Investigation of the Shallow Hydrogeologic System on St. Catherines Island Georgia

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Geological Setting (shoreline systems)

St. Catherines Island

(After Hoyt & Hails, 1968)
St. Catherines Island geology and previous research

Sampling and analyses of wells in Upper Floridan indicate saline water contamination from below (Reichard et al., 2014).

Salt water intrusion by upconing and the past occurrence of Artesian springs requires structural access through confining layers.

Chloride spikes (salt water intrusion) detected in the surficial aquifer.
Current Research:

* **Determine** salt water intrusion pathways into surficial aquifer.

* **Evaluate** potential for communication with the Upper Floridan aquifer.

**Methods:**

- Monitoring head and water chemistry
- Resistivity & GPR surveys
- Coring as needed.
“Snapshot” of chloride concentration in the surficial aquifer on St. Catherines. Circled wells have concentrations that change rapidly in response to peak tides.

How do you explain the irregular occurrence of chloride spikes with respect to well proximity to shorelines?
Potential salt water intrusion pathways:

- Flooding and vertical infiltration
- Lateral intrusion via tabular sand aquifers or buried channel sands
- Lateral intrusion through structural pathways (joints, faults)
- Vertical intrusion through structural pathways (joints, faults, sag structures)
Coastal Plain joint trends after Bartholomew et al., 2007

Yellow Banks Joint Trend

N24°E trend (M1) is same as interpreted Brunswick fault trend of Maslia and Prowell (1988) and Atlantic Coast Fault System

Structural Components of St. Catherines Island

Pre-1960 wetlands in blue overlay.
Faults and joints focused dissolution in Upper Floridan carbonates. Collapse of deep caverns produced sag structures and basins in surficial sediments.
Sinkhole near Middleground Community, Bulloch County
Diameter ~ 26 meters, Depth to Upper Floridan carbonates > 300 ft
Exploring Links Between Structure and Hydrology

Resistivity profile (above) and 100 MHz GPR profile (right).

aquifer

aquitard
Core data at well site S4:
0 – 1.5 m: hydric black sandy top soil
1.5 m – 7.3 m: fine–very fine, subang., well sorted qtz sand
7.3 m – 11 m: muddy very fine qtz sand and mud
11 m – 13.4 m: fine to very fine qtz sand

• Shallow wells < 7.3 meter depth
• Sharp density increase at ~ 5 m depth (2% compaction)
• SCI clays are kaolinite dominant
Radar element offsets suggest minor faults and fractures (sag structure with offsets highlighted)

100 MHz GPR profile
A Test: Are offsets true subsurface displacements or artifacts from surface irregularities? They appear in same location for parallel profiles run from opposite directions.
100 MHz GPR profile across NW-SE trending lineaments, offsets highlighted in lower figure.
Notes on faulting in soft sediments:

• Transition from single fault at depth to splays at ~ 35 to 20 m depth (see Basson et al., 2002 combined seismic & GPR study of Dead Sea Rift.

• Faults may produce deformation bands through combined effects of cataclasis and compaction (See Fossen et al., 2007; Cashman and Cashman, 2000)

• Deformation bands may result in increased density and reduced permeability in band (See Bense, 2004 – Roer Valley Rift study)
Summary

• Faults and joints and associated sag structures permitted former Artesian springs and now allow salt water intrusion via upconing.
• The surficial aquifer is also experiencing local salt water intrusion linked to peak tides.
• Structural features (faults and sags) are probably responsible for irregular nature of salt water intrusion in shallow wells.
• Buried channels may be another potential salt water intrusion source.
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