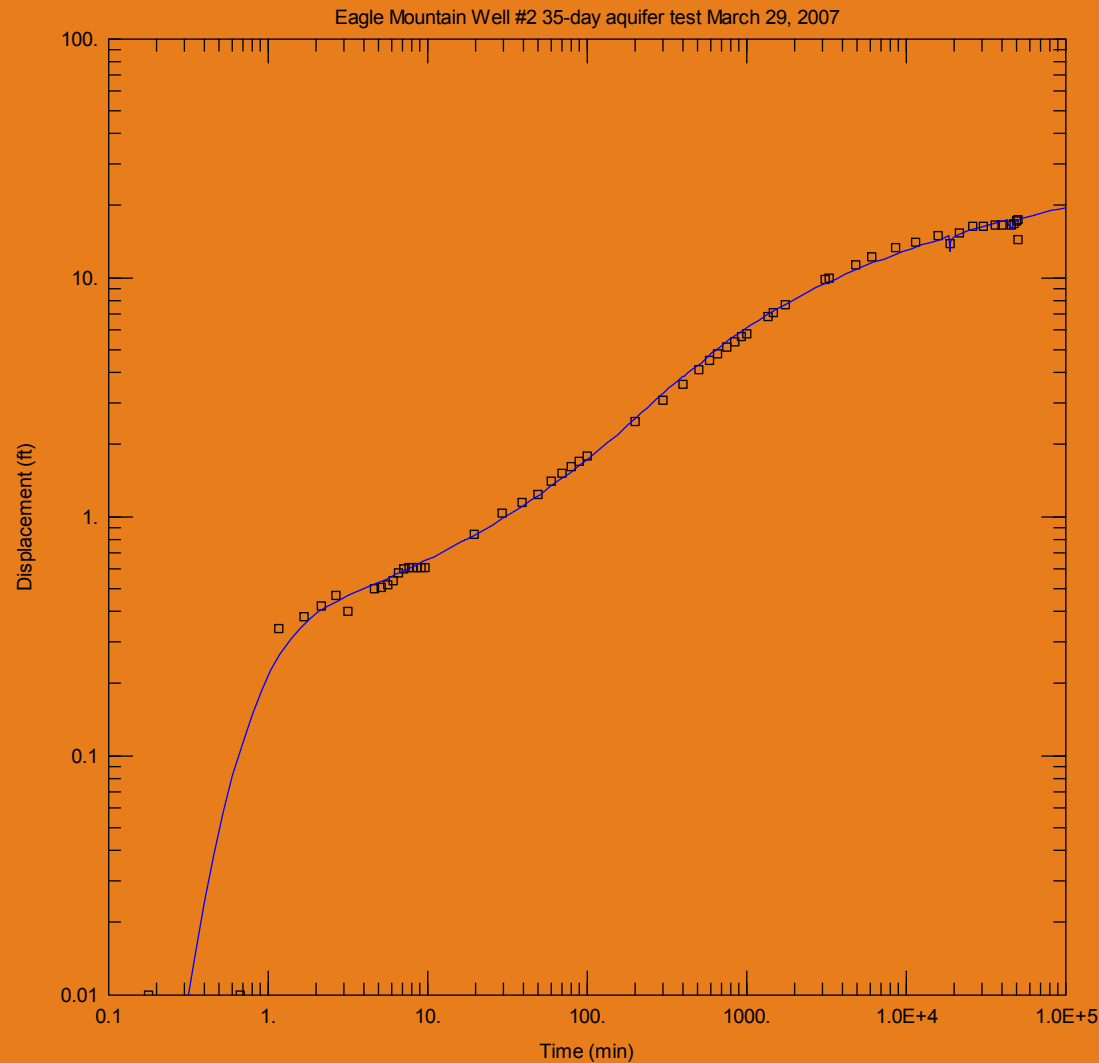
- 35 days, 2500 gpm
- Aquifer thickness ~525 feet
- One observation well 660' distant
- Both wells in Great Blue Limestone







Well #2 test results



Obs. Wells

□ 997

Aquifer Model

Fractured

Solution

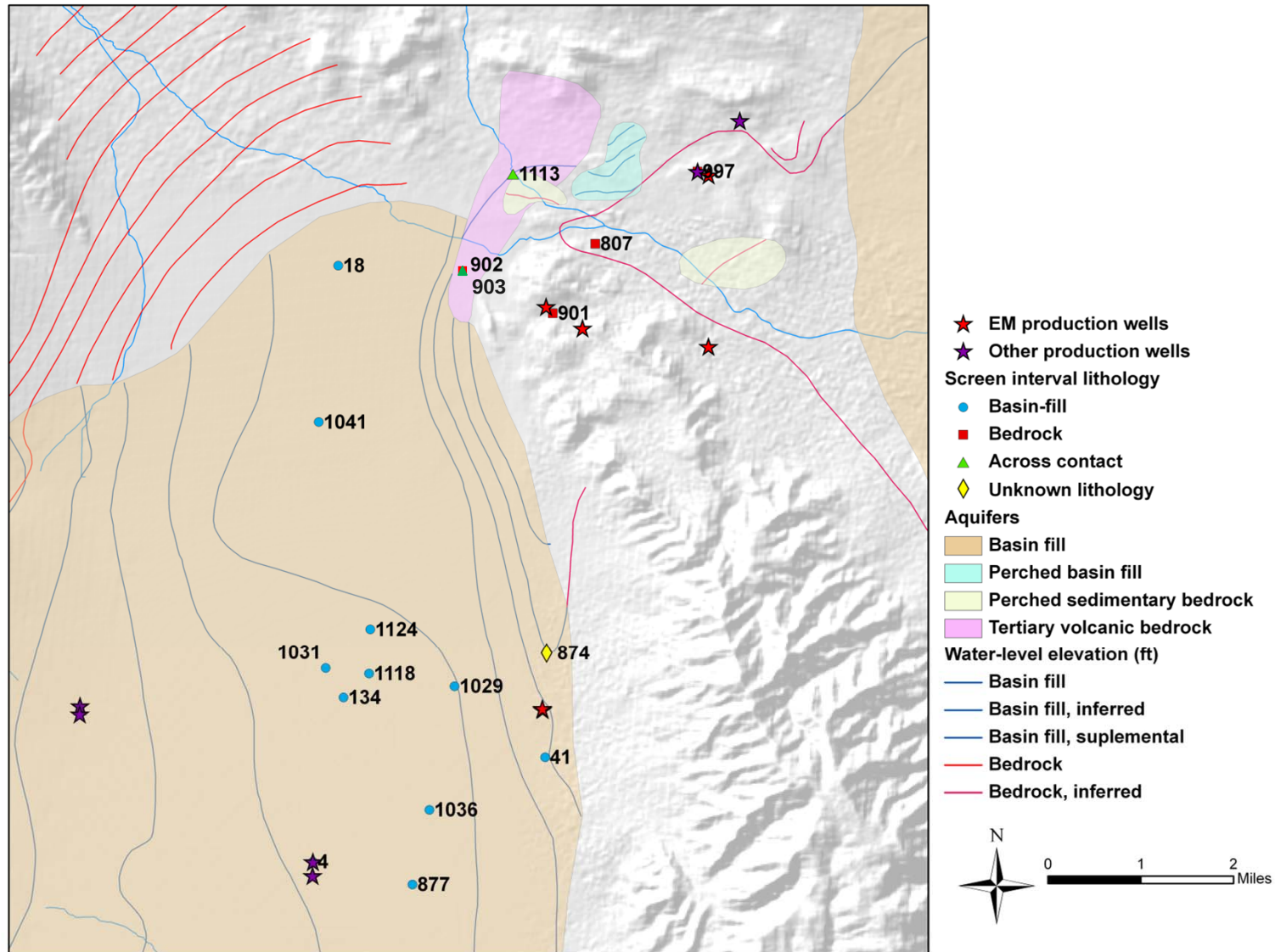
Moench w/spherical blocks

Parameters

$K = 24.01$ ft/day
 $Ss = 1.198E-7$ ft⁻¹
 $K' = 6.36E-5$ ft/day
 $Ss' = 1.29E-5$ ft⁻¹
 $Sw = 10$
 $Sf = 0.2158$

- Max s 660' away = 17.45'
- Good match to fractured solution
- Confined, double porosity system
- K of fractures 10^7 greater than K of blocks
- $T = 12,600$ ft²/d
- $K = 24$ ft/d
- $S = 7 \times 10^{-3}$
- No flow boundary at fault

Water-level monitoring locations





The graph displays the water-level elevation (feet) for four UGS wells (807, MW1, MW2b, and 997) from 3/1/04 to 3/1/11. The y-axis ranges from 4475 to 4515 feet. The x-axis shows dates from 3/1/04 to 3/1/11. Well 997 shows a sharp drop in water level around 3/1/09, labeled 'Pumping water level'.

Date	807 (feet)	MW1 (feet)	MW2b (feet)	997 (feet)
3/1/07	4504	4509	4509	4504
3/1/07.5	4504	4510	4510	4504
3/1/08	4502	4504	4504	4503
3/1/08.5	4498	4498	4499	4493
3/1/09	4499	4500	4500	4503
3/1/09.5	4499	4501	4501	4477
3/1/11	4498	4498	4499	4504



Conclusions

- Geologic controls
 - Disposition of geologic units
 - Tv is a “dam”
 - Perched unconsolidated aquifer OK aquifer
 - Basin fill has no outlet other than to move into Paleozoic bedrock aquifer
 - Structure
 - Folds = anisotropy
 - Faults = impediments and conduits
- Slight aquifer depletion?



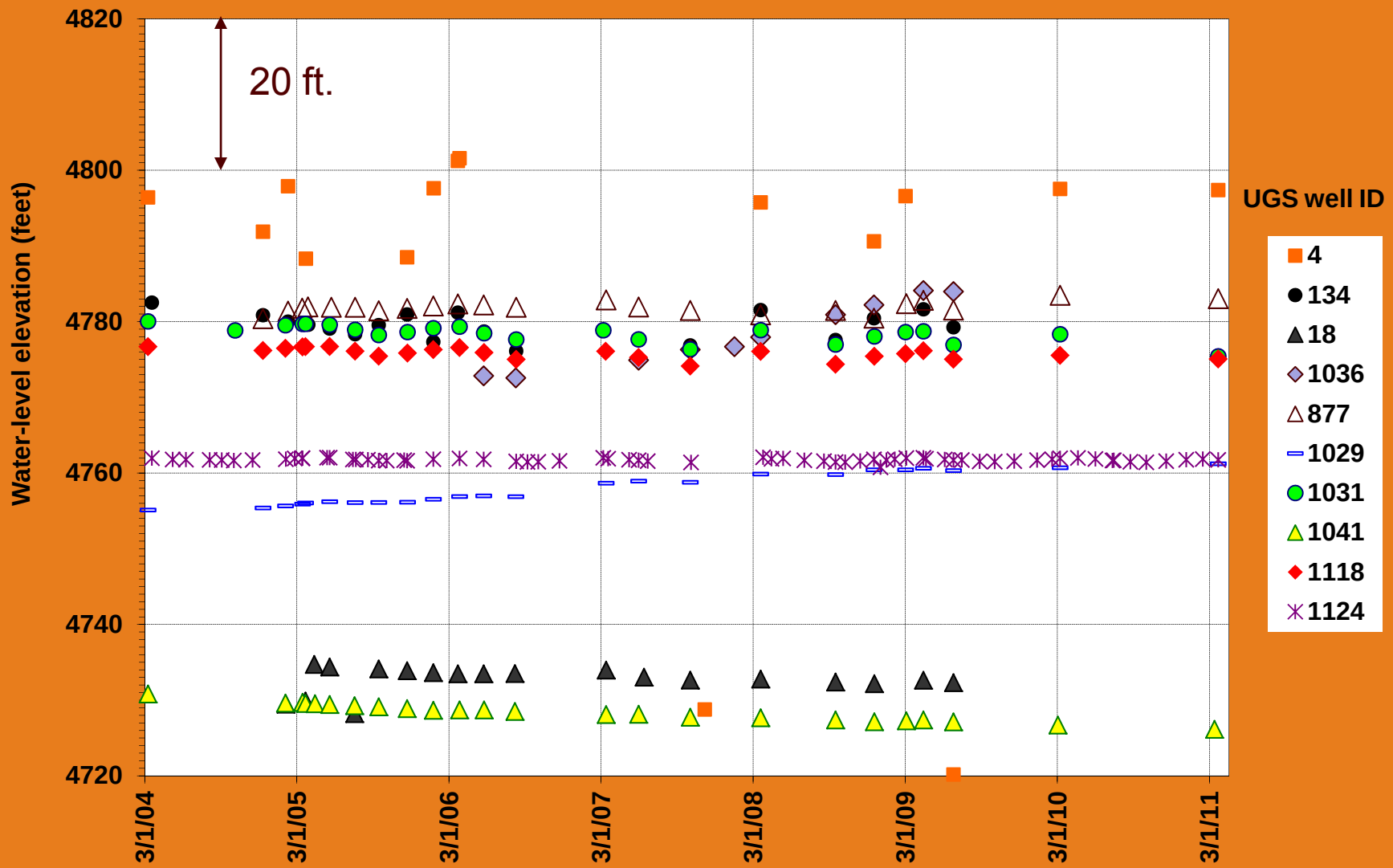
Thank you





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Water-level elevation fluctuation in wells monitored for Eagle Mountain: basin-fill aquifer





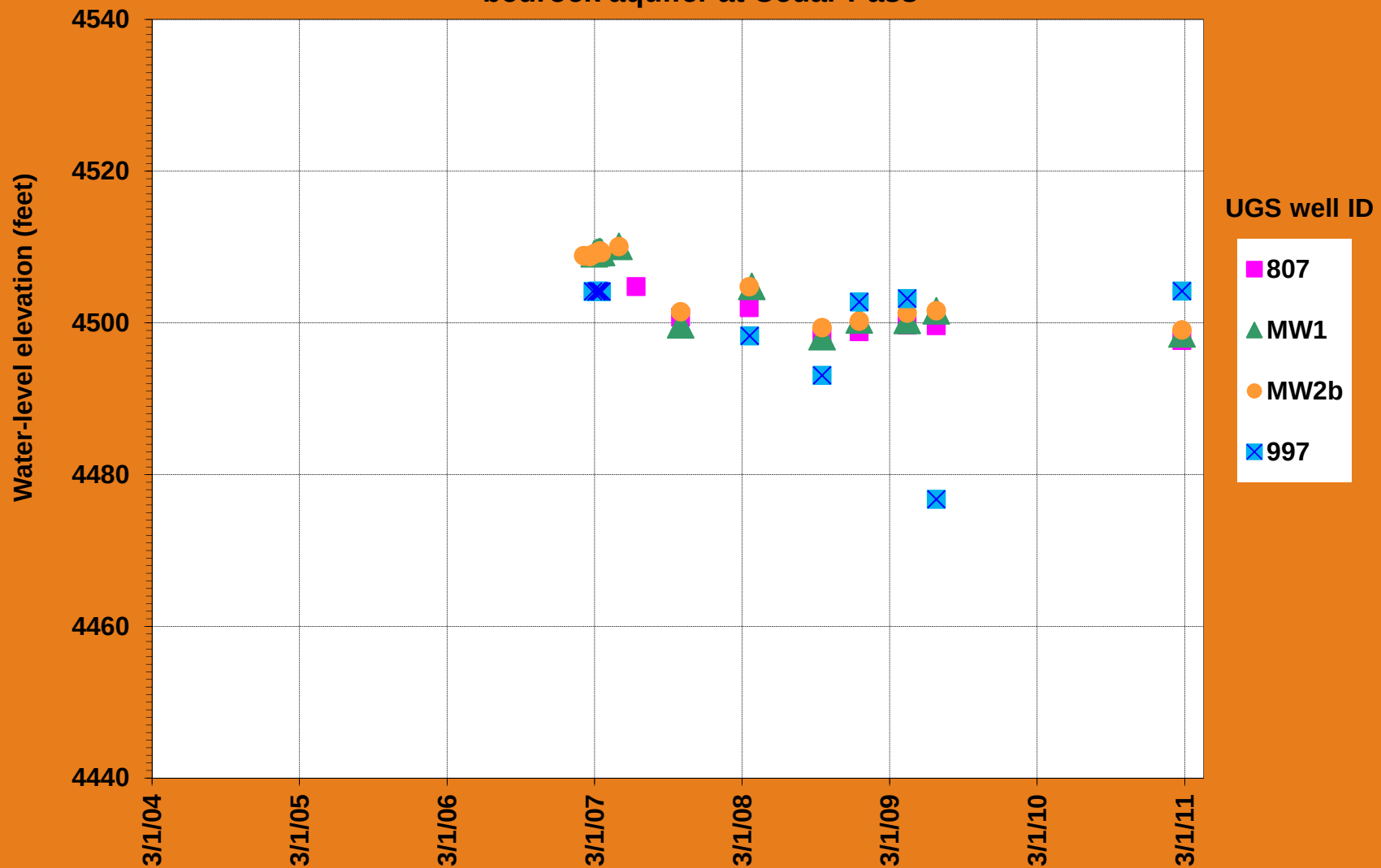
The scatter plot displays water-level elevation in feet on the y-axis (ranging from 4840 to 4940) against dates on the x-axis (from 3/1/04 to 3/1/11). Two data series are shown: UGS well ID 1113 (purple circles) and MW2a (blue squares). Well 1113 maintains a relatively constant elevation of approximately 4900 feet throughout the period. Well MW2a shows a significant drop in elevation starting in early 2007, reaching a low of about 4851 feet, and then gradually recovering to around 4891 feet by early 2011. A blue oval highlights the initial drop, with an arrow pointing to the text 'Recovering from'.

Date	UGS well ID 1113 (feet)	MW2a (feet)
3/1/04	4902	
3/1/05	4901	
3/1/06	4901	
3/1/07	4900	4851
3/1/08	4900	4891
3/1/09	4900	4891
3/1/10	4900	
3/1/11	4897	4891



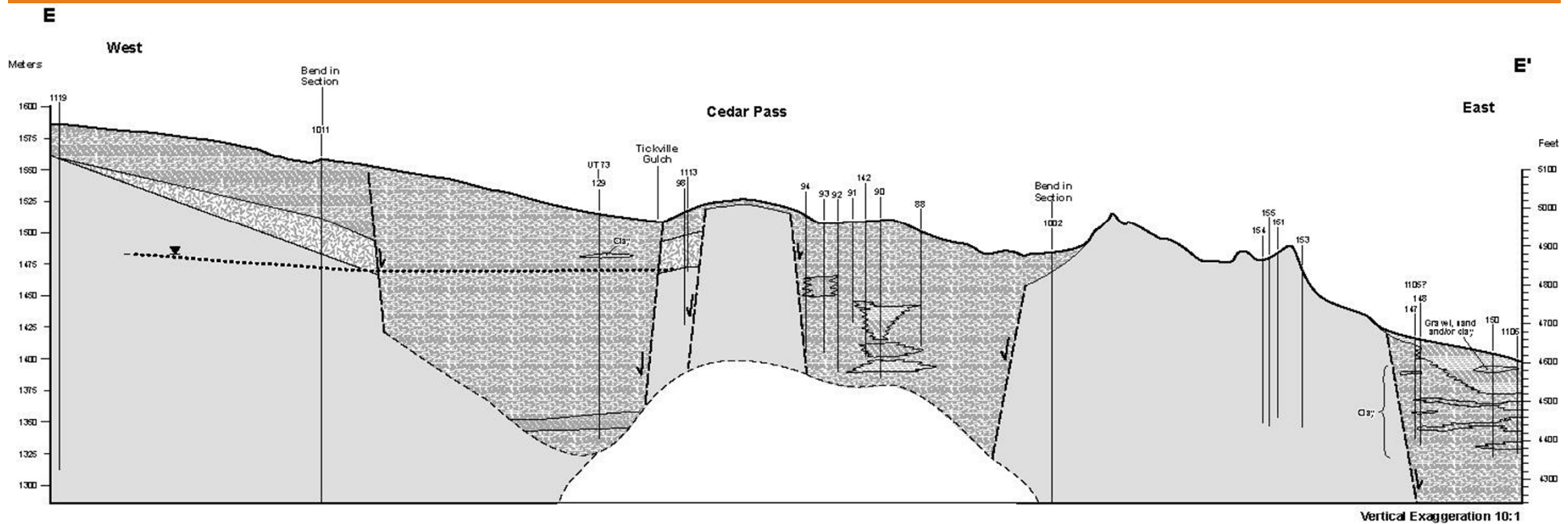
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Water-level elevation fluctuation in wells monitored for Eagle Mountain: bedrock aquifer at Cedar Pass





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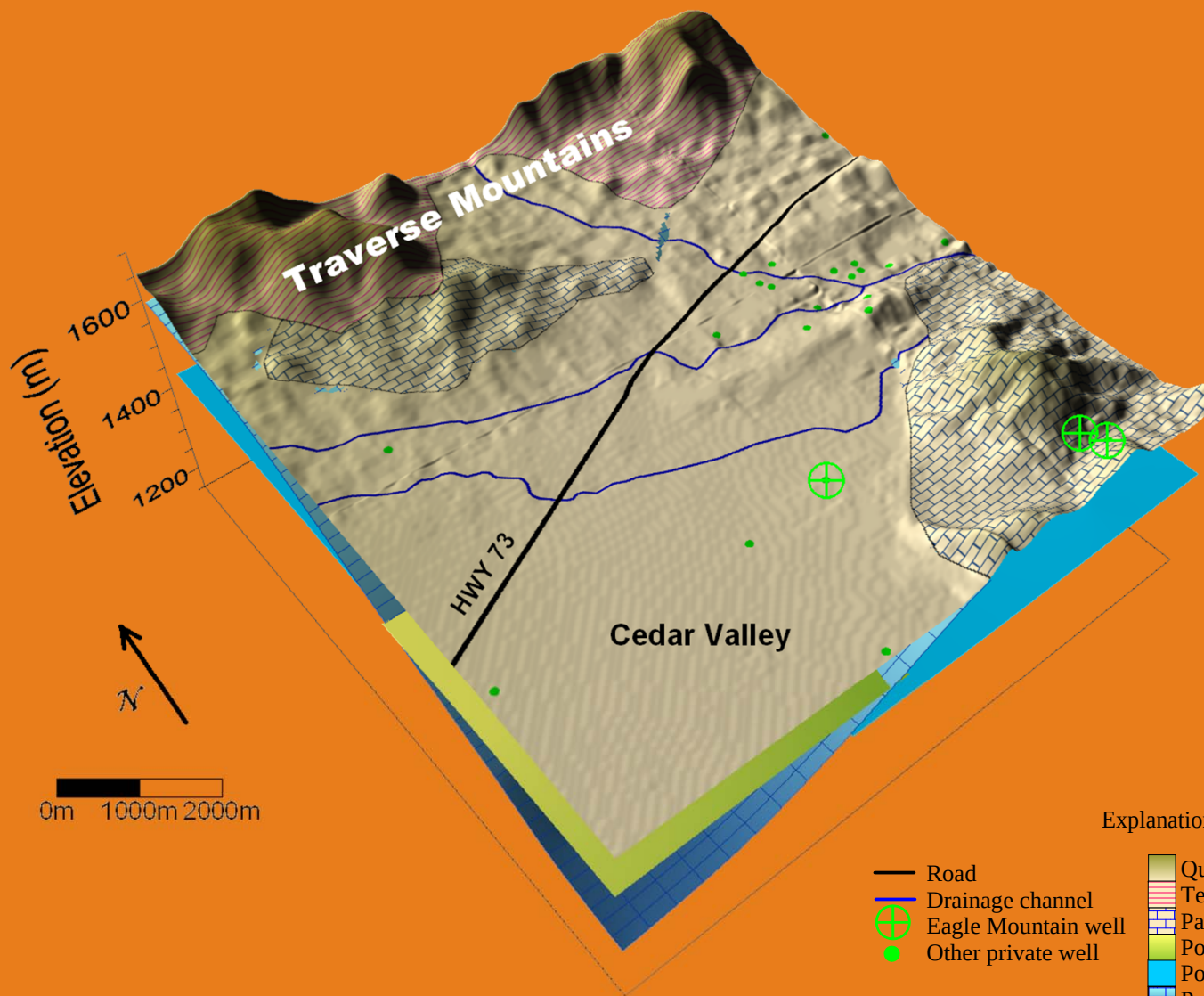


From Hurlow, H.A., 2004, The geology of Cedar Valley, Utah County, Utah, and its relation to ground-water conditions, Utah Geological Survey Special Study, Utah Geological Survey 109.



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Hydrostratigraphy of Cedar Pass area



Explanation

- Road
- Drainage channel
- ⊕ Eagle Mountain well
- Other private well
- Quaternary basin-fill
- Tertiary volcanic bedrock outcrop
- Paleozoic bedrock outcrop
- Potentiometric surface in basin fill
- Potentiometric surface in Paleozoic bedrock
- Paleozoic bedrock surface