Failure mechanism of large landslides triggered by earthquakes in the El Salvador Fault Zone: the case of Jiboa landslide







INTRODUCTION

El Salvador, due to its geodynamic setting – close to the Cocos-Caribbean segment of the Middle American subduction zone, where the plates converge-, is affected by a great number of earthquakes of large and moderate magnitudes. One of the secondary effects associated with these events, which contributes to a large extent to increase the destruction and loss of life, are the landslides. During the last decades El Salvador has suffered many earthquakes of large and moderate magnitude that have caused severe damage and great social alarm, fundamentally because of the associated geological process (landslides, liquefaction, etc.) (White and Harlow, 1993; Bommer et al., 2002). The impact that these phenomena can have was demonstrated with the two 2001 El Salvador earthquakes, that produced and reactivated a great number of landslides across the whole country but with a great concentration in the central part of the El Salvador. During these two events nearly 1,200 people died, from which around 500 casualties were directly related with the induced landslides (CEPAL, 2001).

In this work, we present the slope stability analysis and the failure mechanisms proposed for the Jiboa landslide, which is one of the greatest co-seismic slope instabilities triggered during El Salvador earthquake of February 13, 2001, with a magnitude of M_w 6.6. The seismic events that took place at the beginning of 2001, besides inducing a large number of shallow slope instabilities, triggered large landslides putting hundreds of thousands of cubic meters of material into motion. These large landslides produced the greatest damage in both economic losses and human lives.







Location of the two 2001 El Salvador earthquakes (M_w =7.7 and M_w =6.6).



The Jiboa landslide volume is estimated > 12 million cubic meters with a vertical drop about 170 m. The landslide dammed with debris about 700 m long and 250 m

wide of the valley of Jiboa river and the upstream lake could be about 60 m deep

and 2 km long (Baum et al., 2001).



SLOPE STABILITY ANALYSIS OF JIBOA LANDSLIDE



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a) Real PGA = 0,408g(Zacatecoluca seismic station)

STABLE (FS = 1.36)

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JIBOA LANDSLIDE

The Jiboa landslide occurred near the fault rupture zone in the confluence of the El Desagüe and Jiboa rivers. In the area where Jiboa landslide took





Simplified geological profile of Jiboa landslide caused by the February 13th 2001 El Salvador earthquake and detailed view of materials (C) at the slope.

2) Pseudostatic analysis



The Jiboa landslide slope comprises blocks of andesitic breccia (frequently interbedded with pyroclasts) and rhyolitic tephra from Cuscatlán Formation. The upper part of the slope is covered by a thin recent pyroclastic deposit of Tierra Blanca and paleosoils. The breccia (> 160 m thick) consisted of sand-sized to 50 cm-diameter rock fragments mainly basaltic and andesitic, and had poorly sorted interiors and coarse tops. The underling ryolitic tephra is a poorly compacted and very porous fine grained deposit. In situ and laboratory investigation has shown that these ryolitic tephra is frequently interbedded by thin highly porous and soft volcanic ash layers that are formed by silt and sand particles with a few amount of clay. These thin volcanic ash layers probably form a horizontal sliding surfaces which were essential for the generation of such deep landslide.

- - – Slope prior landslide Present slope and Landslide mass Observation point of contact between mobilized and in situ Present visible scarp ____ Possible failure surface •••• Contact between mobilized and in situ material Landslide Dam L = 700 m A = 200 m

CONCLUSIONS





rystalline pores.

SEM photograph showing an open squeletallike microstructure with a grain-grain contact and large pores, including important intra-

The geological and geotechnical studies carried out in situ, the laboratory tests, as well as the slope stability analysis, reveal that the failure of the Jiboa landslide took place in three independent phases:

Phase 1. The triggering factor of the landslide is attributed to the existence of very porous volcanic ash intercalations, which collapsed due to the released seismic energy and lost their strength parameters. The thin volcanic ashes have a very open texture with a meta-stable honeycomb like microstructure, high void ratio and low unit weight (1.2 kN/m³). These materials have a high apparent strength due to primary weak chemical and silty-clay cementation, but they are susceptible to large reductions in their strength due to shaking and flow like a semi-liquid mass (quick-silt).

Phase 2. The collapse produced a large settlement and, consequently, a tension crack was generated in the head of the slope.

Phase 3. Finally, the movement propagated by shearing failure along a sub-horizontal surface located in such volcanic ash intercalations.

The suggested failure mechanism of Jiboa landslide could be valid for similar large landslides triggered by earthquakes in the El Salvador, and it is coherent with previous model developed for co-seismic landslides (Huang, 2015).



