

Geomorphic paleogeodesy and intraplate deformation associated with the Mineral, VA earthquake and surrounding Central Virginia Seismic Zone (CVSZ) Frank J. Pazzaglia and Helen Malenda<sup>1</sup> Earth and Environmental Sciences, Lehigh University fjp3@lehigh.edu

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In the past two decades, geodesy and geochronology have endeavored to quantify crustal strain and paleoseismology in the central and eastern U.S (CEUS). We have learned that:

Seismic zones have little to no measurable crustal strains when using GPS geodesy.

Seismic zones have measurable crustal strains when using geomorphic and stratigraphic markers.

Seismic zones have clear evidence of past earthquakes.

We know virtually nothing about the potential for earthquakes in between seismic zones.

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Wang et al., 2012, Nature

### No GPS geodetically measurable crustal strains at New Madrid



Smalley et al., 2005, Nature



Calais et al., 2005, Nature Calais and Stein, 2009, Science Craig and Calais, 2014, JGR-SE Boyd et al., 2015, JGR-SE

#### Systems-level model for CEUS intraplate earthquakes



Kenner and Segall, 2000, Science

Crustal strains by Bill Holt, unpublished data





Ν 20 KM Run fault 0 38.25°N OI **€**Zm NK NIT N SOg **€**Zm Louisa QN SOg 38.00°N QN A CZm Mineral earthquake PzYm Ashland Fault SOa 37.75°N Pzgp Pzg N South Cross-section courtesy of QN David Spears and Amy Gilmer Oc Richmond Division of Geology and Yp 37.50°N **Mineral Resources PzYm** Yp Virginia Department of Mines, QN Pzgp Minerals, and Energy 1 78.25°W | 78.00°W 77.75°W 77.50°W Interpretive Geologic Cross-section from the Eastern Blue Ridge to the Eastern Piedmont, Virginia Including the 23 August 2011 Mw 5.8 Louisa County Earthquake Cross-section based on published and unpublished geologic mapping and the I-64 seismic line. No vertical exaggeration. Mountain Run Long Branch Shores-Brooknea Fault Zone Chopa Faul Hylas Fault Zor Chopawamsic Spotsylvania Α Fault (?) Fault (?) Fault Zone €c: O€ml\ €cgs Oct Ybrb ZI O€ml\ ted Oceanic Crust (' Och Mw 5.8 23 Aug 2011 Basement Duplex (?)





1.4e+9 (a) 200 ~300 m drainage divide Gordonsville Elevation (m) 051 1.2e+9 1.0e+9 Gordonsville Two large 8.0e+8 (2) knickpoints Bon Air 100 Area Area 8+90.0 oindexter Choptank related to Fm ~12 Ma South Anna anceyville Many smaller Orchid Rockville base level 4.0e+8 knickpoints that Yorktown 50 fall at the Fm~3 Ma "decorate" the Gilmans Bridge 2.0e+8 1-95 Fall Zone middle reaches 0.0 0 80 90 100 110 120 of the stream, 70 160 0 10 20 50 60 140 150 30 40 130 distance (m \*10<sup>3</sup>) perhaps related 0.1 to rock-type (C) ks = 2.438 = 274 5000000 (b) 200 slope theta = -0.68 or intraplate  $r^2 = 0.88$ seismicity 0.01 180 4000000 1000 10000 Area (m<sup>2</sup> \*10<sup>3</sup>) 30000000°E elevation (m) 160 Area 140 20000000 120 10000000 100 0 150 152 154 146 148 156 158 distance (m \* 10<sup>3</sup>)













Helen Malenda, EDMAP, G13AC00115

69,200 +/- 16000

~ 2 m

a

colluvium 2

colluvium 1

overbank

channel alluvium

strath

saprolite

(b)



### Horseshoe Farm

# Approximate surface projection of fault that ruptured in 2011









Figure 11. (A) Map of permanent vertical displacement expected at Earth's surface from the Mineral earthquake for a focus depth of 7 km at the center of the rupture. Calculation is based on empirical relations for same slip and dimensions (Wells and Coppersmith, 1994) and a Coulomb elastic dislocation model. Benchmark locations are from the National Geodetic Survey. Geographic Coordinate System: World Geodetic System (WGS 1984). Projection: Datum WGS 1984. VE—vertical exaggeration; USGS—U.S. Geological Survey. (B) Rupture plane and corresponding displacement model for slip generated by an  $M_w$  5.8 earthquake with a depth of 7 km (preferred depth). Maximum upward displacement at the surface along transect A-A' crossing the rupture is ~7 cm. Maximum downward displacement is ~1 cm. (C) Profile of vertical displacement at the surface for a range of depths. Walsh et al., 2015, GSA SP 509.



(1) Hanging wall incision~34 m since mid Pleist.(~115 m/m.y.)

assuming that fault slip is driving ~8-16 m of rock uplift in ~400 ka,

and Mineral earthquake
-scale events generate
~7 cm of co-seismic
fold growth
(Walsh et al, 2015),

That is **~114-228** earthquakes with a R.I. of **~3,500 - 1750 yrs.** since the middle Pleist.

55(2) Hanging wall incision
~26 m since mid Pleist.
(~65 m/m.y)

Great Falls of the Potomac....large base level fall and river incision in the VA-MD Piedmont that is currently located between recognized (dense) clusters of seismicity in CVSZ and RLSZ.....



## Conclusions

Direct evidence for local crustal deformation of the Appalachian landscape is evident in the warped terraces of the South Anna River that flows above the epicenter of the 23 August, 2011 Mineral earthquake.

GPS geodesy may struggle to measure crustal strains in the seismic zones, but maybe this is the wrong place to capture the elastic phase of the viscoelastic oscillator. Geomorphic geodesy can measure the accumulated strain of permanent co-seismic deformation, from many earthquakes.

What we do not know, can hurt us. Nearly all intraplate earthquake investigations have been focused in seismic zones; we should also be looking in the aseismic gaps for evidence of crustal strains in CEUS; geomorphic paleogeodesy is one tool that can help identify crustal strains in places that are currently "aseismic".



6

0

(cm) 0.2 0.5 ppe 0.7 0.85 unit CO VA 03B OSL owel ed 67 ka (saturated) 1.05 Alluvial (?) gravel VA 03A IRSL 1.3 212.3+/-29.2 ka saprolite VA 02 IRSL 1.5 391.8+/-44.4 ka

meters

2.5







### Eastern U.S. Dynamic topography (Moucha et al., 2008, EPSL)





Α













Contoured second invariant of strain rate for the lithosphere in parts of ENAM, together with depth-integrated deviatoric stresses, obtained from a global model and using a craton viscosity and old ocean (Müller et al., 2008, 1997) lithosphere viscosity of 1x10<sup>24</sup> Pa-s. Bold vectors are principal axes of compression and open vectors are principal axes of tension. The craton distribution was obtained from the tomography model in Simmons et al. (2006).

