

Use of Zircon U-Pb Ages And Bulk Trace Element and Nd-Hf Isotope Compositions to Determine Soil Provenance in a Limestone Terrane, Middle TN, USA

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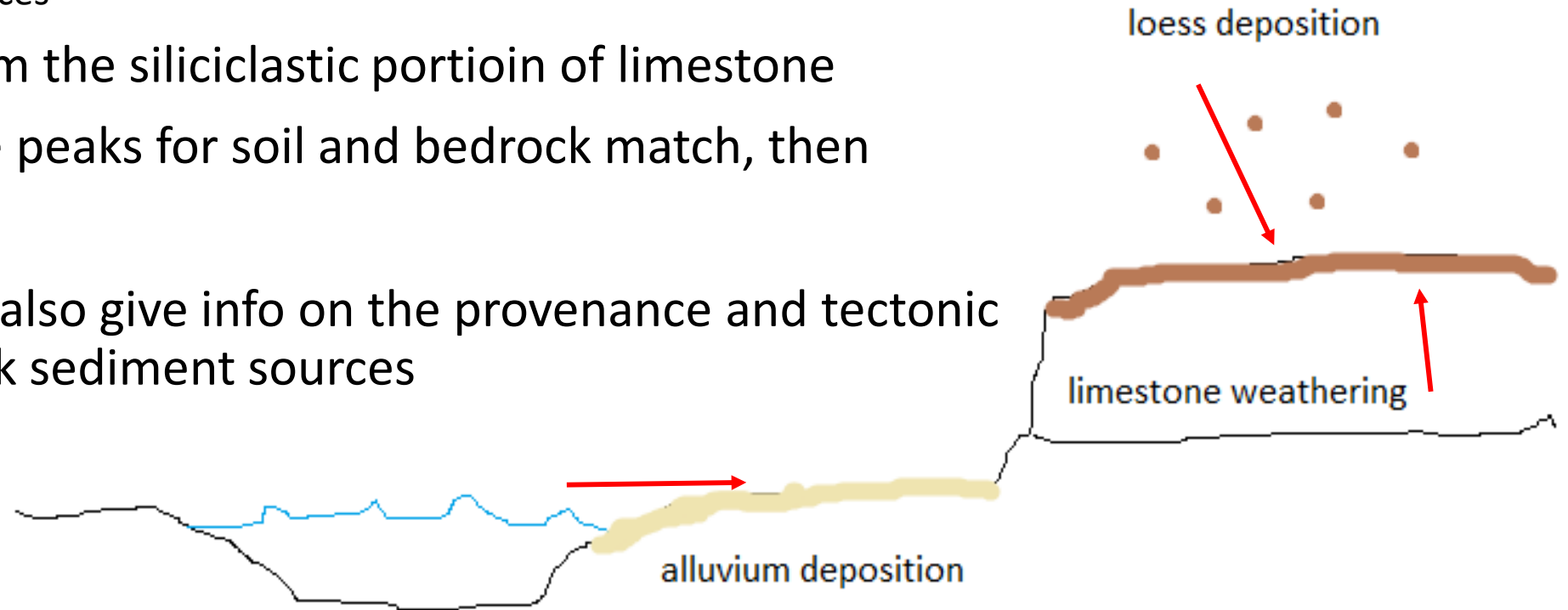
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How did soil form in middle TN?

- In-situ chemical weathering of limestone?
- Input of exotic material such as alluvium or loess?
- Identify the source (provenance) of soil-forming material by studying samples of
 - undisturbed soil (high, flat, non-agricultural areas)
 - potential soil sources
- Obtain zircon from the siliciclastic portion of limestone
- If zircon U-Pb age peaks for soil and bedrock match, then soil formed in-situ
- Zircon U-Pb ages also give info on the provenance and tectonic setting of bedrock sediment sources



Why zircon U-Pb dating?

- Elemental and isotopic composition and mineralogy not reliable provenance indicators
 - Transport of loess and river alluvium can change composition and mineralogy of deposit
 - Chemical weathering during soil formation changes composition and mineralogy
- Zircon age spectra less likely to be affected by transport or chemical weathering
- Only two previous studies of soil provenance using zircon U-Pb dating; the first was:

Brimhall, G.H., Lewis, C.J., Compston, W., Williams, I.S., Reinfrank, R.F., 1993. Darwinian zircons as provenance tracers of dust-size exotic components in laterites: mass balance and SHRIMP ion microprobe results, in Soil Micromorphology: Studies in Management and Genesis. Elsevier, pp. 65–81. [http://dx.doi.org/10.1016/S0166-2481\(08\)70398-2](http://dx.doi.org/10.1016/S0166-2481(08)70398-2)

Middle Tennessee Geology

Site 1: Fort Payne
Formation



Site 2

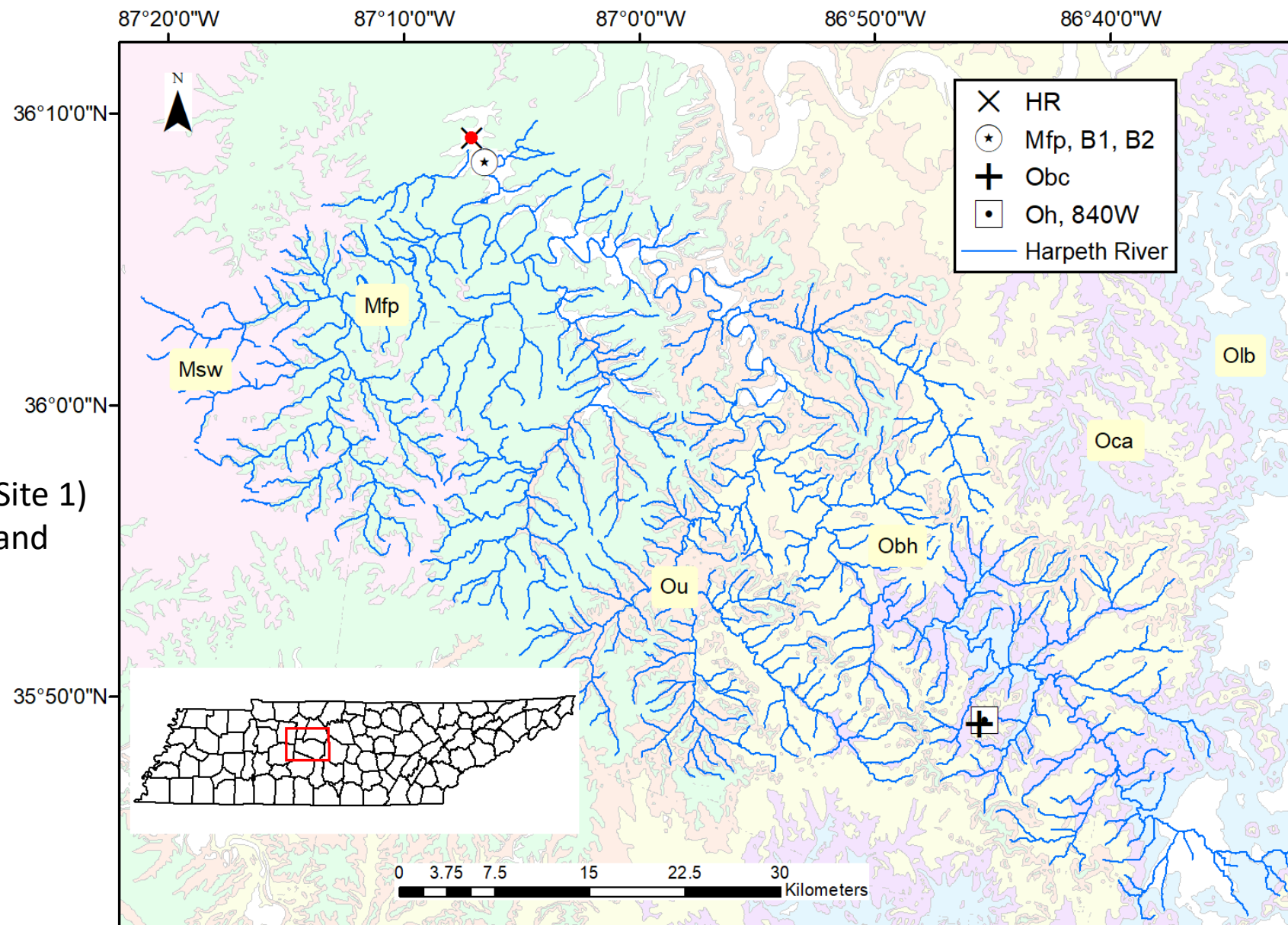
Bigby-Cannon
Limestone
Hermitage
Formation



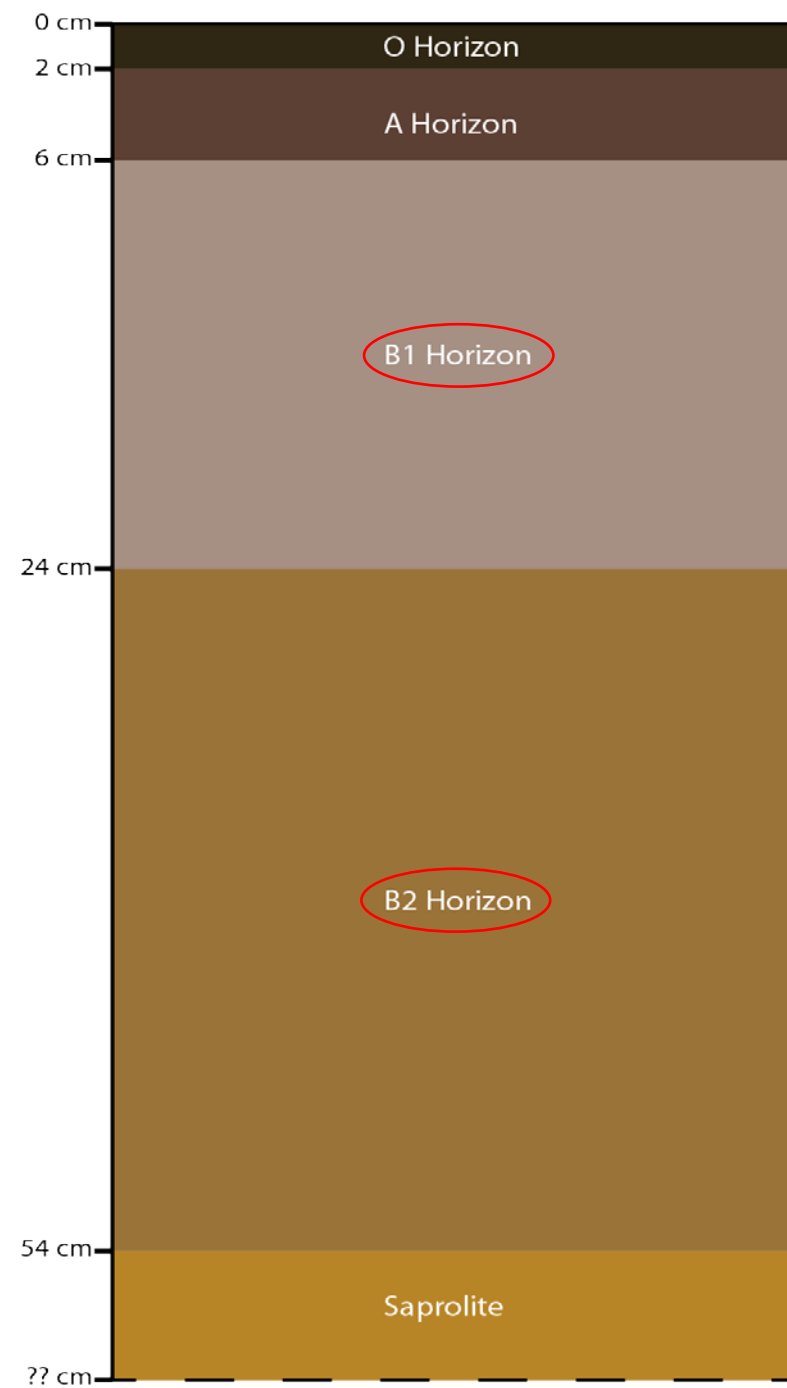
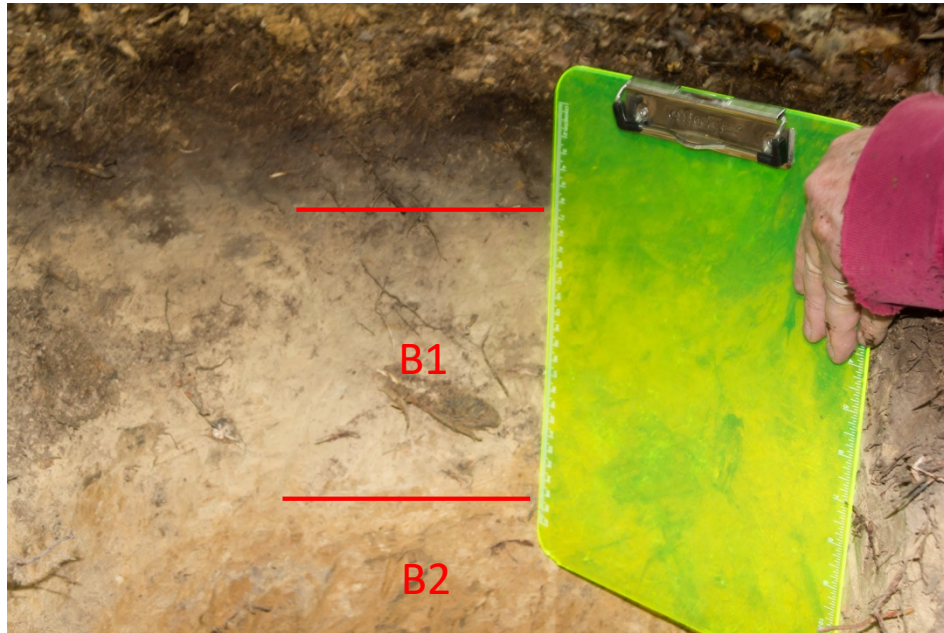
SYSTEM	GROUP	FORMATION	ROCK STRATA	AVERAGE THICKNESS ft.	RANGE OF THICKNESS ft.
MISSISSIPPIAN		STE. GENEVIEVE (MONTEAGLE)		250	180-350
		ST. LOUIS		180	100-280
		WARSAW		100	40-150
		FT. PAYNE		250	200-400
DEVONIAN		CHATTANOOGA		20	10-70
		PEGRAM		17	0-30
		CAMDEN		95	0-220
		FLAT GAP		20	0-55
		ROSS		45	0-110
SILURIAN	WAYNE	DECATUR		VARIABLE	0-250
		BROWNSPORT			
ORDOVICIAN		BRASSFIELD			
		SEQUATCHIE		55	0-275
	MAYSVILLE	LEIPERS		70	0-160
	EDEN	INMAN		50	0-70
	NASHVILLE	CATHEYS		130	10-250
		BIGBY CANNON		80	50-100
		HERMITAGE		120	70-180
	STONES RIVER	CARTERS		60	37-93
		LEBANON		92	74-120
		RIDLEY		110	110-115
		PIERCE		27	23-28
		MURFREESBORO		70*	200-400

Sample Locations

- Mfp = Mississippian Fort Payne (Site 1)
- Obh = Ordovician Bigby Cannon and Hermitage limestones (site 2)



Field Site 1: Harpeth River Terrace: Ultisol



Field Site 2: Highway Outcrop

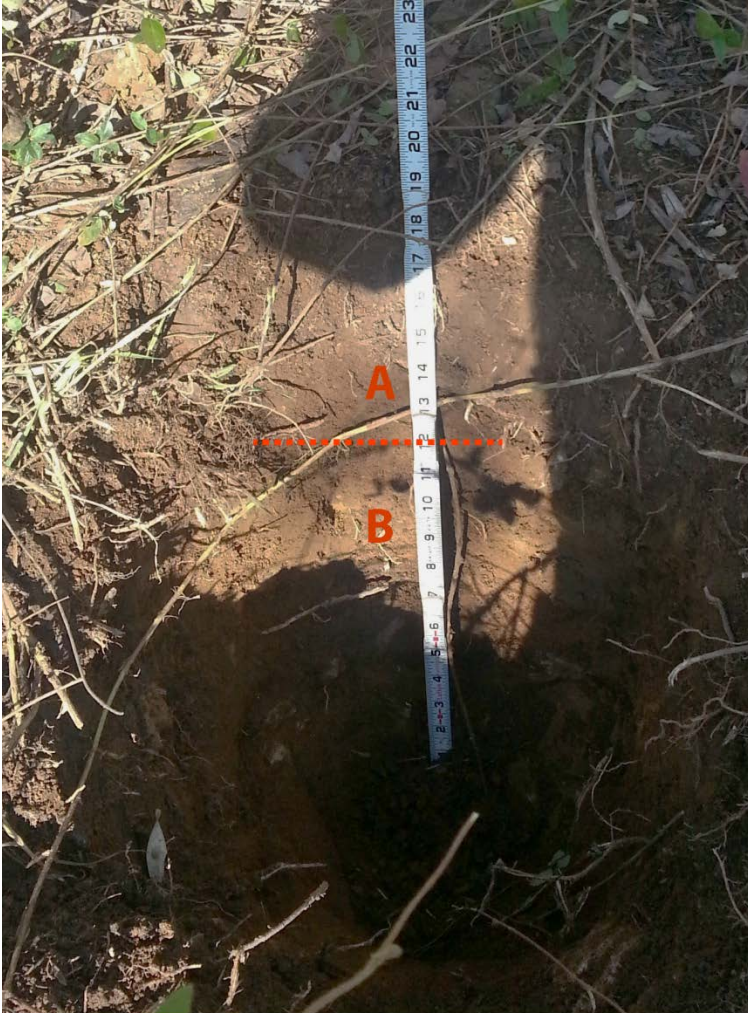
Sampling

- Alfisol soils atop Hermitage Fm. (Oh) Ordovician sandy and argillaceous limestone bedrock



Field Site 2:

Sampling of alfisol



Field Site 2: Obc sample

- Bigby-Cannon Limestone (Obc)
Ordovician phosphatic calcarenite/limestone
- Located 0.15km away from the Hermitage-Alfisol sample site



Xiaomei Wang and Jenna Nam

Loess
outcrop in
Meeman
Shelby Park,
Memphis,
TN



Bulk Soil Properties

Site	Sample	Munsell Color	ρ (g/cm ³)	Mean grain size (μ m)	Soil Texture (USDA)
1	B1	10YR 6/3	1.5 ± 0.2	43	Silt
1	B2	10YR 5/6	1.5 ± 0.2	19.9	Silt
2	W840	7.5YR 5/4	1.8 ± 0.2	29.3	Silt Loam

Bulk Properties

Sample	% Insol. Residue	€ _{Zr, w}	CIA %	% OC
B1	NA	0.81	73	0.69
B2	NA	0.76	84	0.63
W840	NA	0.68	76	1.81
Mfp	99%	NA	78	
Oh	28%	NA	70	
Obc	15%		60	
PL	82%			3.92
RS	99%			0.25
HR	99%			1.02

Major element compositions

Original name	Loveland loess	Roxana silt	Peoria loess	Harpeth R.	Mfp	B1	B2	Oh	Obc	W840
SiO ₂	88.92	77.26	69.46	81.87	87.33	86.14	78.12	24.2	5.67	67.32
TiO ₂	0.32	0.87	0.71	0.99	0.33	1.03	0.99	0.2	0.04	0.64
Al ₂ O ₃	5.29	10.05	9.10	7.16	5.13	5.14	8.76	4.02	1.04	6.34
TFe ₂ O ₃	1.75	3.91	3.21	3.42	2.58	2.36	4.39	1.62	0.49	3.21
MnO	<0.01	0.06	0.08	0.04	<0.01	0.02	0.02	0.16	0.05	0.15
MgO	0.21	0.68	2.83	0.37	0.46	0.24	0.43	1.19	2.68	0.47
CaO	0.48	0.49	3.89	0.39	0.19	0.1	0.09	36.57	48.2	7.7
Na ₂ O	0.04	0.48	1.33	0.32	0.09	0.37	0.27	0.13	0.11	0.25
K ₂ O	0.19	1.66	2.07	1.1	1.05	1.03	0.95	1.19	0.3	1.09
P ₂ O ₅	0.03	0.06	0.11	0.24	0.05	0.04	0.08	1.69	2.81	5.96
LOI	2.61	4.30	7.17	4.29	2.97	3.31	5.76	29.28	38.31	6.75
TOTAL	99.84	99.81	99.95	100.19	100.18	99.78	99.86	100.25	99.7	99.88

Powder XRD Results

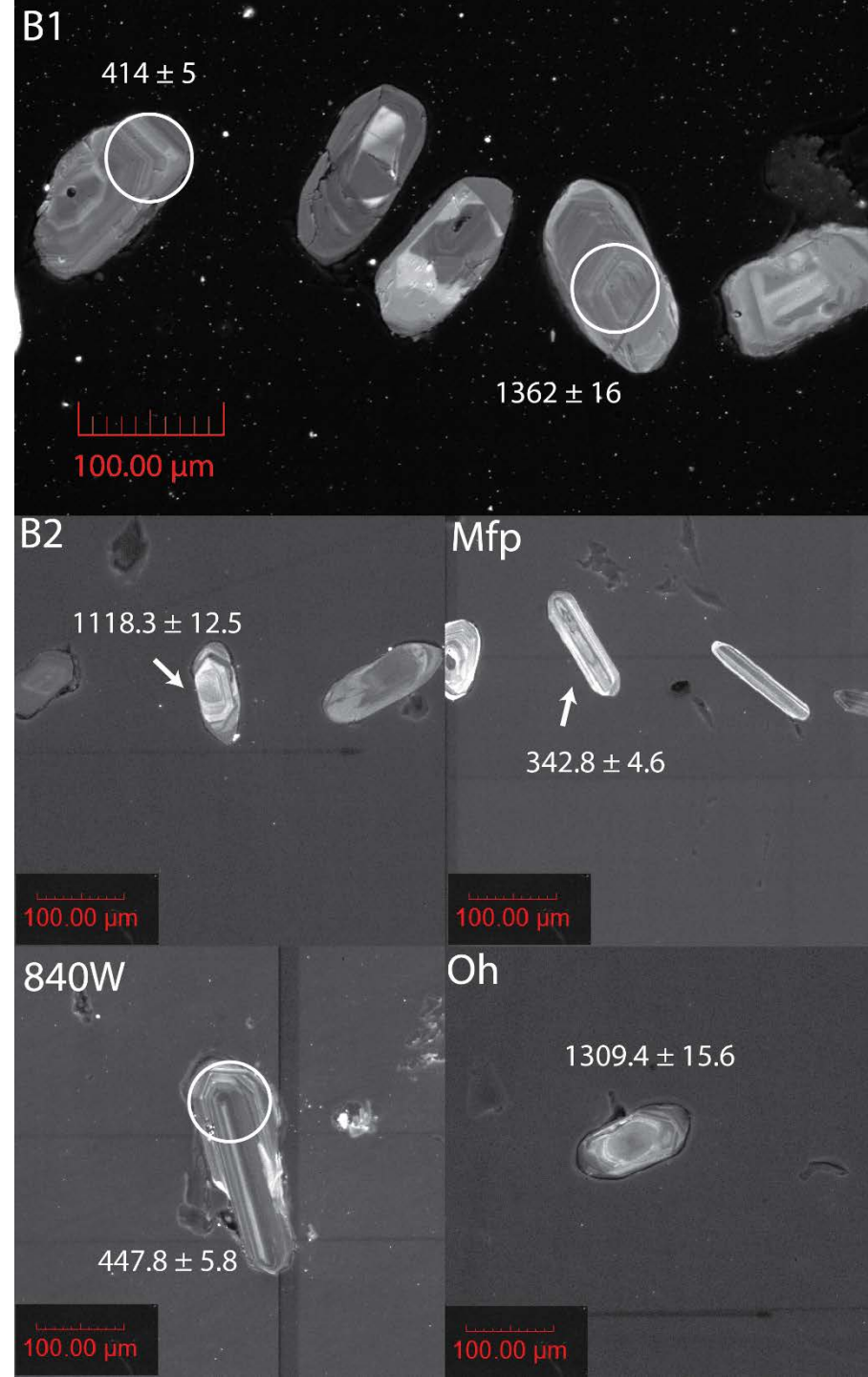
- W840 soil has same minerals as Oh bedrock
- B1 and B2 soils do not have the same minerals as Mfp bedrock

Sample	Quartz	Microcline	Albite	HAp	Chlorite	Amphibole	Sum
B1	95	4	1	0	0	0	100
B2	95	2	1	0	2	0	100
Mfp	100	0	0	0	0	0	100
Obc	38	8	11	35	8	0	100
W840	92	2	2	4	0	0	100
Oh	87	3	1	9	0	0	100
HR	97	2	1	0	0	0	100
PL	90	2	5	0	2	1	100
RS	97	2	1	0	0	0	100
LL	98	0	0	0	2	0	100

Geochronological Methods

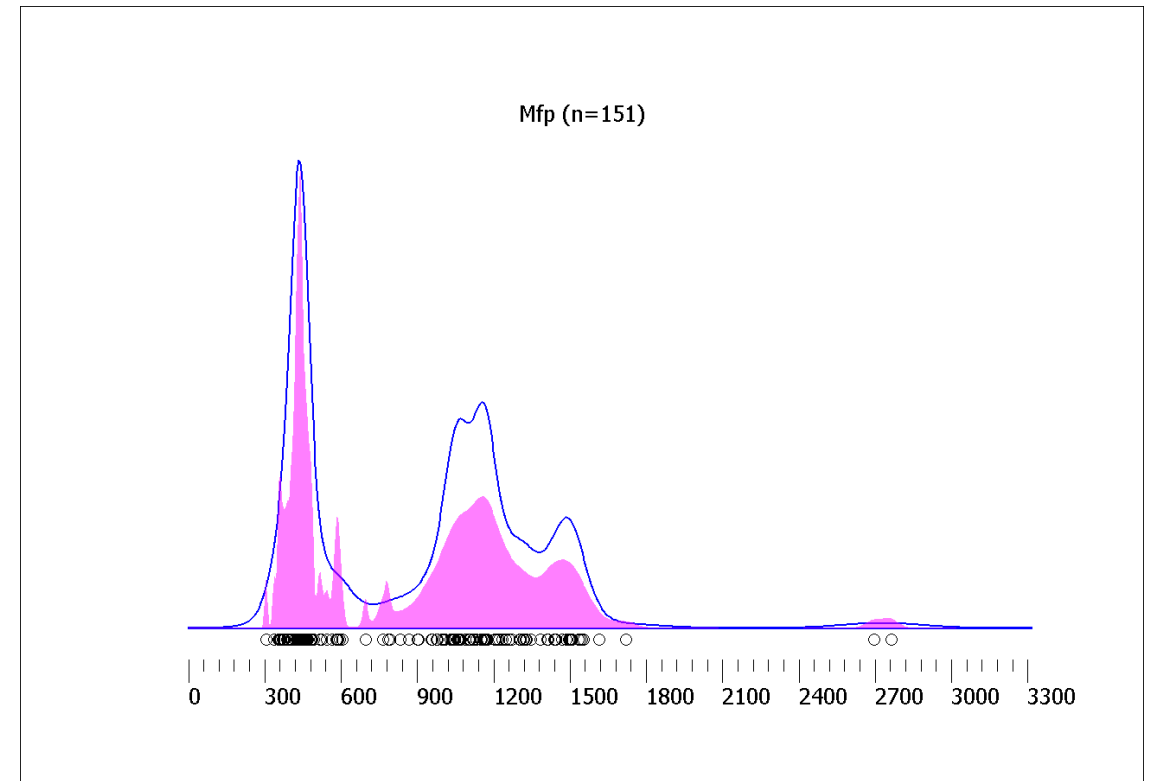
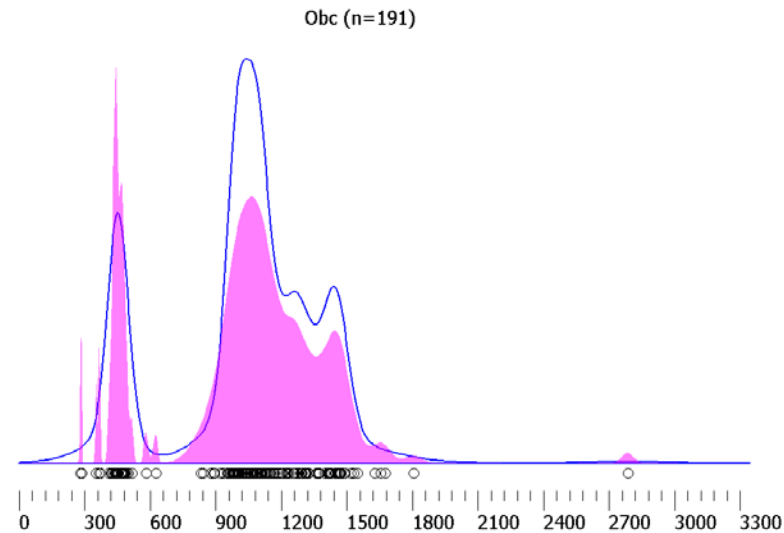
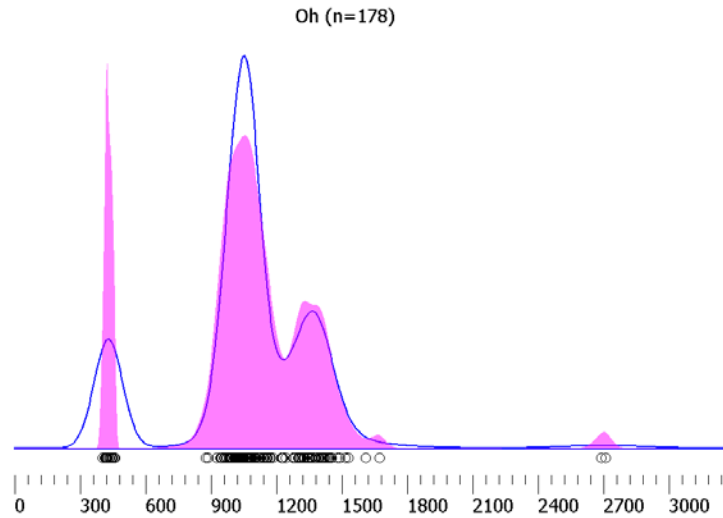
- CL imaging – zoning and inclusions
- ThermoFisher iCAP Qc quadrupole ICP-MS with CETAC autosampler and Photon Machine Excite 193nm excimer laser ablation system
- 50 x 50 μm spot size
- Data processed in Glitter, then ET_Redux v. 3.6.25
- Analyses that were > 20% positively or negatively discordant were rejected, where % discordance was calculated as
$$100 - (100 * (^{206}\text{Pb}/^{238}\text{U} \text{ date}) / (^{207}\text{Pb}/^{206}\text{Pb} \text{ date}))$$
- Age spectra plotted as kernel density estimates using “DensityPlotter” program of Vermeesch (2012)

Zircon CL Images



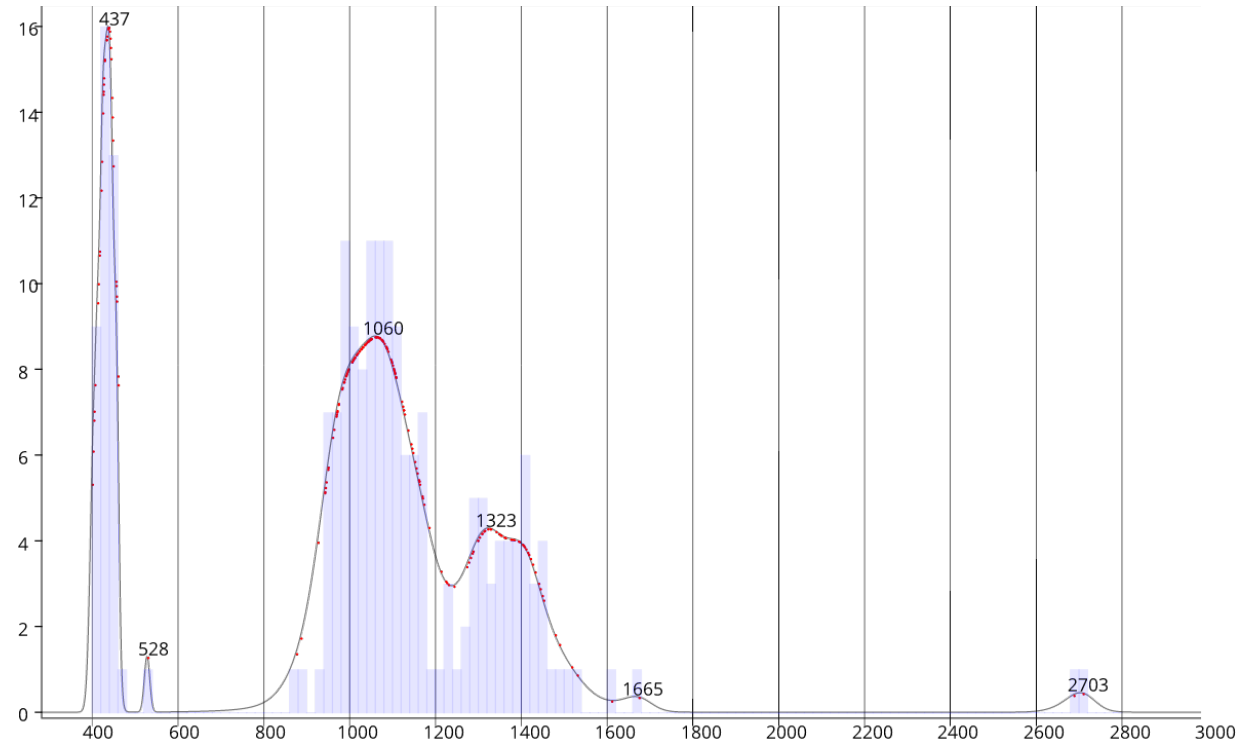
Bedrock Units

- Taconic orogeny: 430-450 Ma
- 1000-1100 Ma ages consistent with previous ages from southern Appalachians



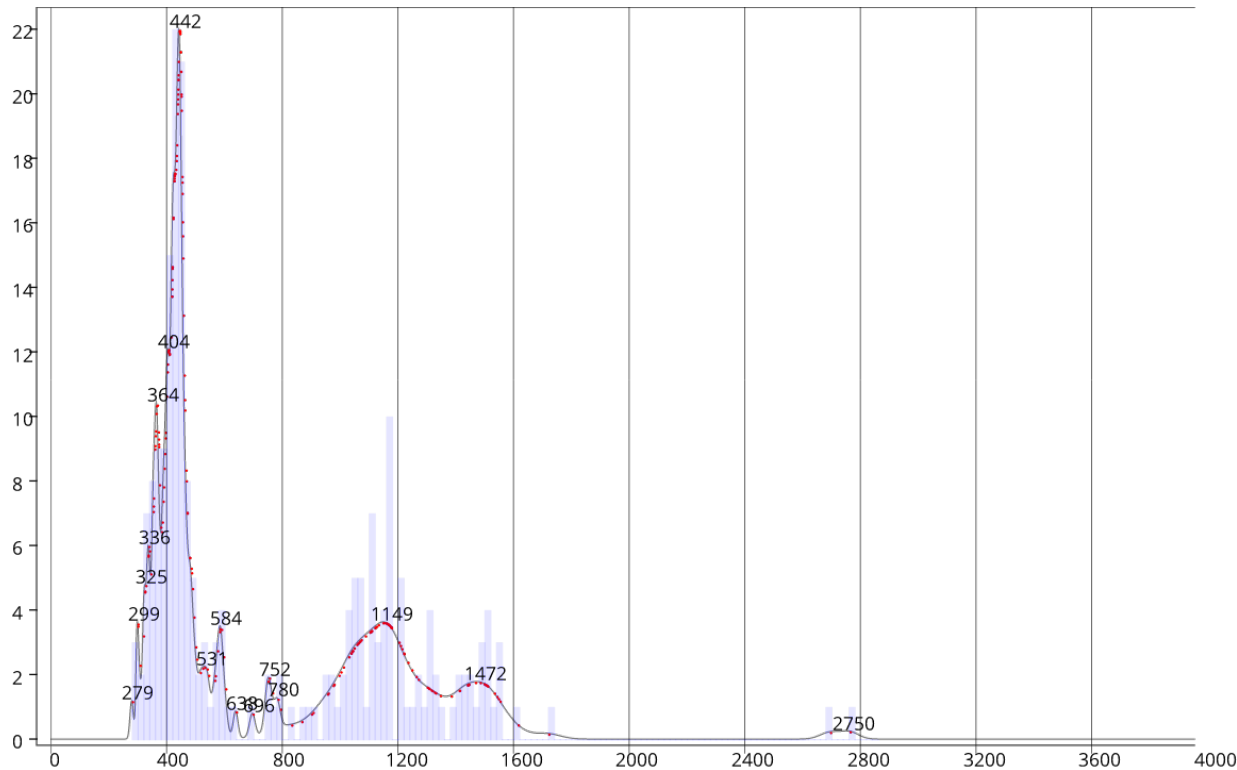
Pink represent probability density function, blue line represents the kernel density estimate (Vermeesch, 2012).

Tectonic Setting: Oh sediments



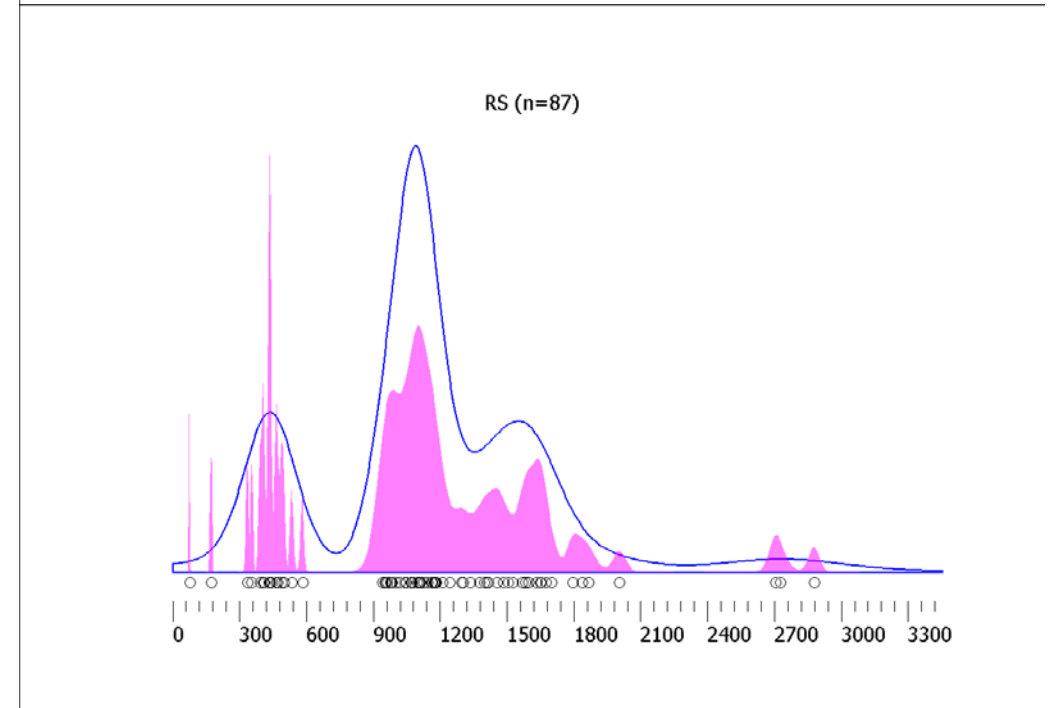
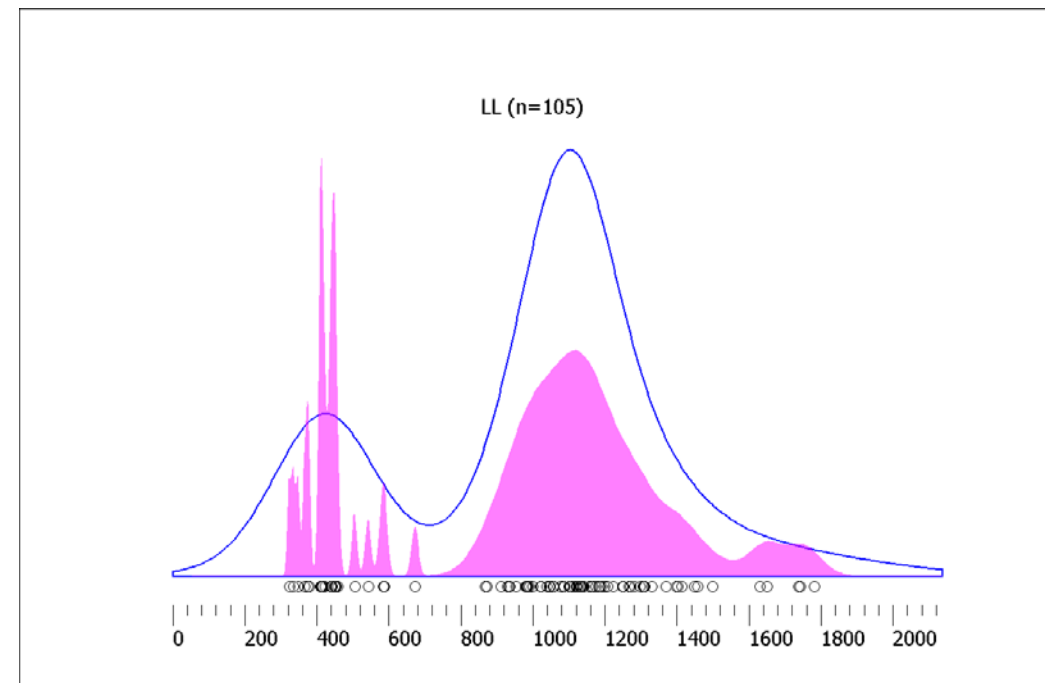
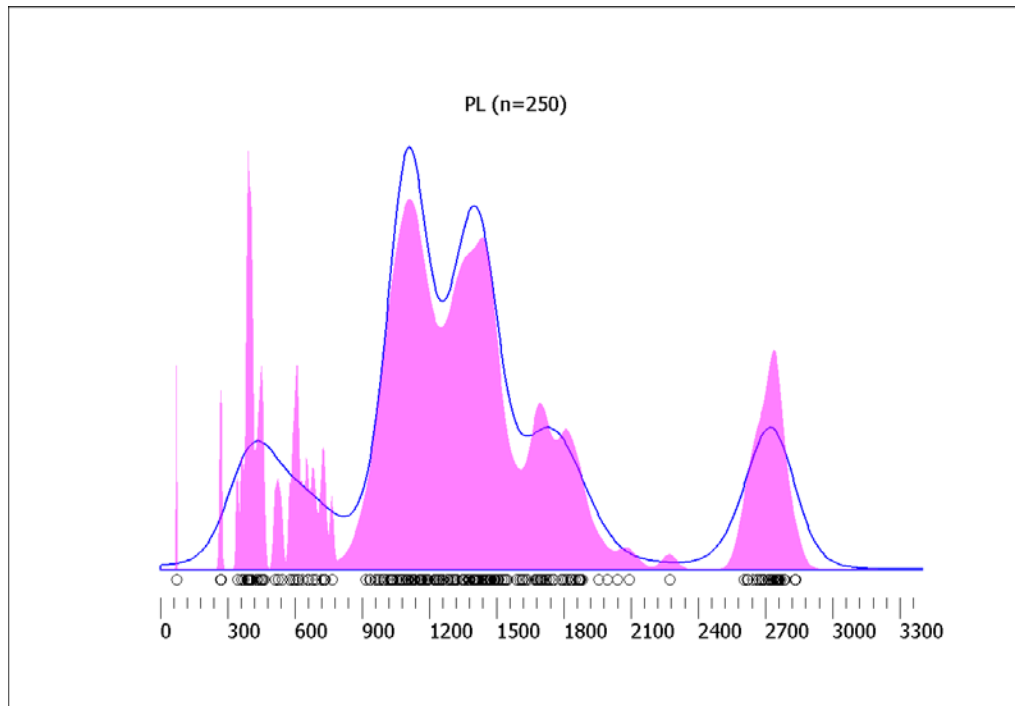
- Oh depositional age 453 Ma (Holland and Patzkowsky, 1997)
- Youngest age peak close to depositional age
- Consistent with a convergent plate margin (supra-subduction zone) setting during the Taconic orogeny.

Tectonic Setting: Mfp Sediments



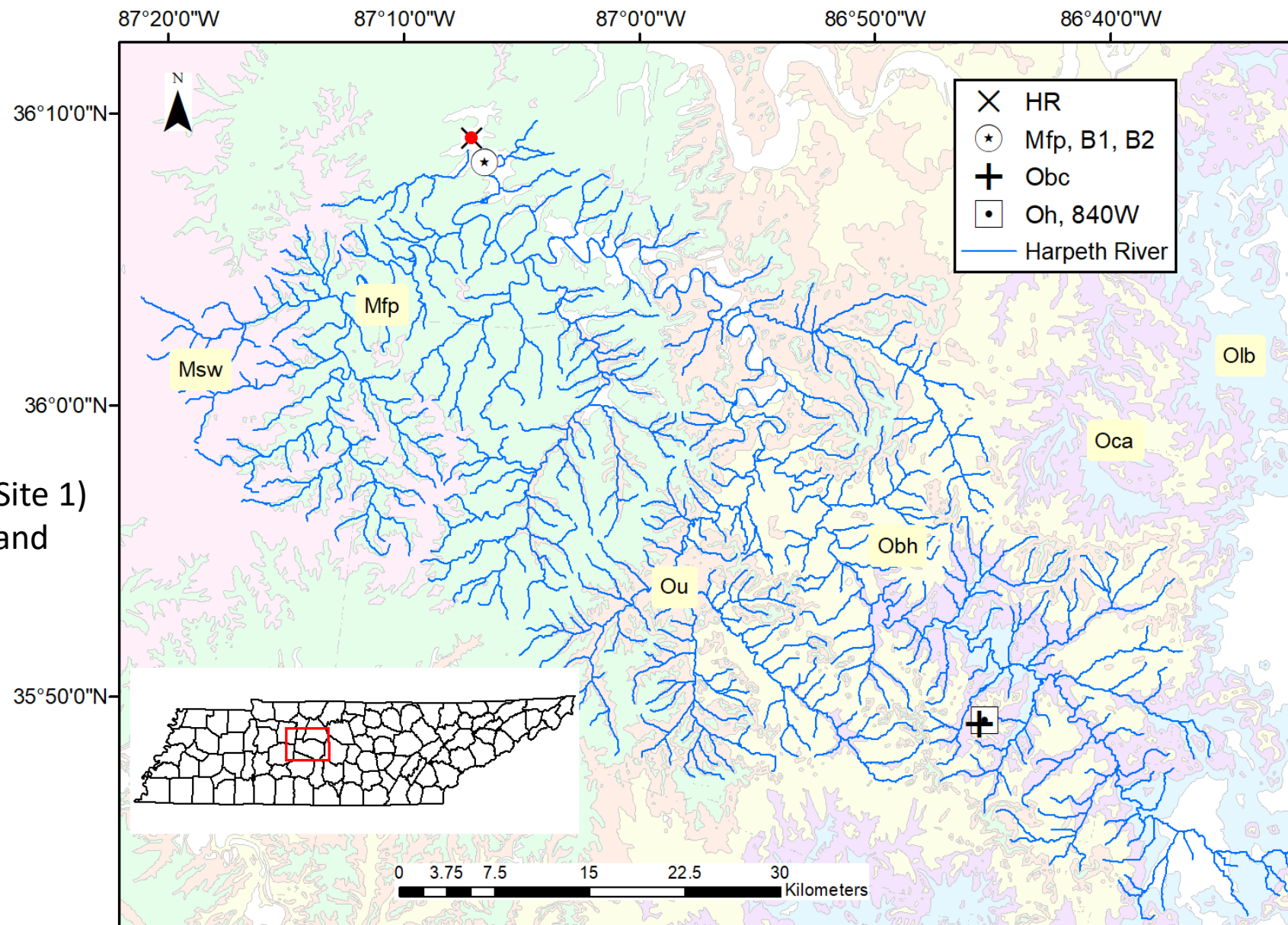
- Mfp is lower Mississippian, so depositional age ~ 350 Ma
- Youngest prominent age peak ~ 440 Ma
- Zircon crystals deposited ~ 90 Ma after crystallization
- Suggests a collisional setting, specifically the foreland basin of the Appalachians (Cawood et al., 2012).

Loess samples

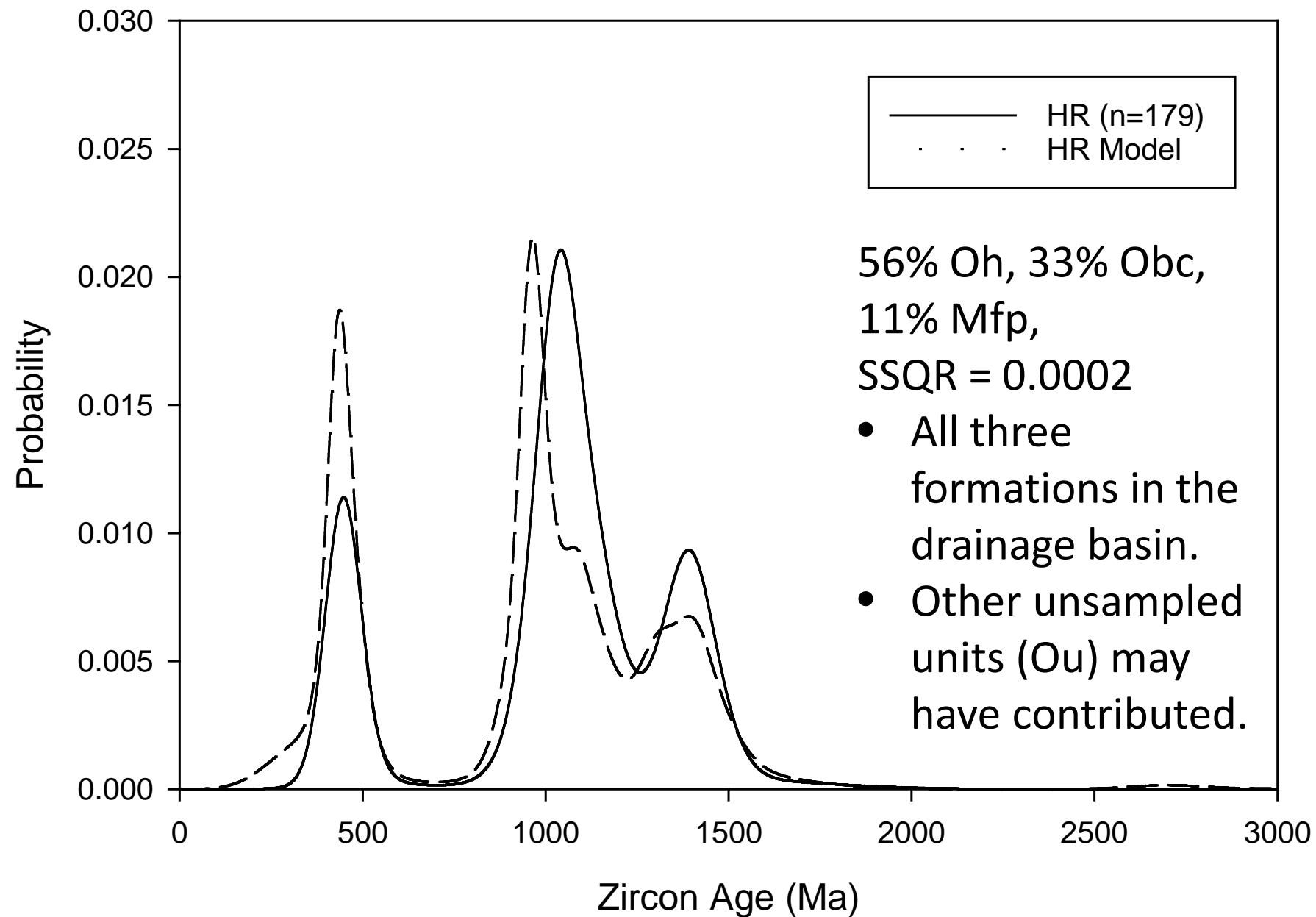


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- Obh = Ordovician Bigby Cannon and Hermitage limestones (site 2)

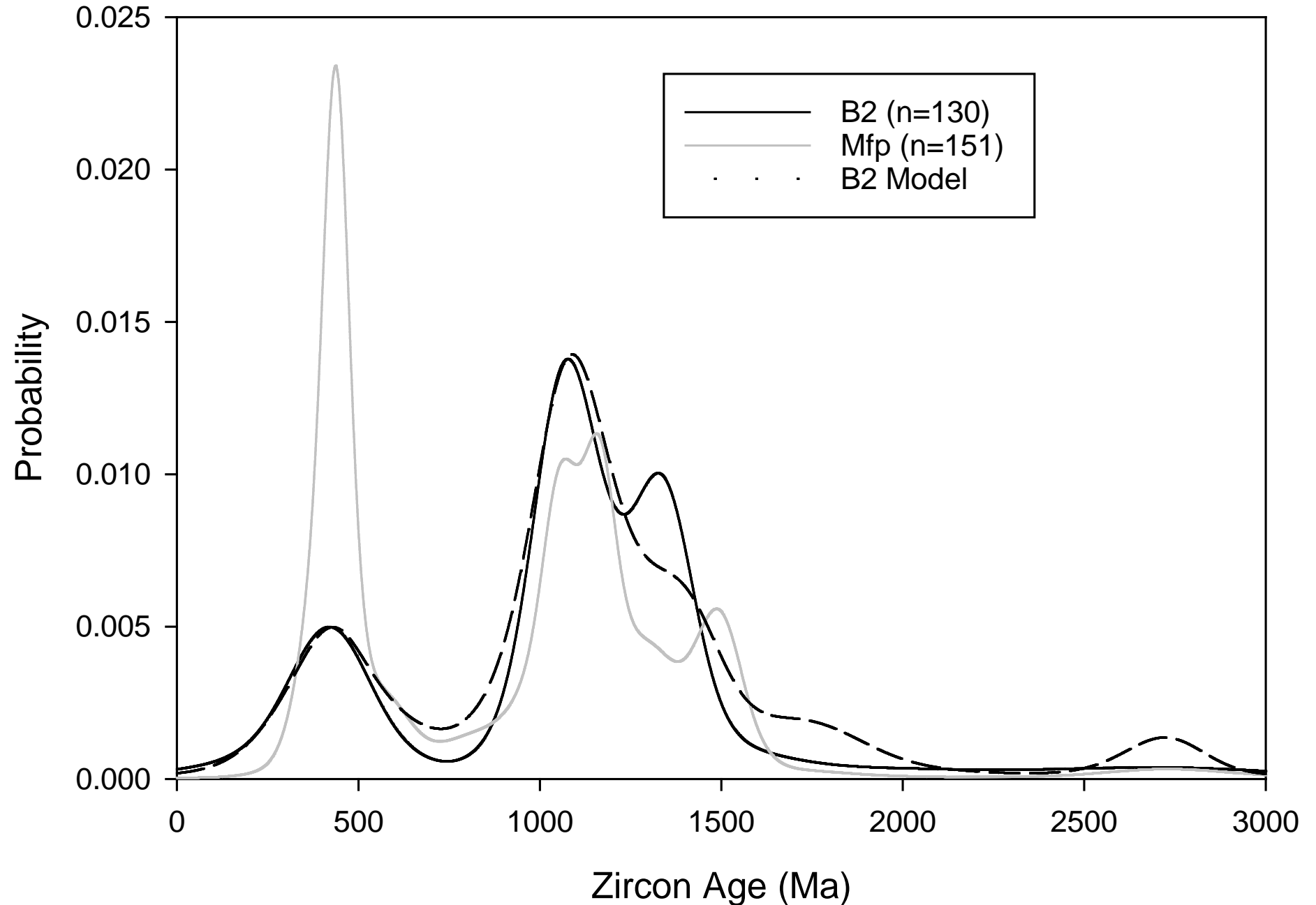


Harpeth River Age Spectra



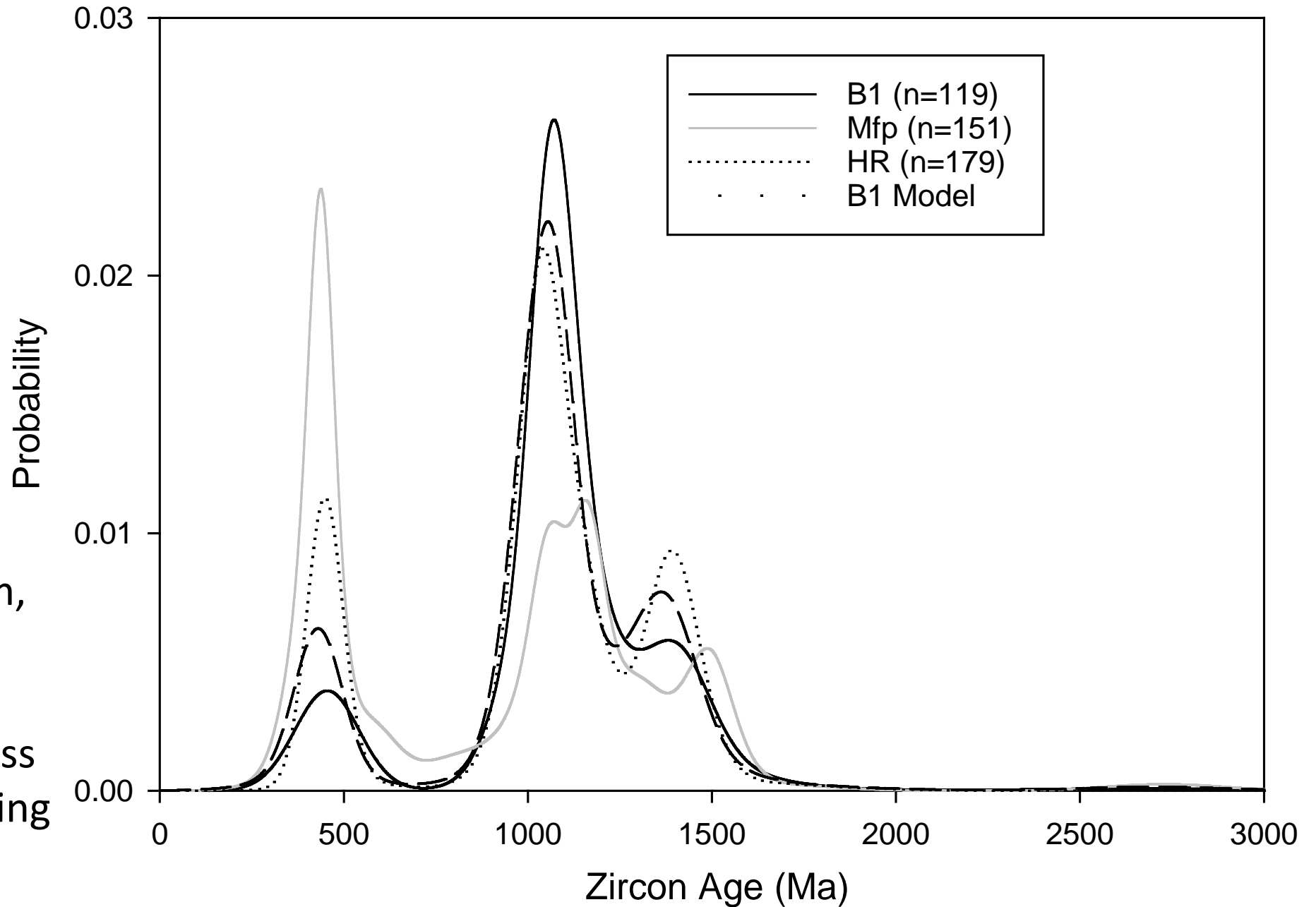
Deep layer B2 Age Spectra

47% LL, 29% PL,
24% Oh
SSQR = 0.00028
Since Oh is deep
beneath the
surface at
site 1, must have
been
deposited as
HR alluvium.

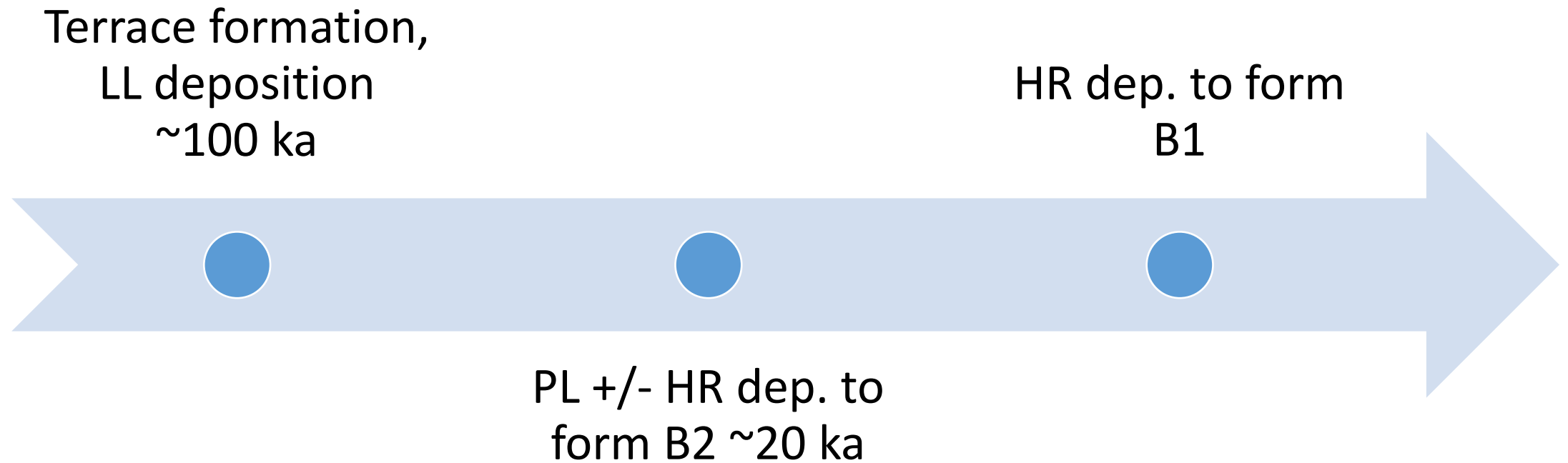


Shallow layer B1 Age Spectra

90% Oh as HR alluvium,
10% LL
SSQR = 0.0006
LL seems unlikely unless
there was vertical mixing

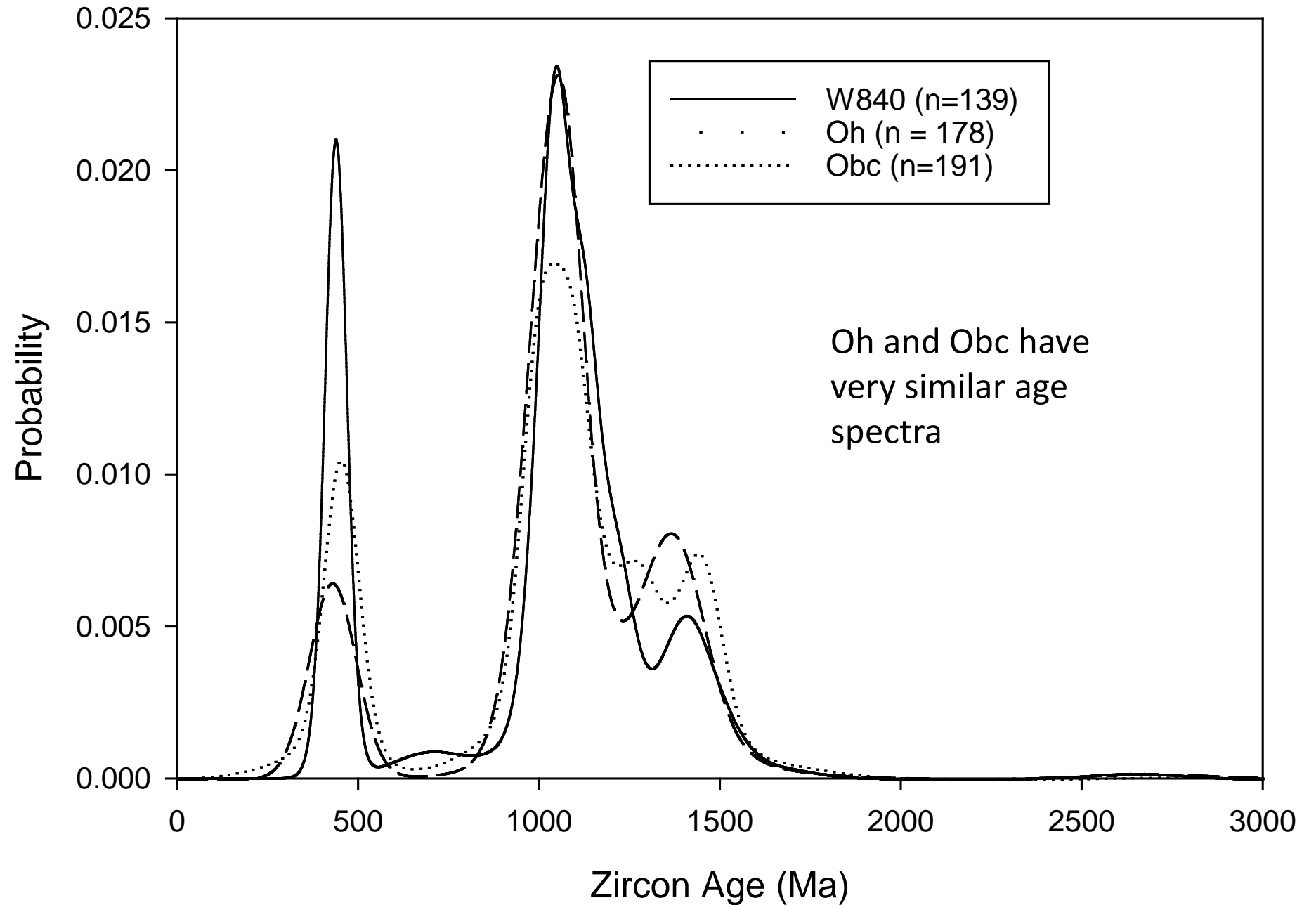


Timeline for Site 1



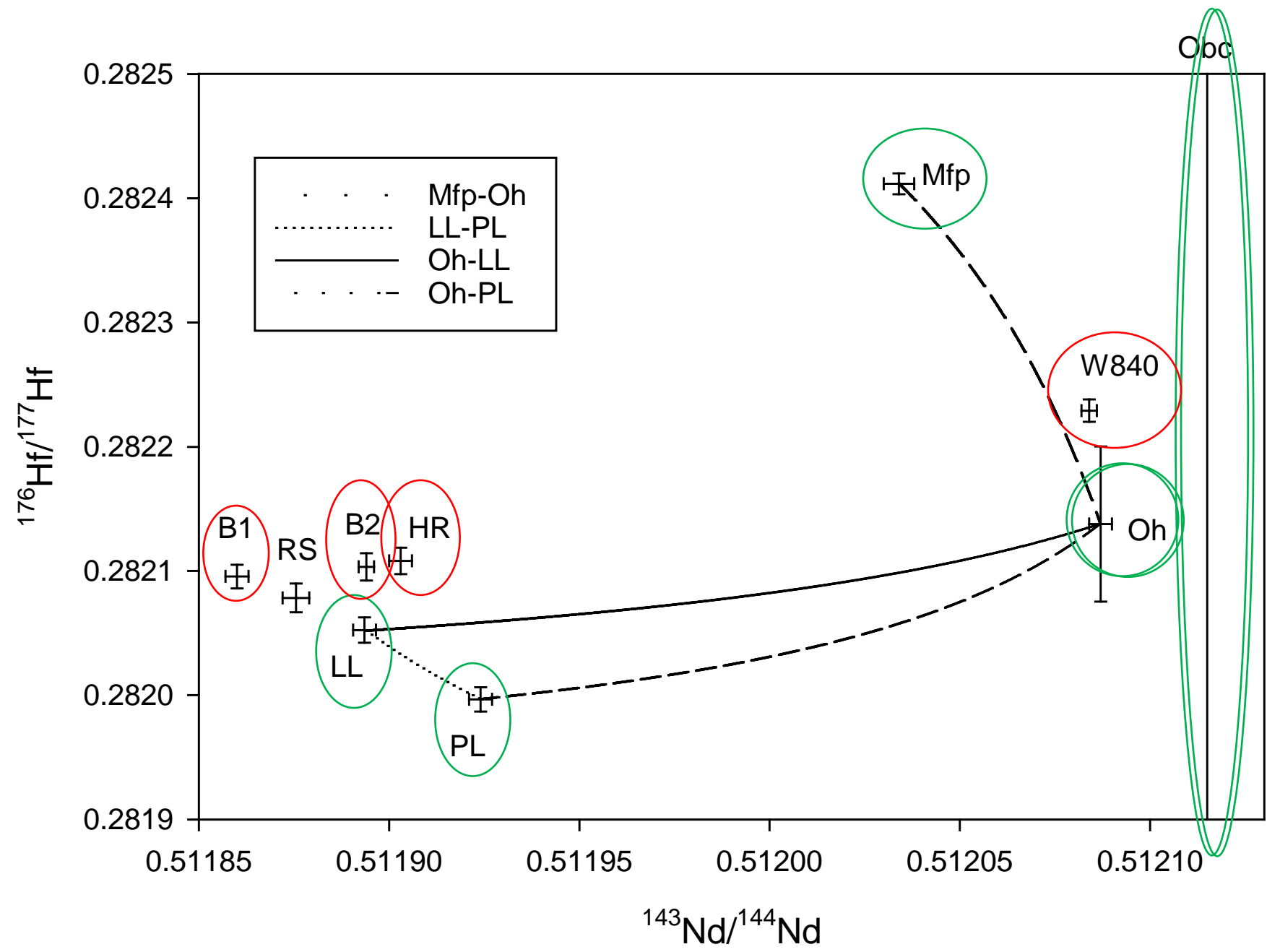
Site 2 W840 Soil Age Spectra

82% Obc,
18% Oh \pm Mfp
SSQR = 0.001

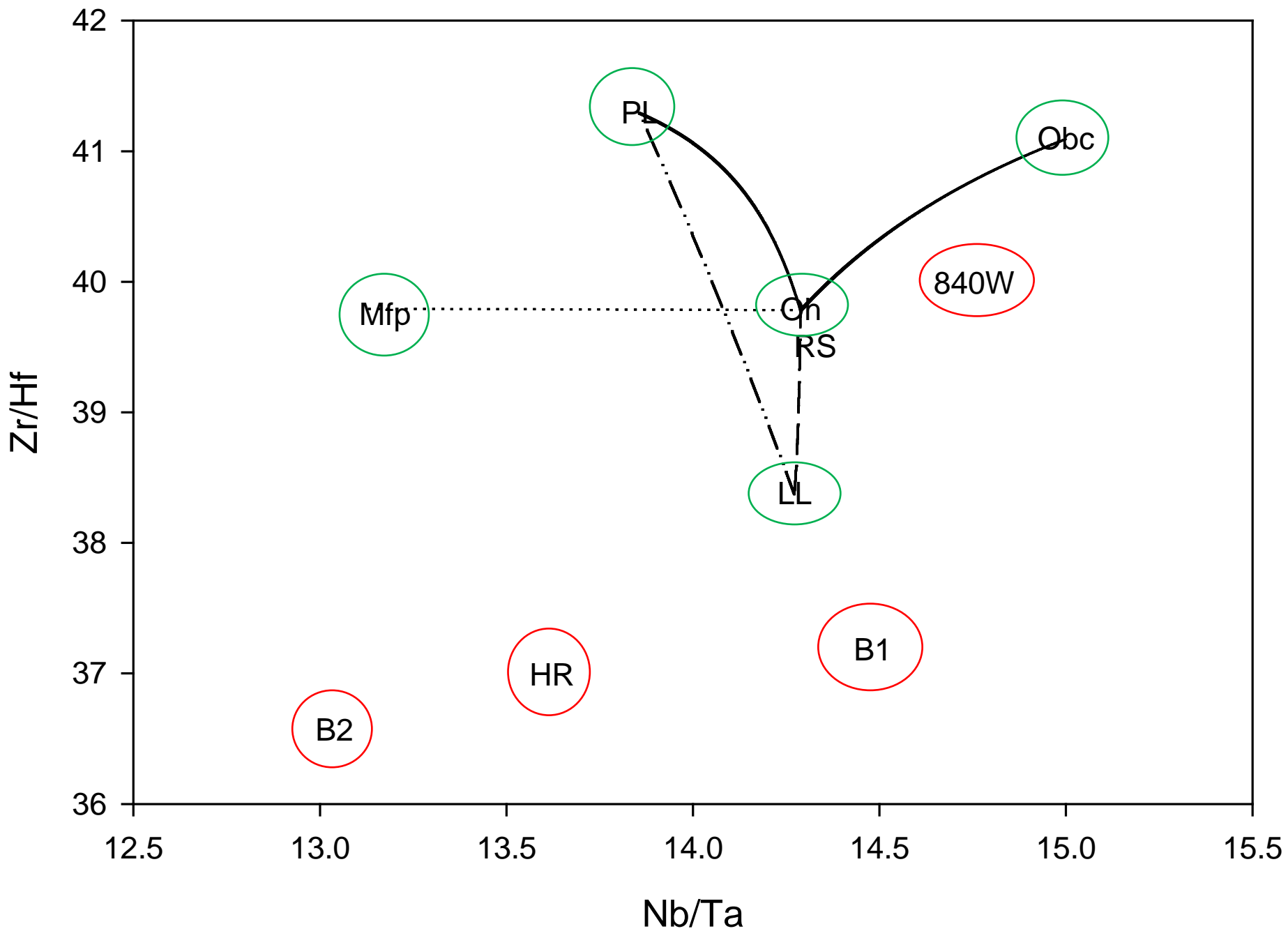


Bulk Isotope Results

B1 and B2 not
close to Mfp



Trace Element Results



Evidence Consistent With Genetic Relationship?

Parent-soil pair	Mineralogy	U-Pb age spectra	$^{143}\text{Nd}/^{144}\text{Nd}$	$^{176}\text{Hf}/^{177}\text{Hf}$	Zr/ Hf	Nb/ Ta	Gd _N / Yb _N	Eu/ Eu*	τ_j^{\S}
Mfp-B1	N	N	N	N	N	N	Tie†	N	Y
Mfp-B2	N	N	N	N	N	Y	Tie	N	Y
HR-B1	Y	Y	N	Y	Y	N	Tie	Y	
HR-B2	Y	Y	Y	Y	Y	N	Tie	Y	
Oh-W840	Y	Y	Y	Y	Y	?	Y	Y	Y
Obc-W840	?	Y	N		?	?	N	N	Y
HR-W840	N	Y	N	N	N	N	N	N	
<u>Binary Mixtures</u>									
(Oh+LL)-B1	N	Y	N	Y	N	Y	N	N	
(Obc+Oh)-W840	?	Y	N	Y	Y	Y	Tie	Tie	
(PL+LL)-B2	Y	Y	Y	N	N	N	N	N	
(Mfp+Oh)-W840	Y	N	Y	Y	Y	N	Tie	Tie	

† "Y" if parent is closest of all parents to soil, or to binary mixing lines for pairs of parents. For mixtures, soil must plot in-between sources. "?" indicates ambiguous results. "Tie": Mfp and HR have equal Gd_N/Yb_N, and W840 has almost identical Gd_N/Yb_N and Eu_N/Eu*_N as Oh.

§ Mass transfer coefficient τ calculated only for soil-bedrock pairs.

Conclusions

- Bedrock age peaks indicate source rocks for sediments formed during the Taconic orogeny, but many older zircons preserved from prior orogenies.
- Above Fort Payne chert bedrock, ultisol with exotic source:
 - Deposition of Loess and Harpeth River alluvium to form ultisol B2 soil horizon + intense weathering.
 - Deposition of Harpeth River alluvium to form ultisol B1 horizon.
- In-situ weathering of Oh + Obc to form W840 alfisol

Use of zircon U-Pb ages for soil provenance

- Zircon is recoverable from limestone.
- Zircon U-Pb geochronology is an effective tool for determining provenance of soil.
- However, best match in zircon U-Pb mixing model does not always agree with other datasets such as Nd and Hf isotopes.
 - Similar age spectra of our endmembers: all rocks formed from Appalachian sediments.
 - Harpeth River alluvium derived by erosion of those rocks with similar age spectra.
 - Loess deposits have many age peaks, some of which overlap with Appalachian sediment age peaks.

Applications to Igneous & Metamorphic Rocks

- Testing for soil formation by in-situ chemical weathering would be easier if bedrock had only one age peak, which is more likely for igneous bedrock.
- Provenance of metamorphic rocks: Compare zircon U-Pb age spectra of protolith and related metamorphic rocks. Could provide unequivocal evidence for metamorphic growth (ages younger than the youngest age in the protolith)
- Especially useful for metamorphic rocks with clastic sediment protoliths.