

Abstract

The c. 200 km long Chopawamsic fault is the most significant structure within the western Piedmont of northern Virginia, separating the pre-Early Ordovician Potomac terrane accretionary complex from the Middle-Late Ordovician Chopawamsic terrane volcanic arc. It has been hypothesized that the fault could represent the main lapetan suture in the southern Appalachians because the Potomac terrane appears to be affiliated with Laurentia, whereas the affinity of the Chopawamsic terrane is unknown, but possibly Gondwanan. Previous field, geophysical, and regional interpretations suggest that the Chopawamsic fault is an east-dipping, west-vergent thrust that is overlain by two small successor basins to the west of Fredericksburg, VA. In order to assess the significance, kinematics and timing of this prominent structure, we have initiated a multidisciplinary investigation of the fault and surrounding rocks.

In the area south of Lake Anna, limited observations of winged sigma feldspar grains indicate a sinistral component of shear along the Chopawamsic fault. Nearby, the Ellisville granodiorite intrudes across the fault, providing a timing constraint on fault motion. New TIMS U-Pb zircon ages indicate that the latest movement on the fault was pre- c. 437 Ma.

Field studies to date confirm the existence of a successor basin atop the fault in the area of Storck, VA; we intend to use detrital zircon analysis of metasediments in the basin in order to further constrain the timing of fault motion. In the Wilderness, VA area, our mapping indicates that the other successor basin previously mapped to overlie the fault, does not exist and that rocks previously mapped as the basin are most likely metasedimentary rocks within the Chopawamsic ter-

Our 'in progress' investigation will involve additional field and laboratory research along the Chopawamsic fault. We plan to assess the cratonic source of the Chopawamsic terrane with the use of Nd and Pb isotopic analyses as well as detrital zircon populations from each side of the Chopawamsic fault. The detrital zircon study should also provide supplementary constraints on the timing and nature of fault motion. High precision U-Pb TIMS data from previously undated intrusive bodies will also augment our understanding of the early Paleozoic evolution of the western Piedmont.



NEW OBSERVATIONS ON THE CHOPAWAMSIC FAULT, AN EARLY PALEOZOIC TERRANE BOUNDARY IN THE WESTERN PIEDMONT OF VIRGINIA HUGHES, Stephen¹, (kshughes@ncsu.edu); TERBLANCHE, Alet¹; NANCE, Dillon¹; HIBBARD, James¹; and MILLER, Brent² (1) Department of Marine, Earth, and Atmospheric Sciences, North Carolina State University, (2) Department of Geology and Geophysics, Texas A&M University

Focus Area (1)

This area was chosen to investigate two proposed successor basins comprised of undivided metasedimentary rocks which may overlie the Chopawamsic fault. These basins could be related to the larger Quantico and Arvonia formations. We informally refer to the feature in the southern portion of the focus area as the 'Wilderness basin' and the northern feature as the 'Storck basin.'

Storck Area



Through preliminary mapping in the Storck area, we have identified various metasedimentary rocks, which may be part of a successor basin thought to overlie the Chopawamsic fault. If this feature is a successor basin, it would be an important factor in helping to constrain timing of the fault. A sample of micaceous quartzite from this area (Fig. 3) has been collected and will be dated using detrital zircon



analysis.

Observations of the heterogeneously deformed Richland Run pluton in the Storck area reveal increasing deformation from E to W approaching the Storck Basin (Figs. 4-6). This deformation array hints at the presence of the Chopawamsic fault beneath the successor basin.

Wilderness Area

The Wilderness area contains rocks that are identical to the known sedimentary lenses within the Chopawamsic Fm. This is evident in outcrop, thin section and XRD (X-ray Diffraction) analysis. The outcrop similarities are displayed in Figures 8 & 9, where Fig. 8 represents the Chopawamsic Fm. and Fig. 9 shows outcrop in the previously mapped Wilderness basin. XRD analysis has aided in distinguishing between phyllosilicates found within these different areas and we conclude that there are only two major different rock types: Mine Run Complex and Chopawamsic Fm.

There is no evidence that a synclinal basin, which was previously mapped as conformably overlying the Chopawamsic fault, exists. While the Storck basin may remain suitable for placing time constraints on the Chopawamsic fault, this area does not present such an opportunity.

Conclusions

Our investigation along the Chopawamsic fault in the western Piedmont of Virginia remains in progress. Our observations have led us to these preliminary conclusions:

the Wilderness basin.

(1) The Storck successor basin remains a valid possible timing constraint over the Chopawamsic fault and further exploration will provide more information

(2) The previously mapped Wilderness successor basin does not exist and cannot aid in providing timing constraints to the fault

(3) The stitching Ellisville pluton represents a petrographically and geochemically coherent pluton that intrudes across the Chopawamsic fault

(4) The Ellisville pluton stitched the Chopawamsic fault and intruded both the Potomac and Chopawamsic terranes at c. 440 Ma

(5) A sinistral component of shear exists along the Chopawamsic fault and we interpret it to be a record of original motion, along with an inferred thrust component



igure 9 a & b: Rocks previously mapped

Fia. 6

Focus Area (2)

This area was chosen to investigate the possible stitching relationship between the Ellisville pluton and the Chopawamsic fault. While the main body of the pluton has been interpreted to intrude the Mine Run Complex (Pavlides et al., 1994), we have directly observed the Ellisville pluton tail in intrusive contact with the Chopawamsic Formation (Fig. 10). Our detailed mapping, and petrographic and geochemical analyses (Hughes, 2011; Hughes & Hibbard, 2011) lead us to conclude that the Ellisville pluton can be confidently used to constrain the timing of latest motion upon the Chopawamsic fault.



Although generally considered a thrust fault (e.g. Pavlides, 1989; 1990; 1995; Mixon et al., 2000; 2005), there has been no published kinematic information pertaining to the Chopawamsic fault. In Focus Area 2, we have discovered σ -type, winged feldspar clasts which indicate a sinistral component of shear along the Chopawamsic fault (Fig. 12). We interpret these indicators, found just south of the South Anna River, to be indicative of original motion along the fault due to their proximity to the stitching Ellisville pluton -- a feature which records no reactivation along the fault. Steeply dipping clast and mineral lineations have also been observed along the trace of the Chopawamsic fault. Collectively, these kinematic observations suggest the fault formed as Northeast. a result of transpressional deformation.

Future Work

To test the first order hypothesis that the Chopawamsic fault may represent the main lapetan suture in the southern Appalachians, we intend to determine the cratonic source for the Chopawamsic terrane. This goal will be accomplished with the use of isotopic fingerprinting and detrital zircon analysis. Comparison with data from the Laurentian Potomac terrane will aid in this endeavor. The dating of volcanics within the Chopawamsic formation and previously undated intrusive bodies will also help to clarify the early Paleozoic evolution of the western Piedmont of Virginia.

References







Our new geochronological data affirm an earlier evaluation of the Ellisville pluton as a texturally composite body (Pavlides and Cranford, 1982). While there are no compositional differences within the pluton, we interpret the composite nature to reflect only a main pulse of magmatism throughout the pluton at c. 444 Ma and a later, subsidiary phase at c. 437 Ma. These new data indicate that movement along the fault occurred sometime prior to stitching at c. 444 Ma. Data from the Chopawamsic formation (Coler et al., 2000; Horton et al., 2010) indicate that the fault initiated sometime after 471-453 Ma.



Fig. 12: Photomicrographs taken at 2x. Field c view is 9.5 mm. From left to right is towards



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Chopawamsic fault and are presently undated

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