Coupled Pressure-Temperature Probes for Monitoring Porewater Fluxes in Coastal Sediment Nicole K. LeRoux^{1,2}, Joseph Tamborski^{1,3}, Sanjana Moodbagil^{1,4}, & Barret L. Kurylyk^{1,2}



1. Introduction

Submarine groundwater discharge (SGD) and porewater exchange is driven by hydraulic gradients between coastal aquifers/sediment and the ocean¹

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- When integrated over the large spatial areas of exchange, SGD flux can be comparable in magnitude to coastal riverine discharge²
- SGD delivers contaminants to the ocean, including nitrogen and heavy metals
- Ocean-aquifer interactions vary in magnitude and direction due to hydraulic oscillations associated with wave and tidal action (Fig. 1)



Figure 1: Conceptual model of ocean-aquifer interactions, including: (1) density-driven circulation, (2) tidal pumping, (3) wave pumping and (4) fresh SGD (modified from Robinson et al.¹)

2. Methods and Theory: Heat as a Groundwater Tracer

- When groundwater flows, it advects heat, disturbing subsurface temperatures
- The thermal effects of groundwater flow enable heat to be used as a groundwater tracer
- Multi-depth temperature sensors in sediment (Fig. 2, left) can reveal how periodic temperature signals propagate into the subsurface
- The thermal sine wave is lagged and damped with depth (Fig. 2d,e), but the precise signal transfer is influenced by the direction and magnitude of groundwater flow





Figure 2: Past setups for temperature monitoring: (a) drive-point piezometer with temperature sensors on a cable, (b) sensors embedded in a rod, or (c) high-resolution temperature sensing (HRTS) approaches (modified from Irvine et al.³). Under discharge conditions (d) the signal decays more than under (e) recharge conditions (modified from Kurylyk and Irvine⁴).

Governing partial differential equation: 1D, transient conduction-advection eq.

Boundary condition: Periodic surface temperature

• Analytical solution: Signal damping and lagging, where d and L depend on the Darcy flux⁵

$$\lambda_{app} \frac{\partial^2 T}{\partial z^2} - qc_w \rho_w \frac{\partial T}{\partial z} = c\rho \frac{\partial T}{\partial t}$$

$$T(z=0,t) = T_m + A\sin\left(\frac{2\pi t}{p} - \phi\right)$$

Damping Phase shift (lag)

$$T(z,t) = T_m + A \exp(-dz) \sin\left(\frac{2\pi t}{p} - \phi - Lz\right)$$

waterstudies.





3. Field Site & Challenges in Coastal Settings

- seldom been applied in tidal settings
- temperature and groundwater flux magnitude or even direction





- The temperature time series from
- The inferred fluxes (Fig. 5) do
- Fluxes estimated from the
- More information on the







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