

# Eocene exhumation of the High Andes differentiated by detrital multimethod U-Pb-He thermochronology and thermal history modeling

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# Acknowledgements

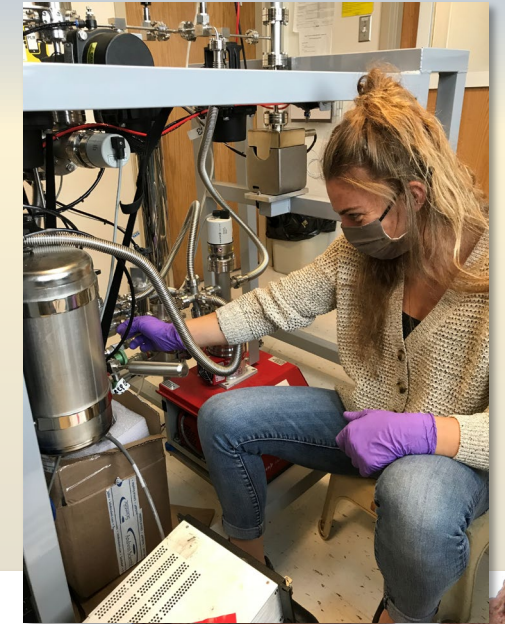
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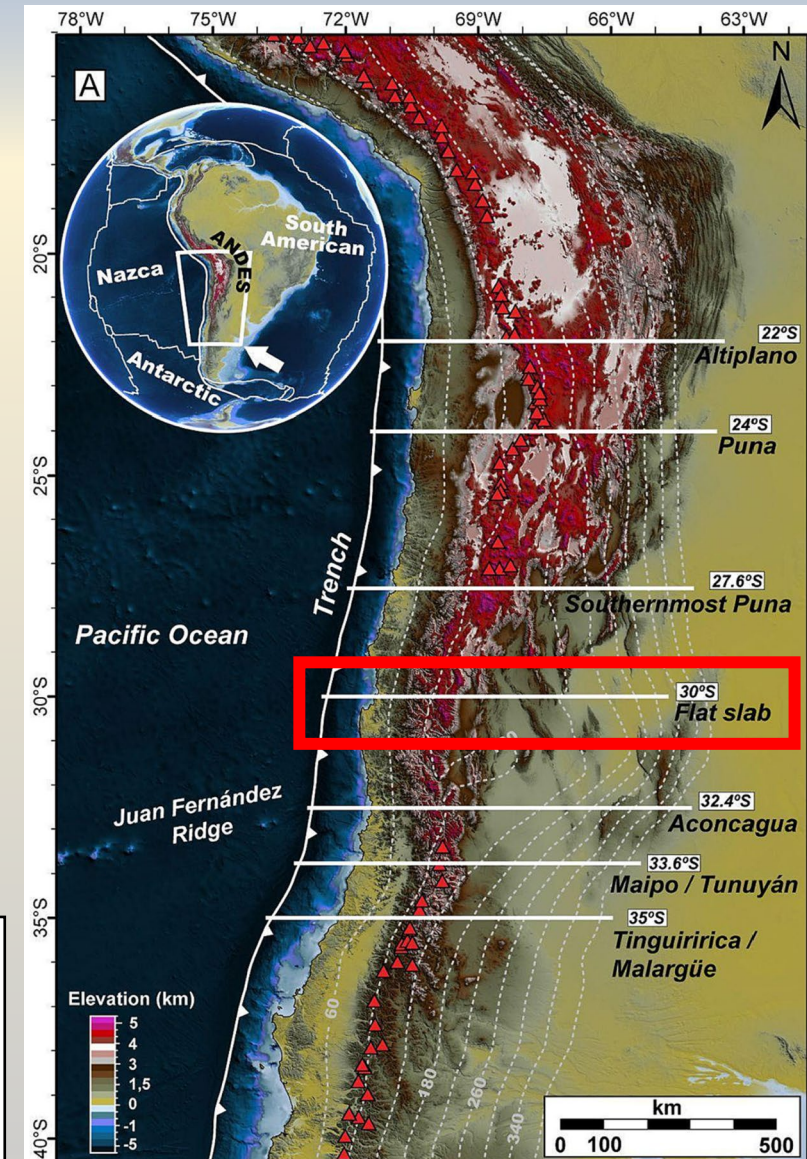


# The Central Andean Orogenic Belt

## Archetype of subduction-related orogenic system

- Linked processes between thrust-belt thickening, surface uplift, basin subsidence, and arc magmatism
  - Orogenic cyclicity *DeCelles et al., 2009; 2015*
  - Evolutionary stages of wedge growth *Giambiagi et al., 2022*
- Paleogene orogenic history remains contentious
  - Tectonic stasis/quiescence? *Horton and Fuentes, 2016; Horton et al., 2018*
  - Recent recognition of Paleogene retroarc strata *Fosdick et al., 2017; Suriano et al., 2023; Ronemus et al., 2024*
  - New structural and thermochronology of hinterland deformation *Lossada et al., 2017; Rodriguez et al., 2018; and Mackaman-Lofland et al., 2024*

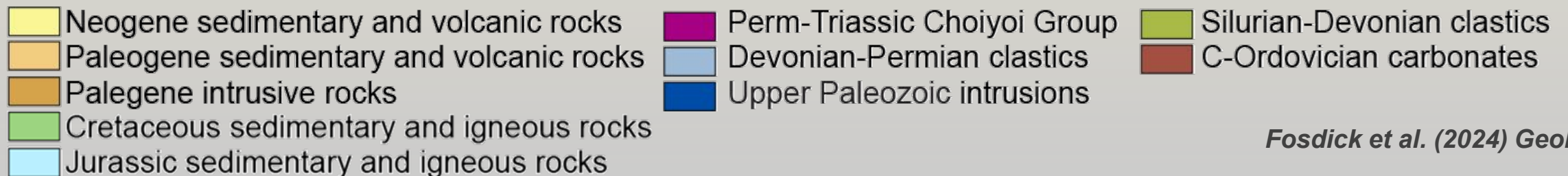
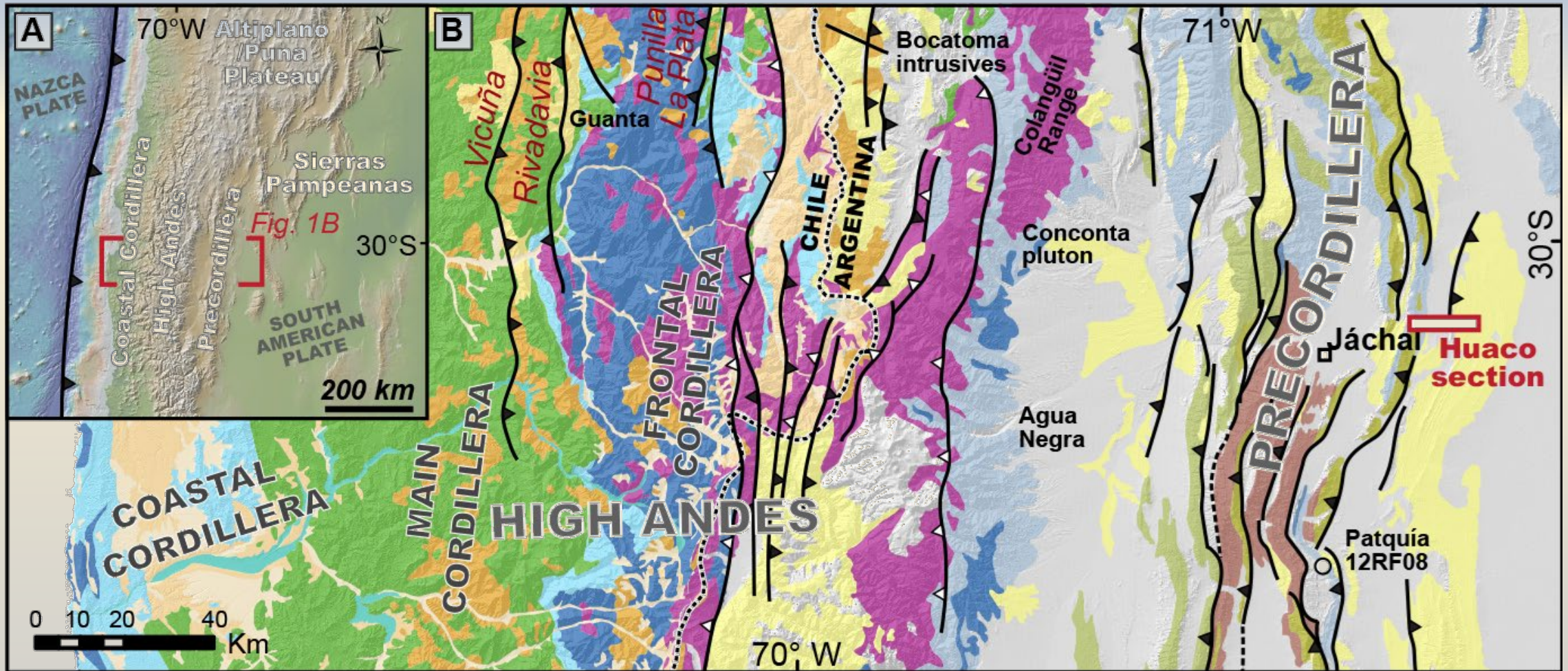
We revisit the detrital signature of Paleogene – Neogene strata to test whether there is evidence of hinterland exhumation and source-basin sediment routing



*Giambiagi et al. (2022)*

# Southern Central Andes 30°S

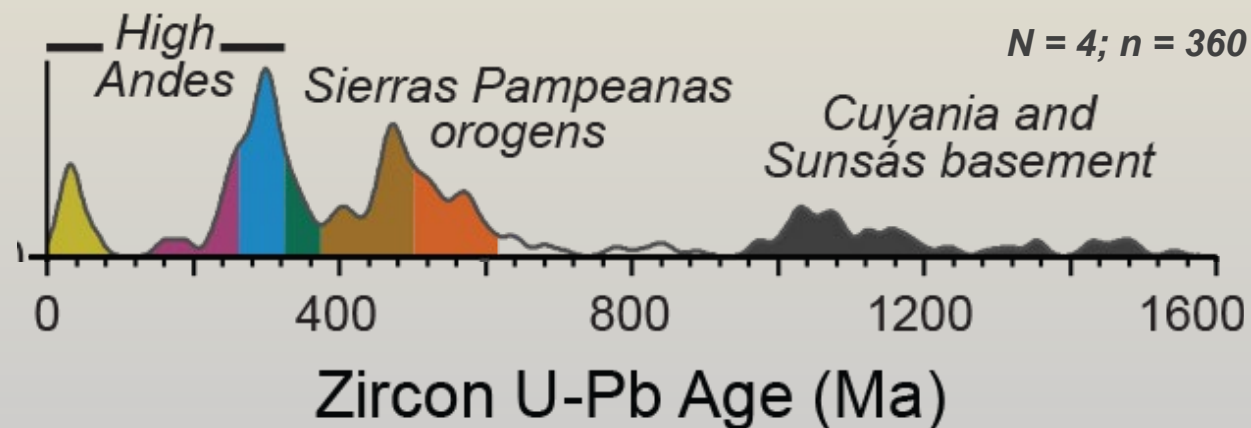
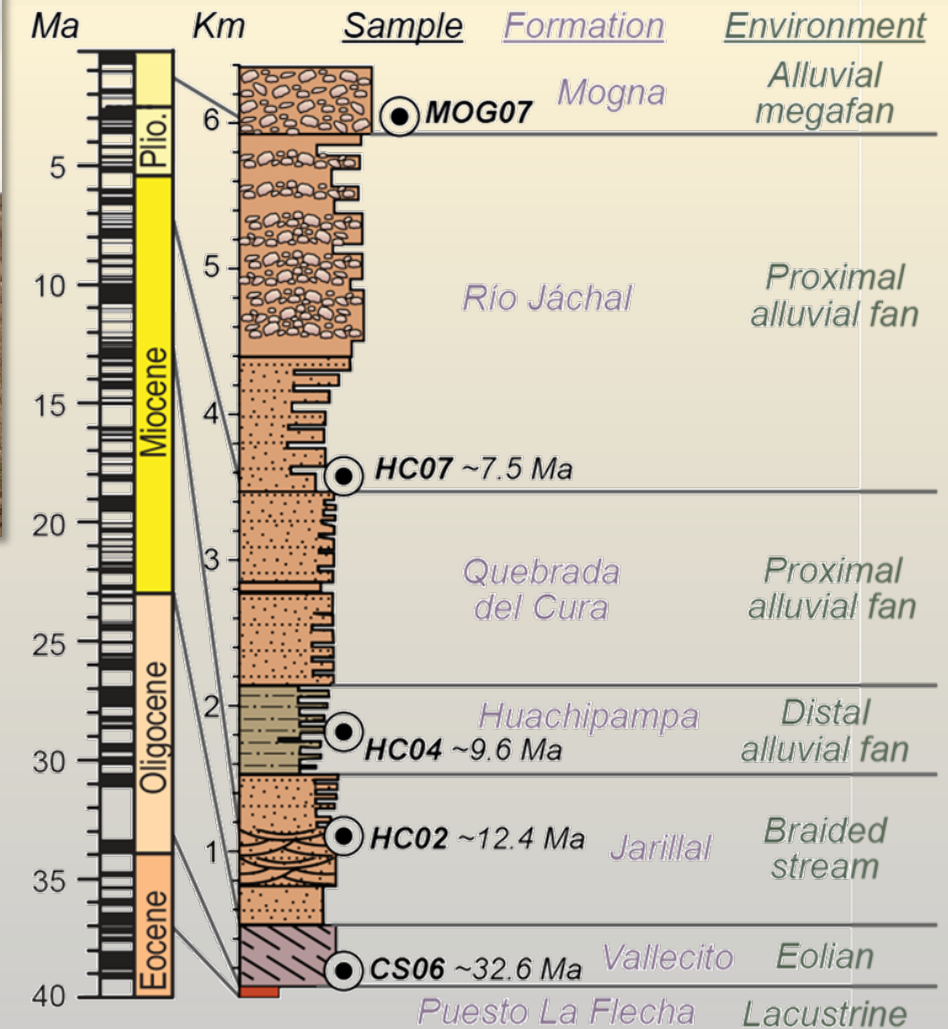
Pampean flat-slab segment



# Bermejo – Vinchina foreland basin



## Huaco Composite Section



# Bermejo – Vinchina foreland basin

**Vallecito Fm. (lower Oligocene to lower Miocene)**

*Eolian and fluvial deposits*

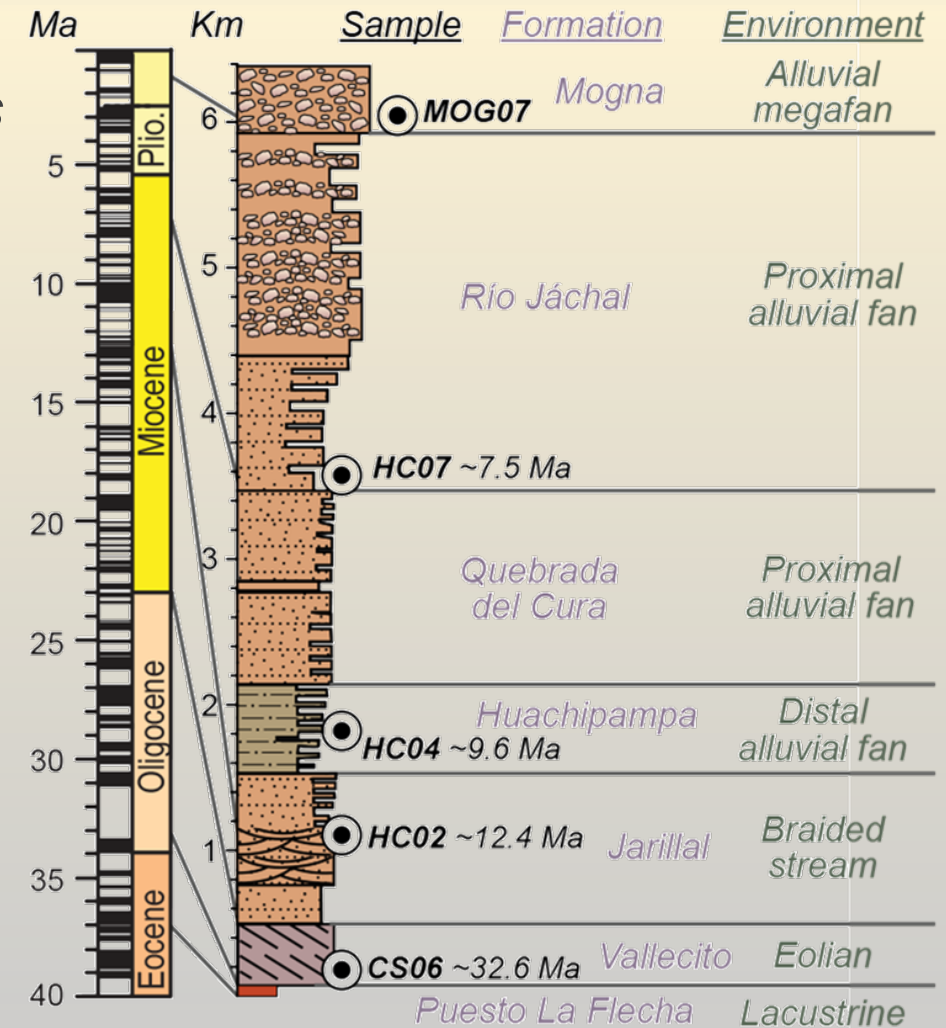
**Puesto de la Flecha Fm. (upper Eocene)**

*Distal lacustrine siltstone, sandstone, and evaporite deposits*



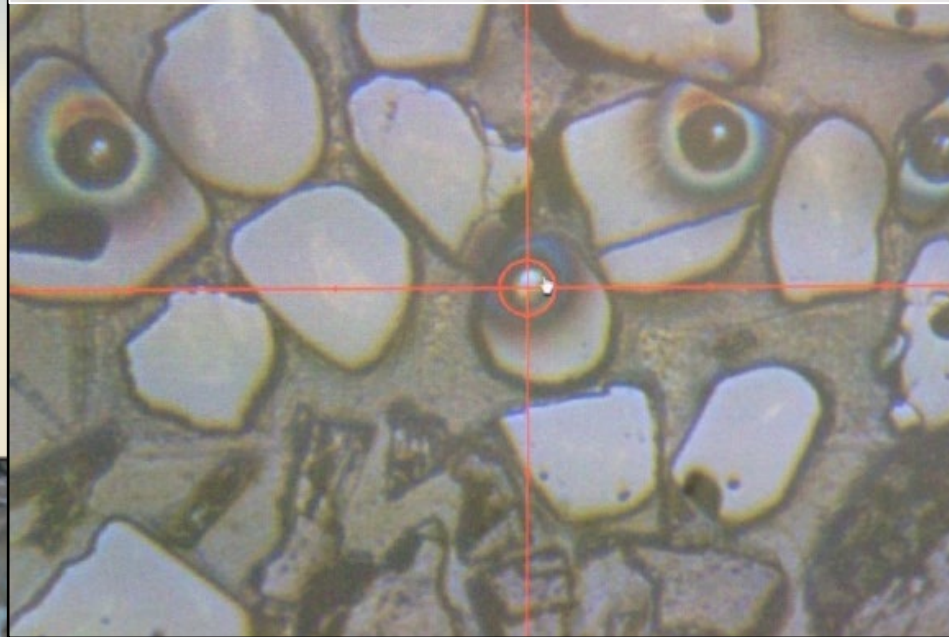
Fosdick et al. (2024)

## Huaco Composite Section

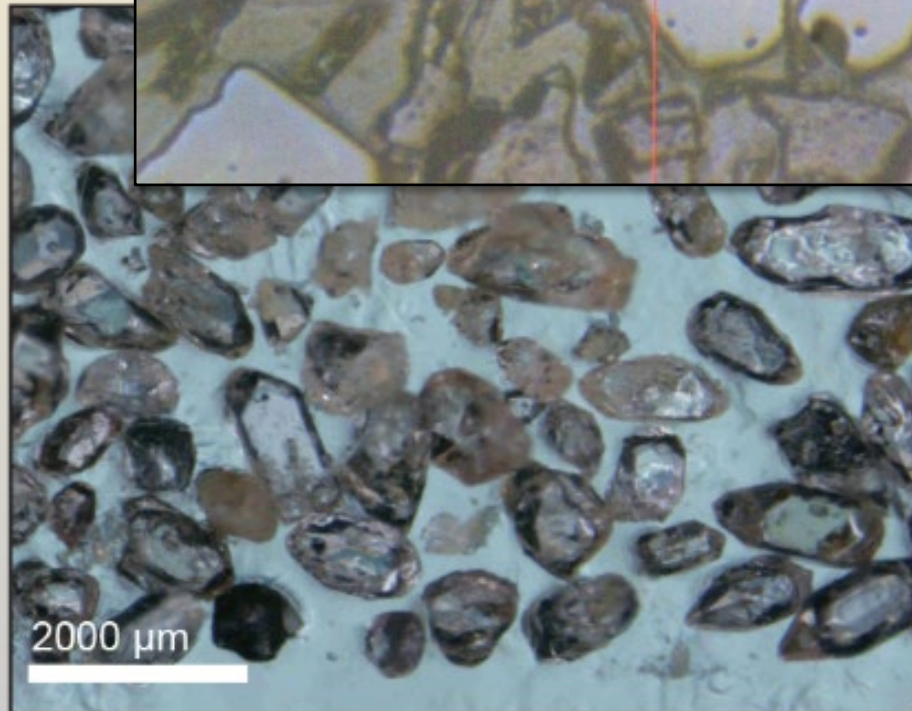
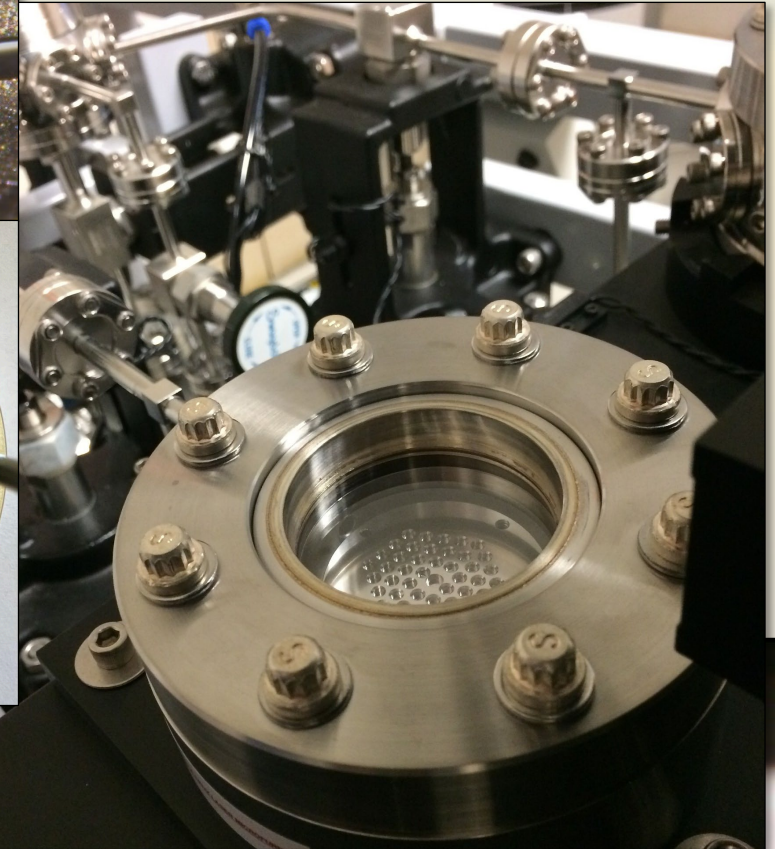
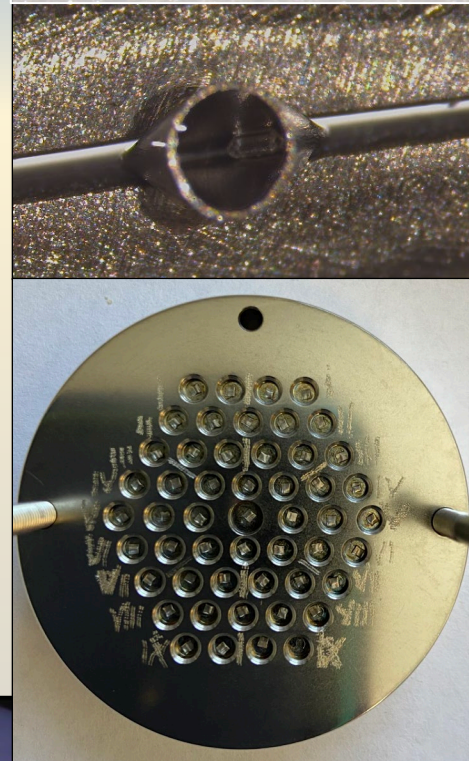


# Paired Detrital U-Pb and (U-Th)/He Thermochronology

LA-ICPMS U-Pb analysis

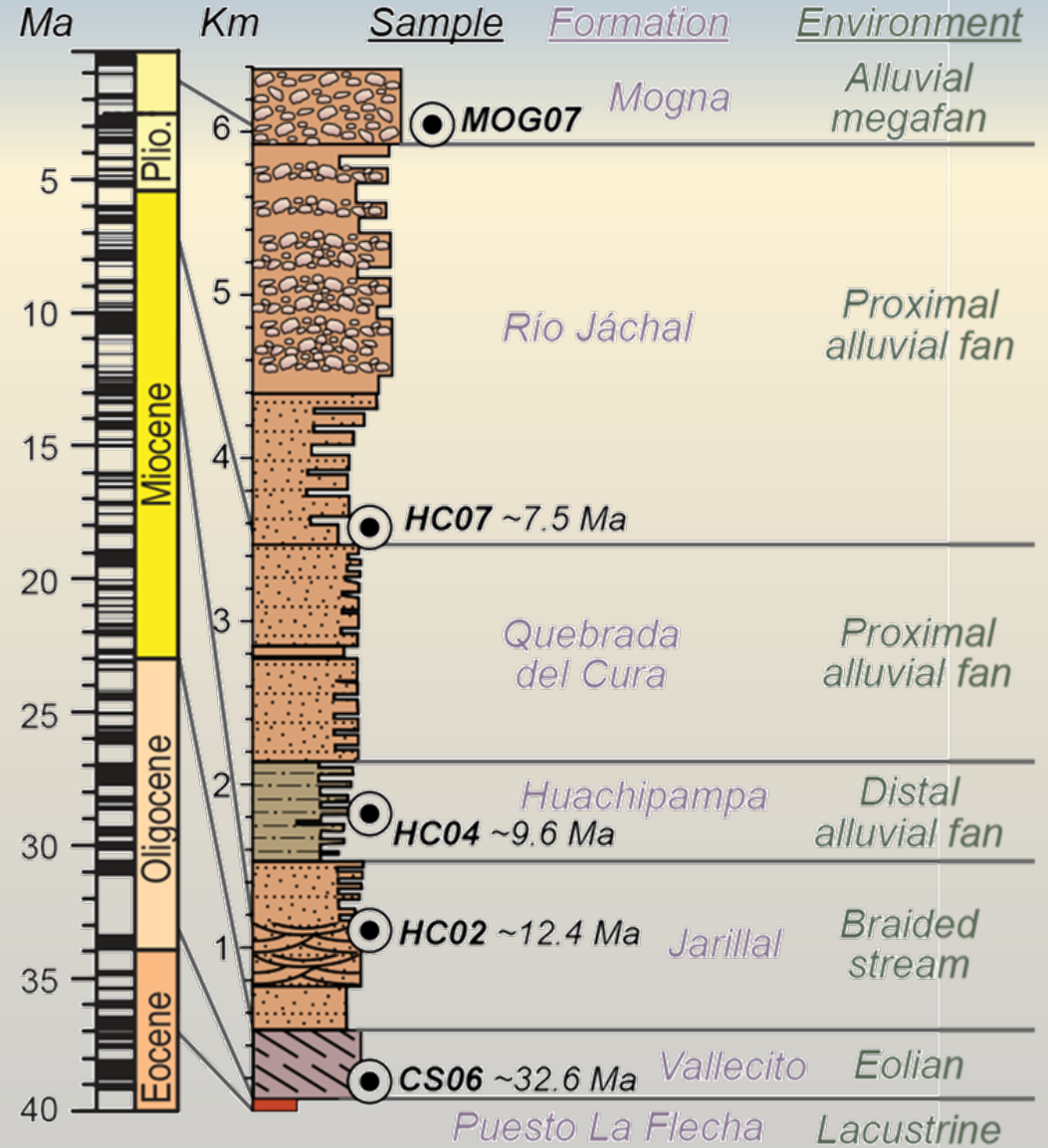
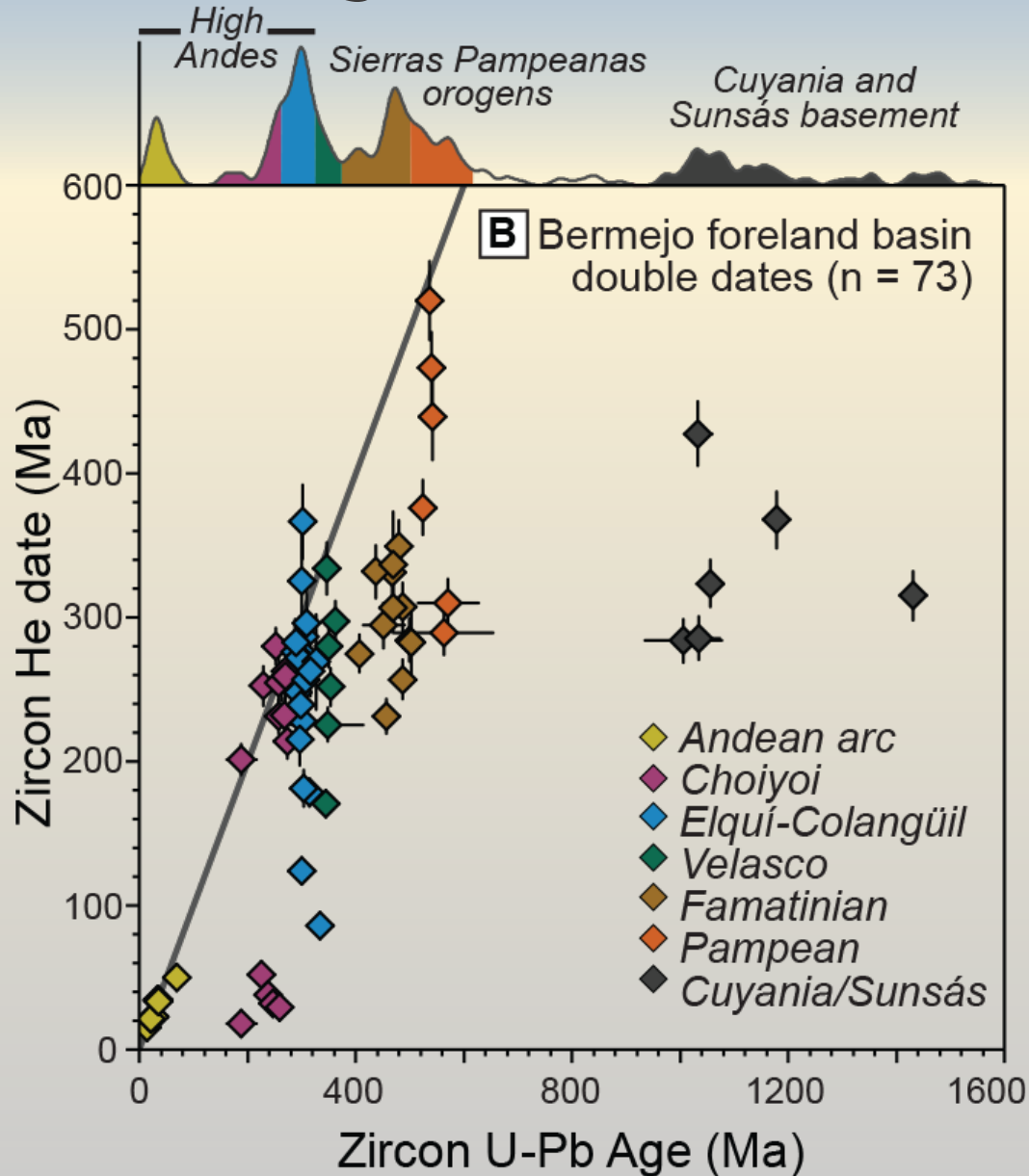


Radiogenic He degassing and measurement



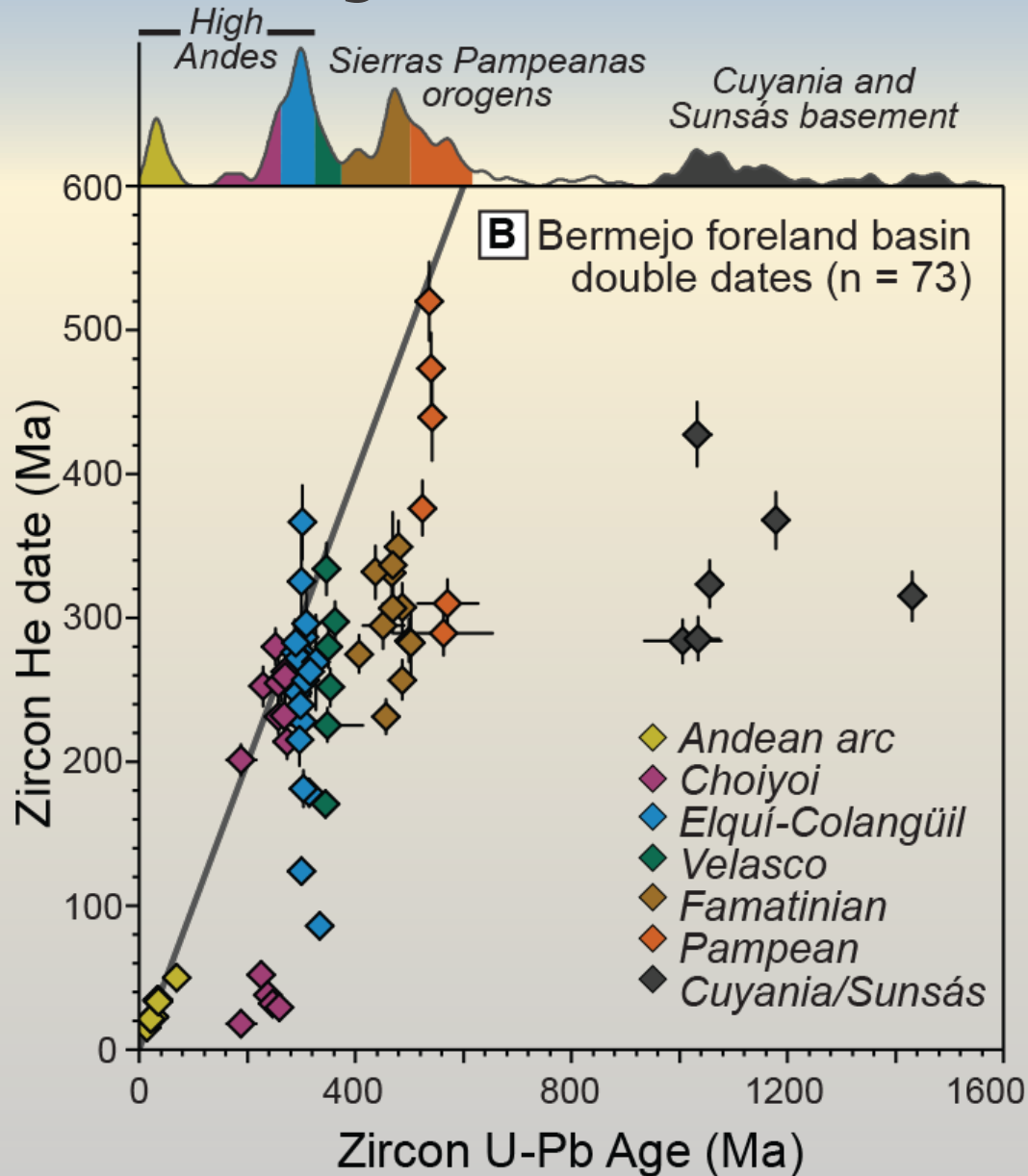
Dissolution and parent U-Th chemistry

# Bermejo Basin: detrital double-dating thermochronology





# Bermejo Basin: detrital double-dating thermochronology



## Challenges with interpreting detrital thermochronology data

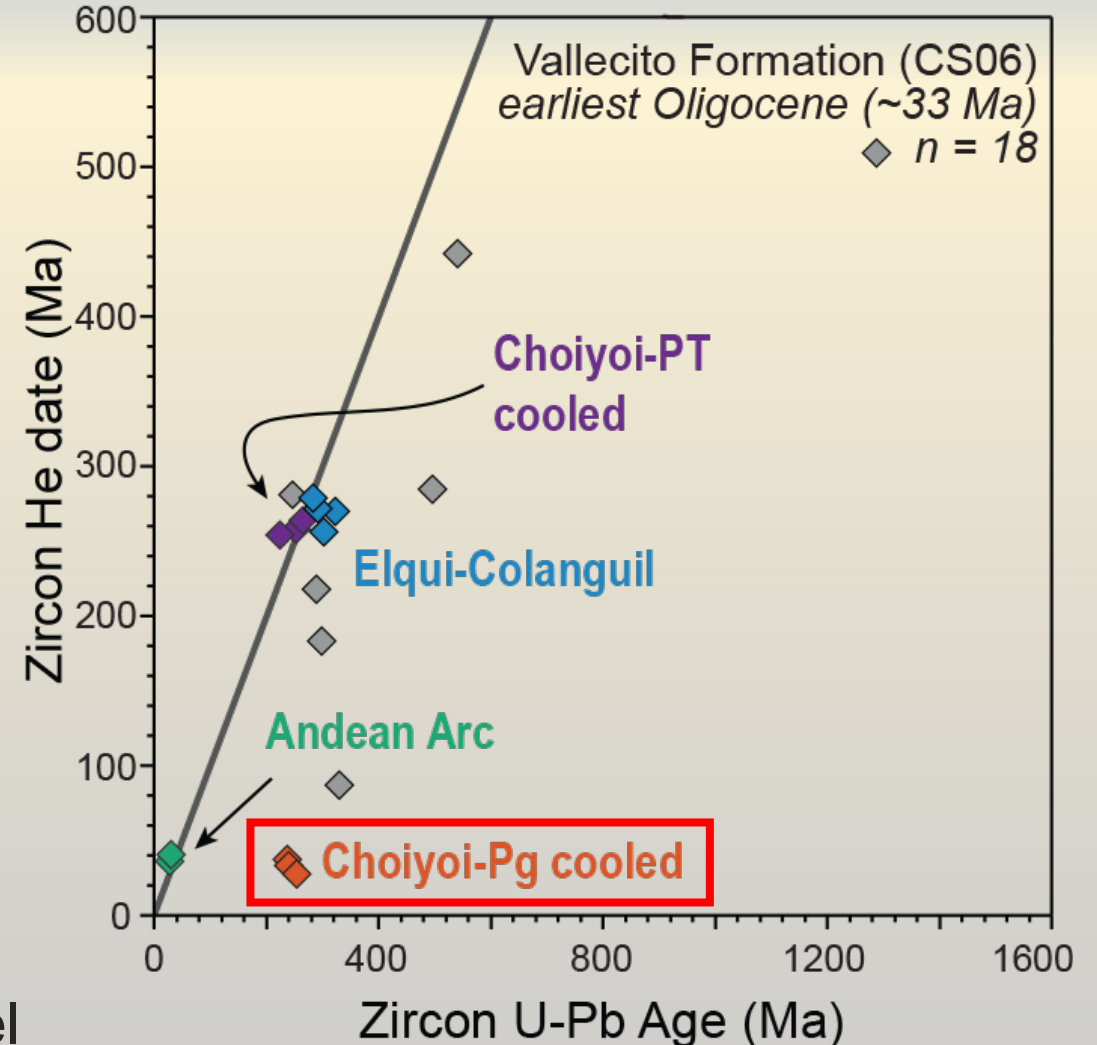
- 1) Missing geologic context available for in-situ samples (bedrock/outcrop/core)
- 2) Potential source areas have insufficient t-T history characterization  
→ Precludes useful MDS or unmixing models
- 3) How to assess effects of post-depositional heating and diffusive He loss

# Thermal history modeling of detrital U-Pb-DZHe Data

HeFTy software (v 2.0 *Ketcham, 2024*)

1. Identify define distinct **U-Pb-DZHe modes**
2. Construct inverse model to **test a specific hypothesis** or t-T scenario, constrained by other geologic information, e.g., U-Pb crystallization age, uniformities, burial histories
3. Invert (U-Th)/He data to resolve possible t-T histories permitted by data and diffusion kinetics
4. Leverage shared basin t-T history and other DZHe modes for whole sample to refine model and evaluate results

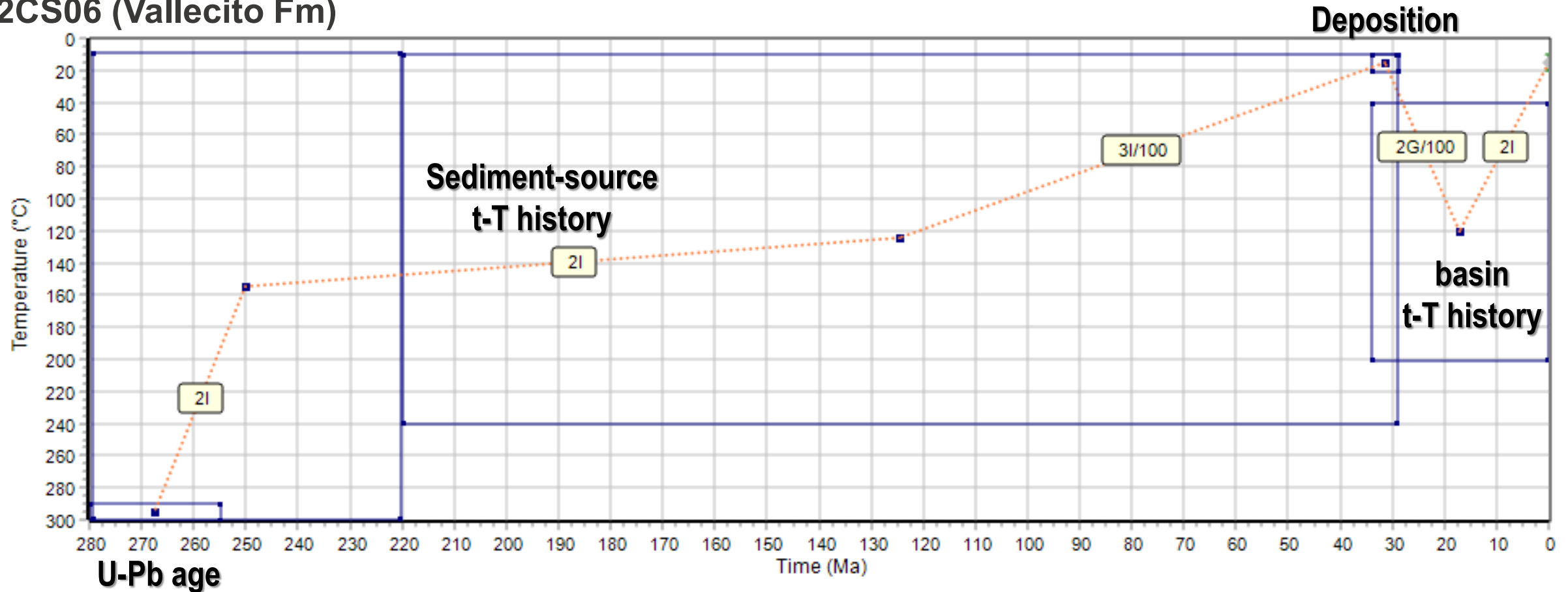
Detrital zircon U-Pb-ZHe modes



# Model 1: undefined burial t-T history

Choiyoi II

## 12CS06 (Vallecito Fm)



**Data:** Choiyoi II ZHe (n=4)  
Sample AHe data (n=2)

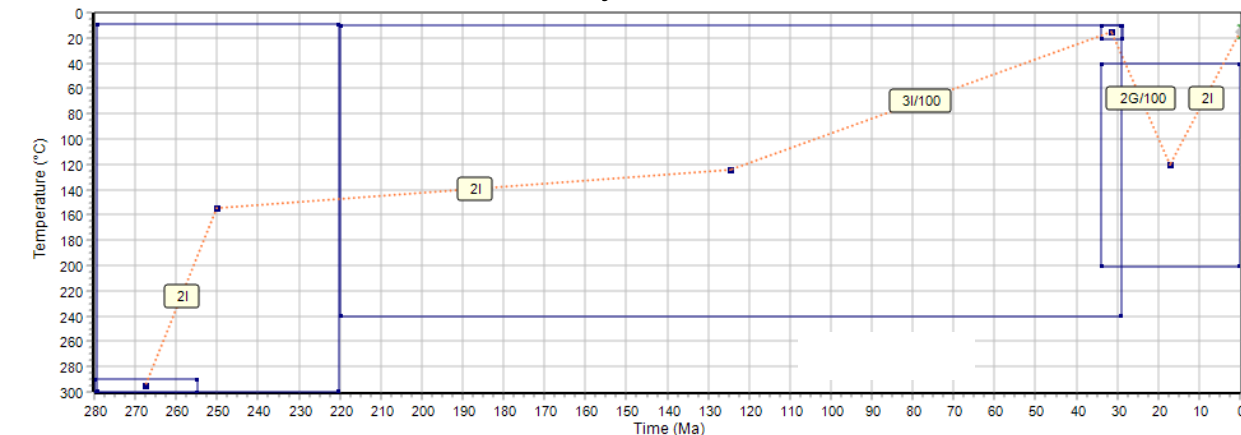
**Constraints:**  
Box 1: DZ U-Pb age @ 300 °C  
Box 2: Cooling  
Box 3: Surface or reheating

Box 4: Sample deposition  
Box 5: Burial reheating (80-200 °C, 34-0 Ma)  
Surface conditions 15 ± 5 °C

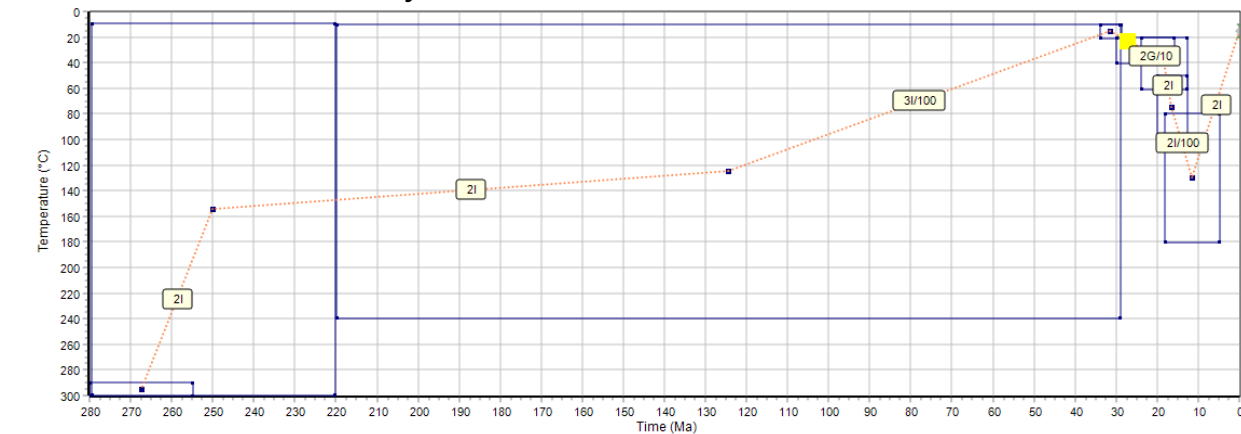
**Kinetics:** RDAMM (Flowers et al., 2009)  
ZDAMM (Guenther et al., 2015)

HeFTy 2.0 (Ketcham 2024)

**Model 1: undefined basin t-T history**

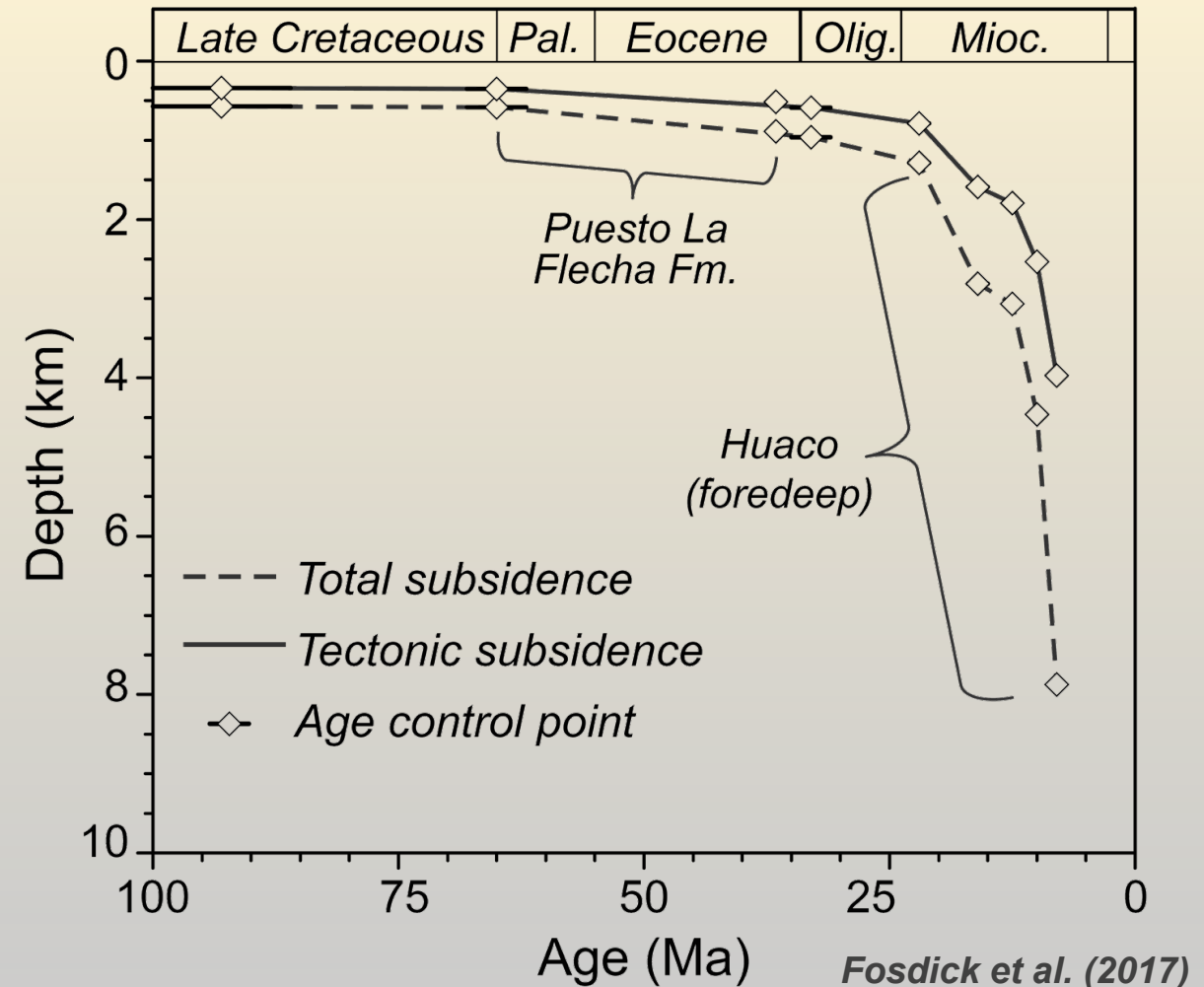


**Model 2: constrained by basin subsidence**

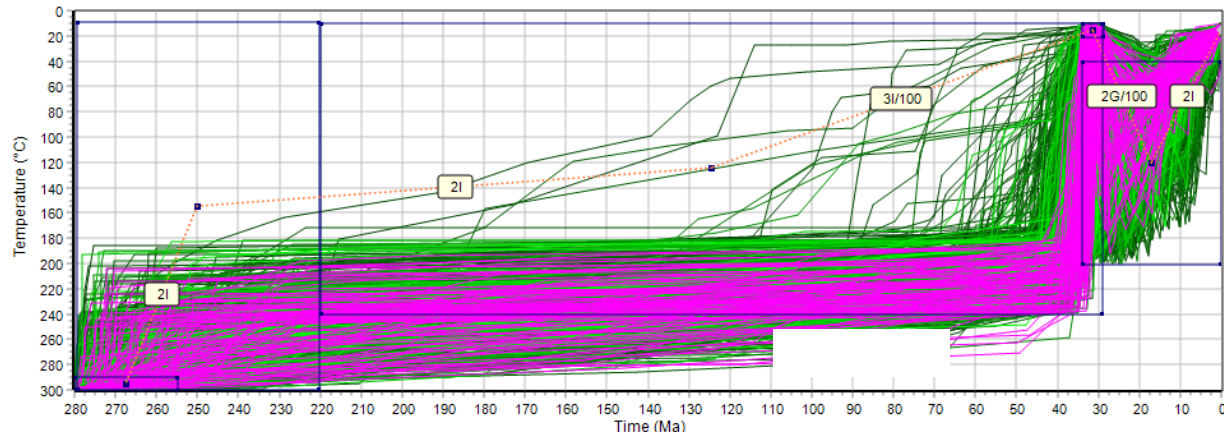


# Model 2: constrained by basin subsidence record

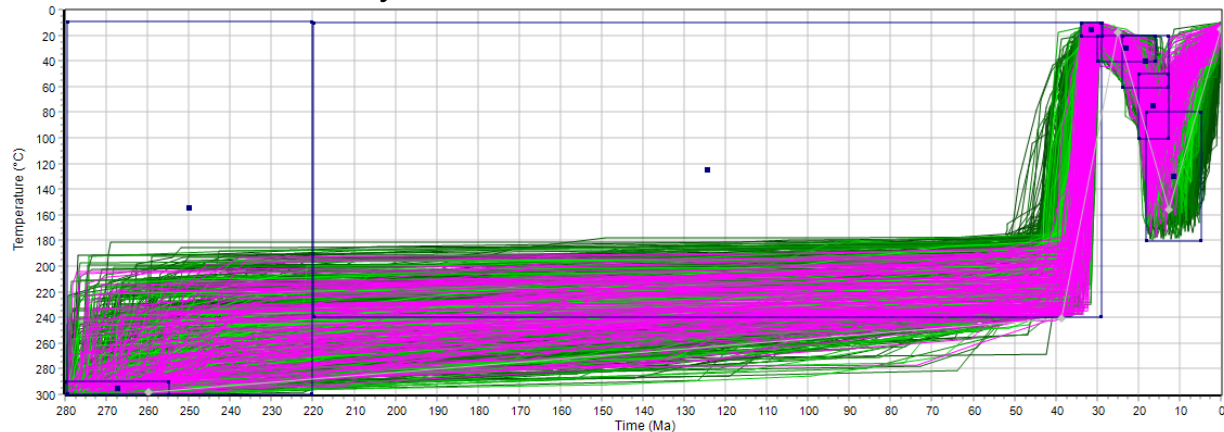
The most complete geologically constrained model



**Model 1: undefined basin t-T history**

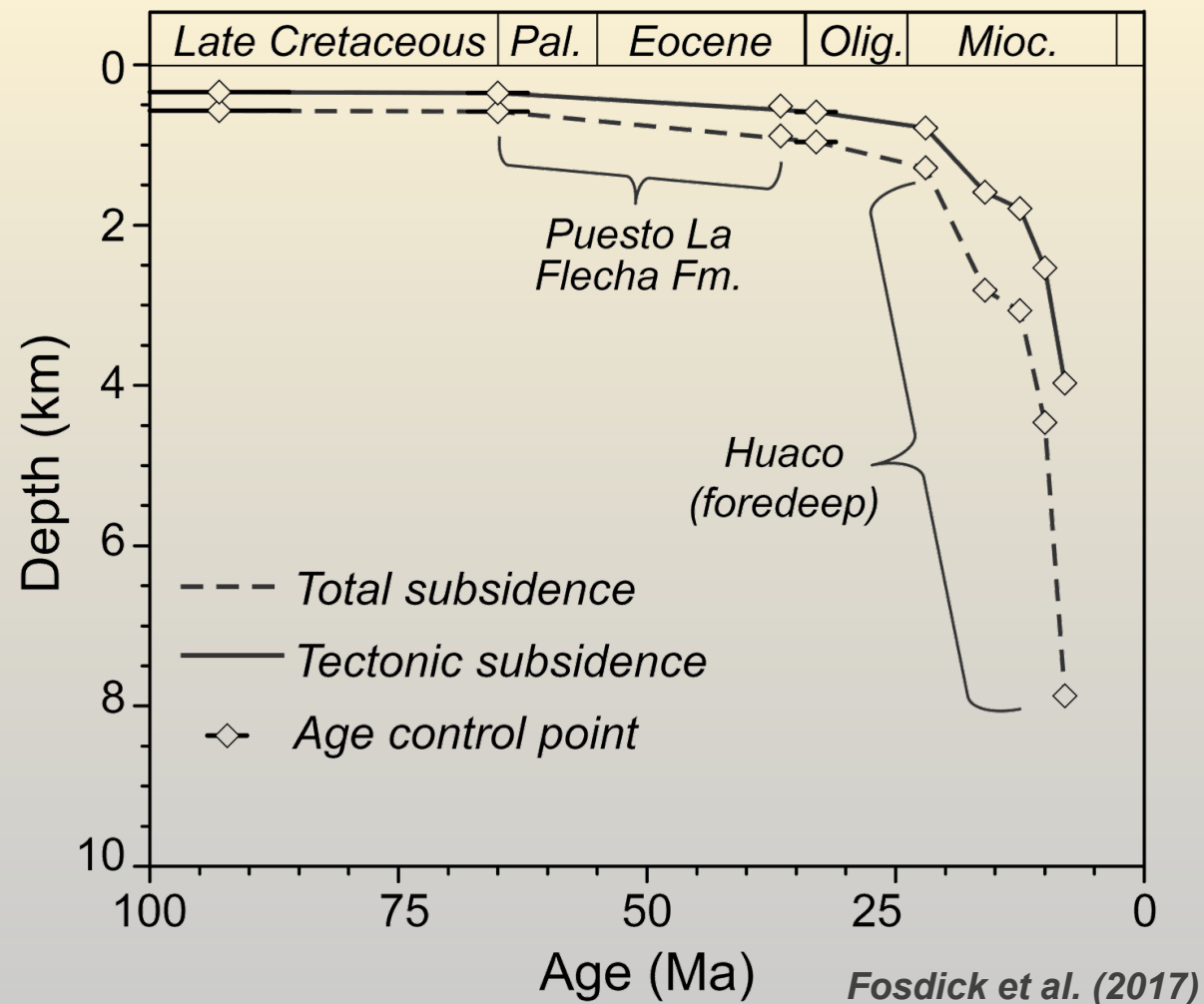


**Model 2: constrained by basin subsidence**

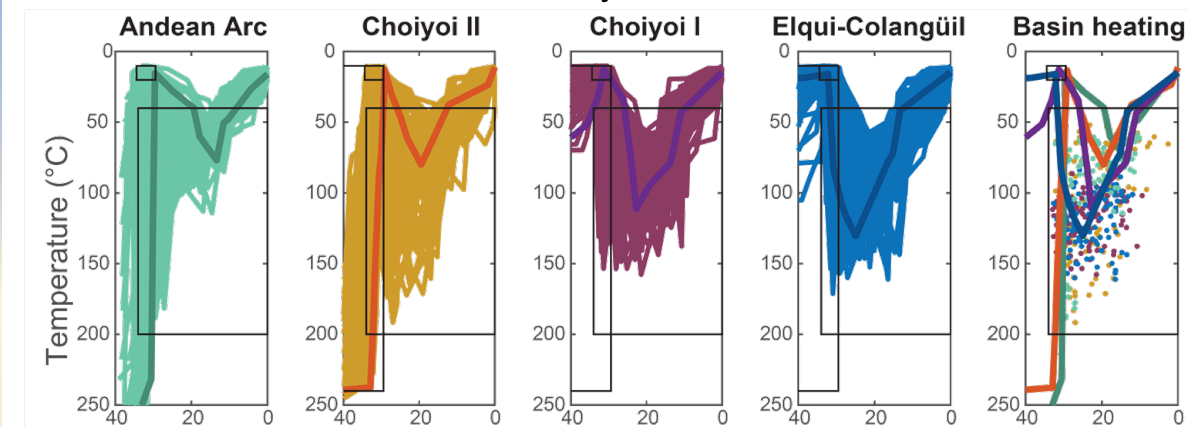


# Model 2: constrained by basin subsidence record

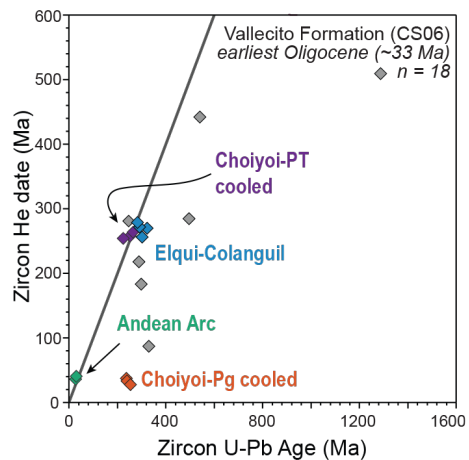
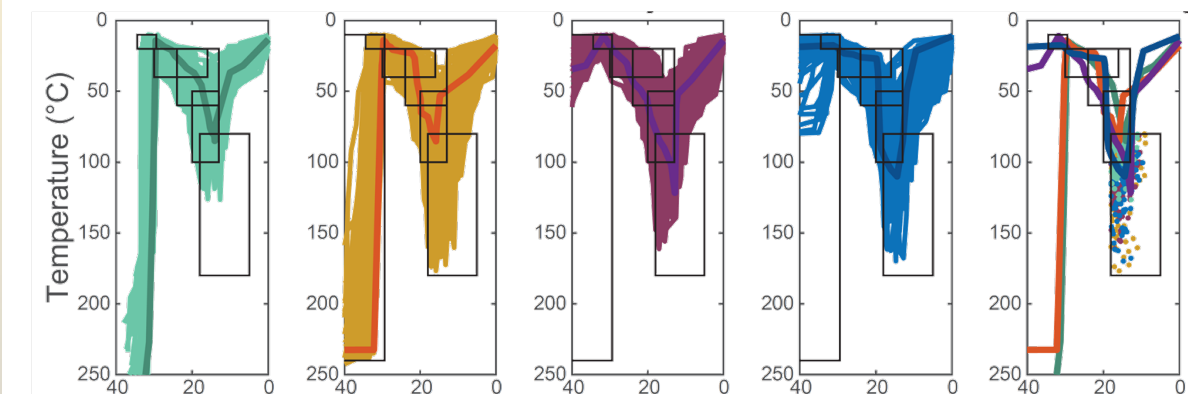
The most complete geologically constrained model



## Model 1: undefined basin t-T history



## Model 2: constrained by basin subsidence

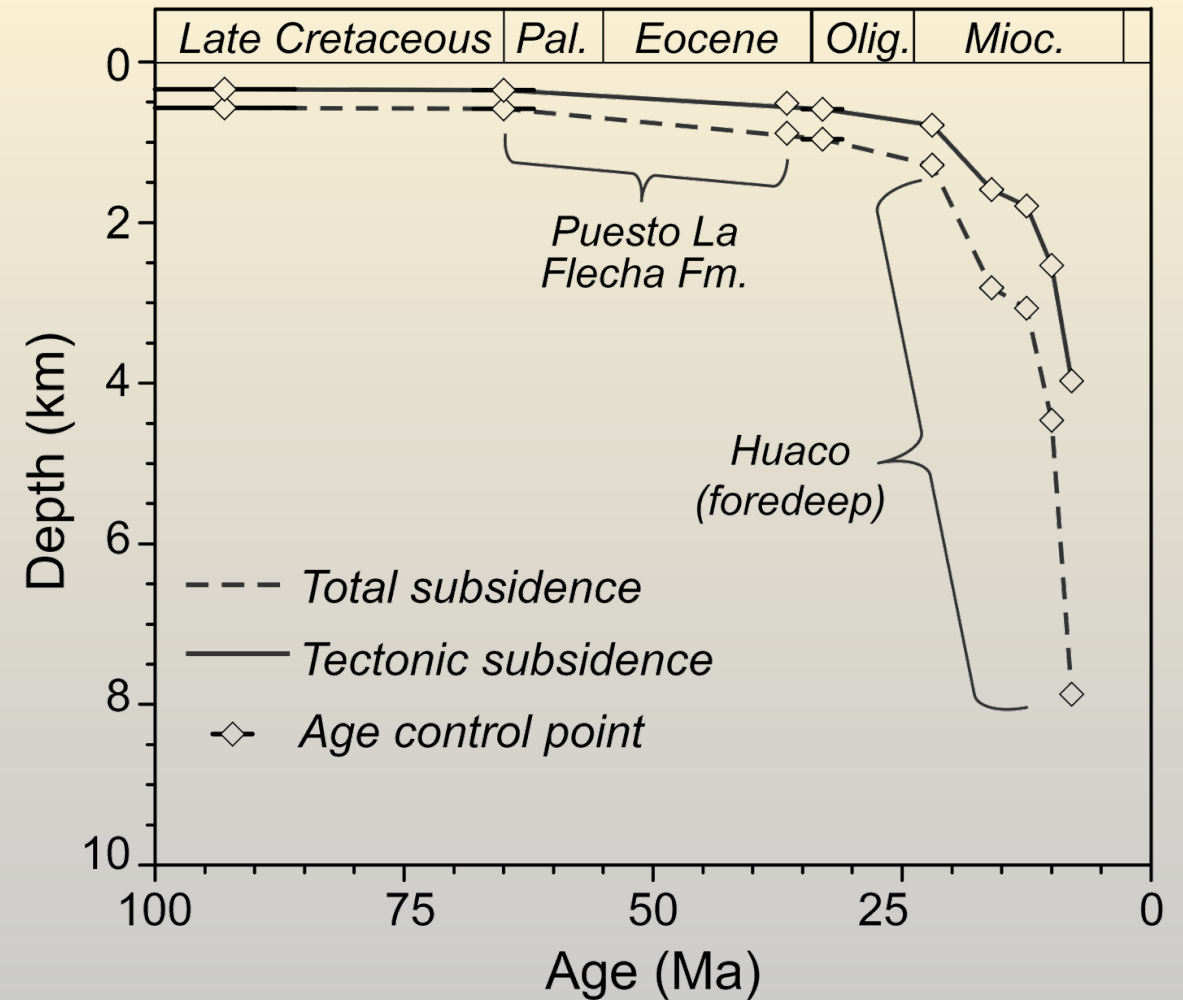


### HeFTy t-T paths

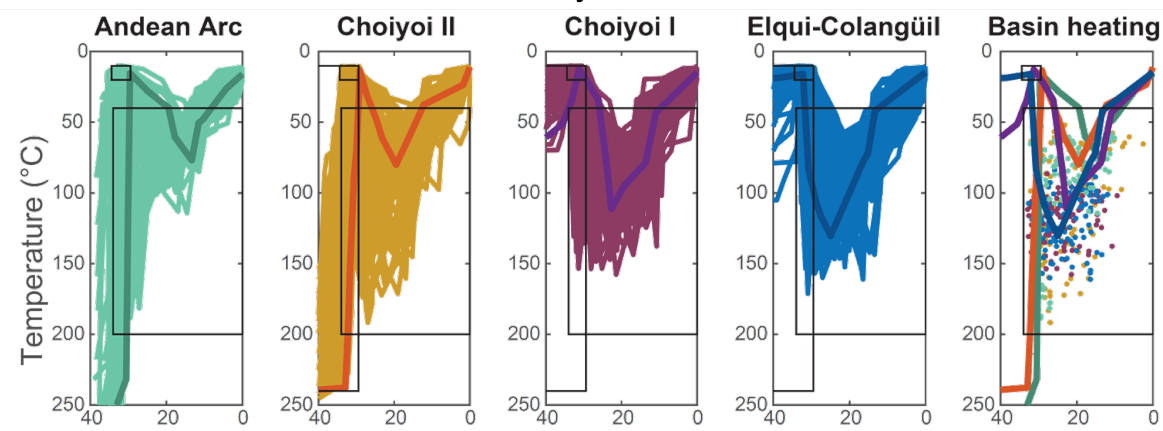


# Model 2: constrained by basin subsidence record

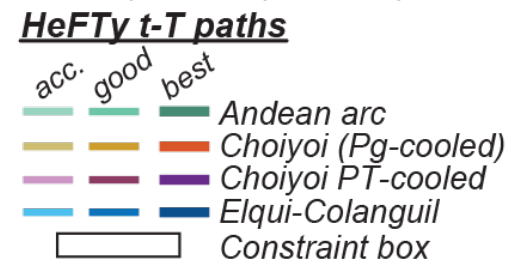
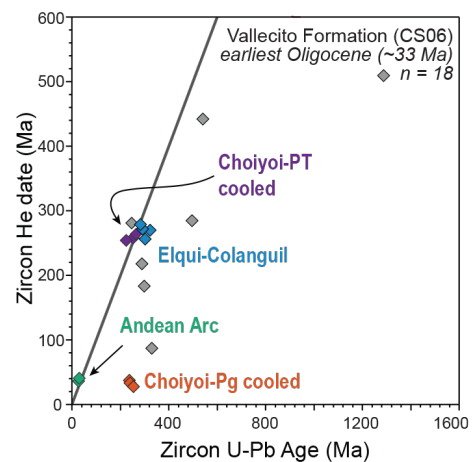
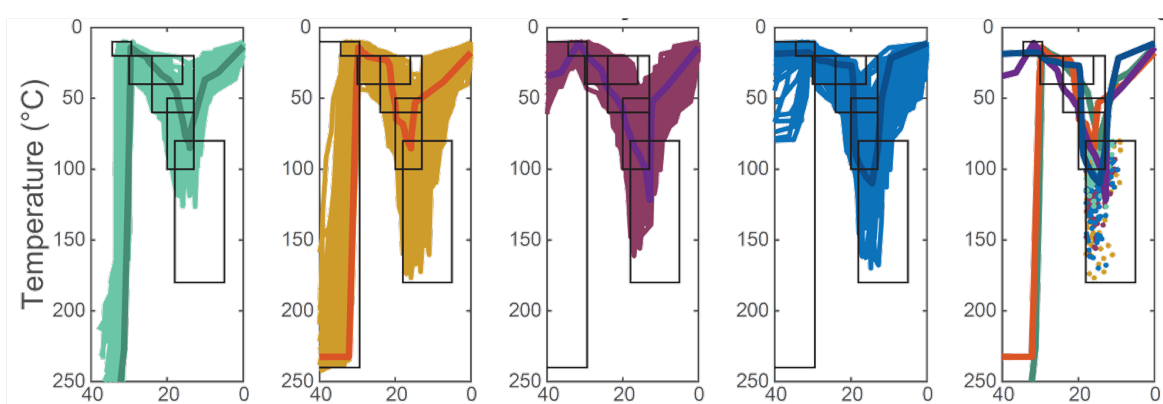
The most complete geologically constrained model



### Model 1: undefined basin t-T history



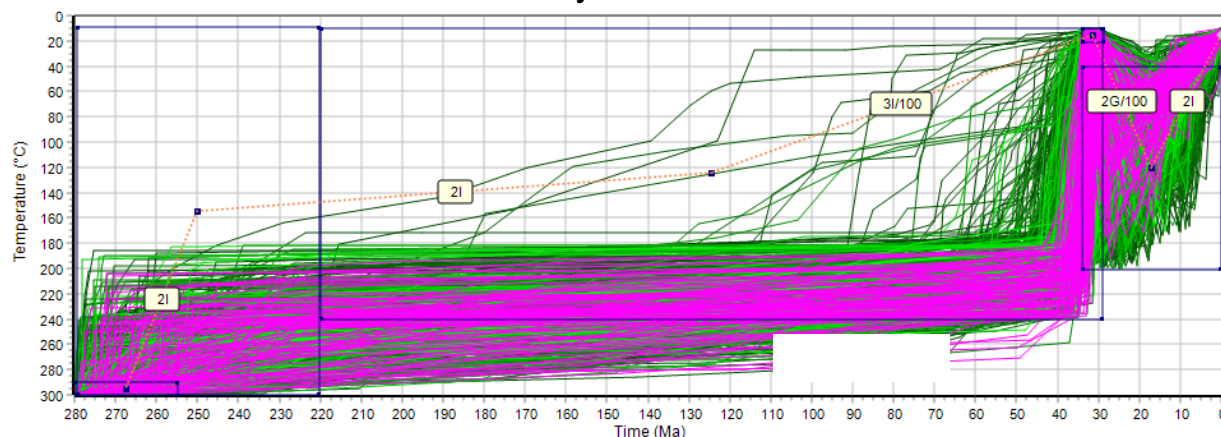
### Model 2: constrained by basin subsidence



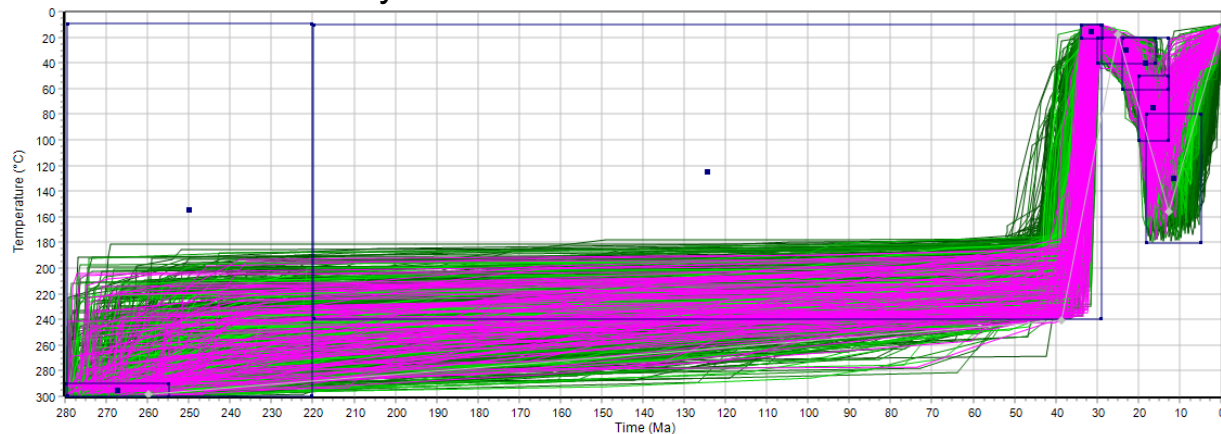
# Model 3: constrained by basin subsidence record *and* requires compatible basin t-T history that satisfies all DZHe modes

1. Evaluate overlapping basin t-T histories from Model 2
2. Refine basin t-T constraints to satisfy all modes
  - most retentive grains?
  - Younger DZHe mode(s)?
  - Depositional age important

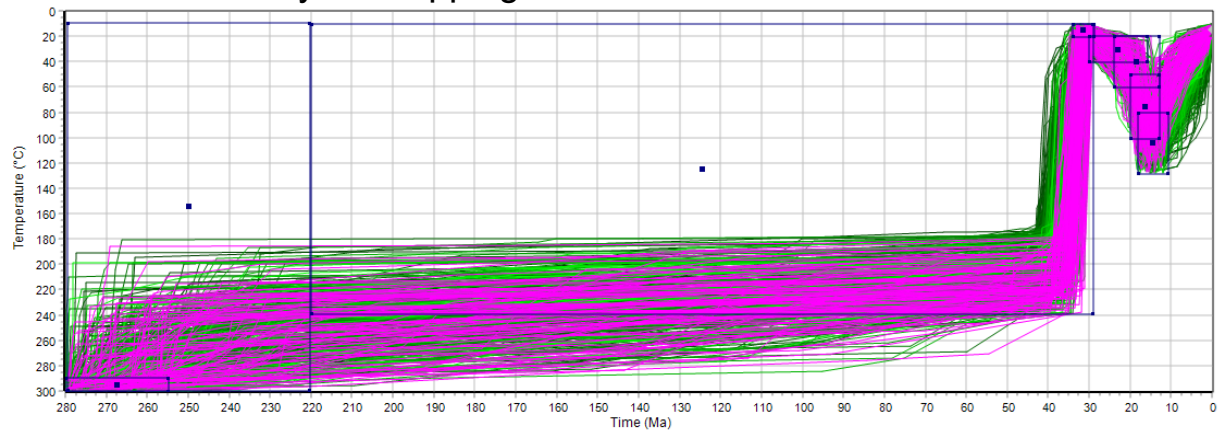
**Model 1:** undefined basin t-T history



**Model 2:** constrained by basin subsidence



**Model 3:** refined by overlapping DZHe modes

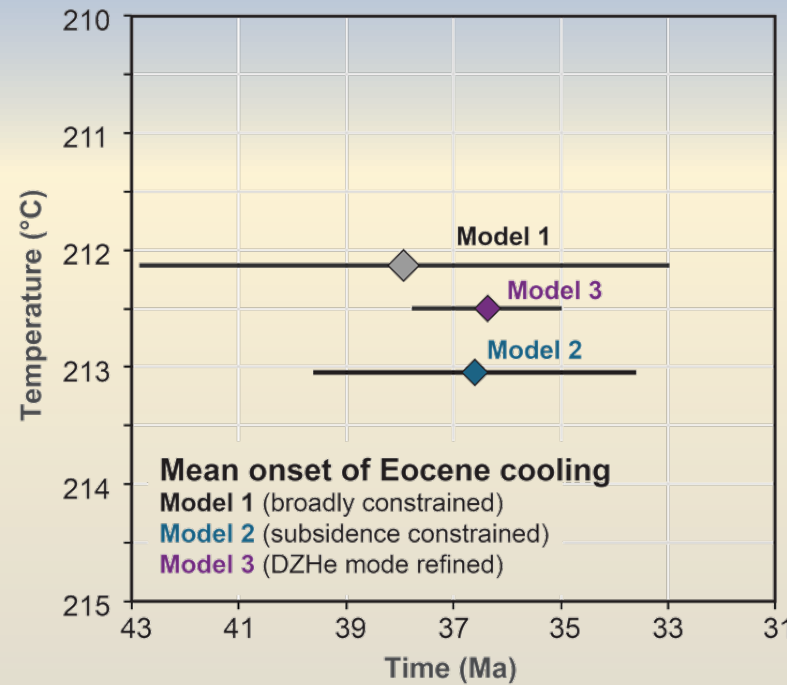
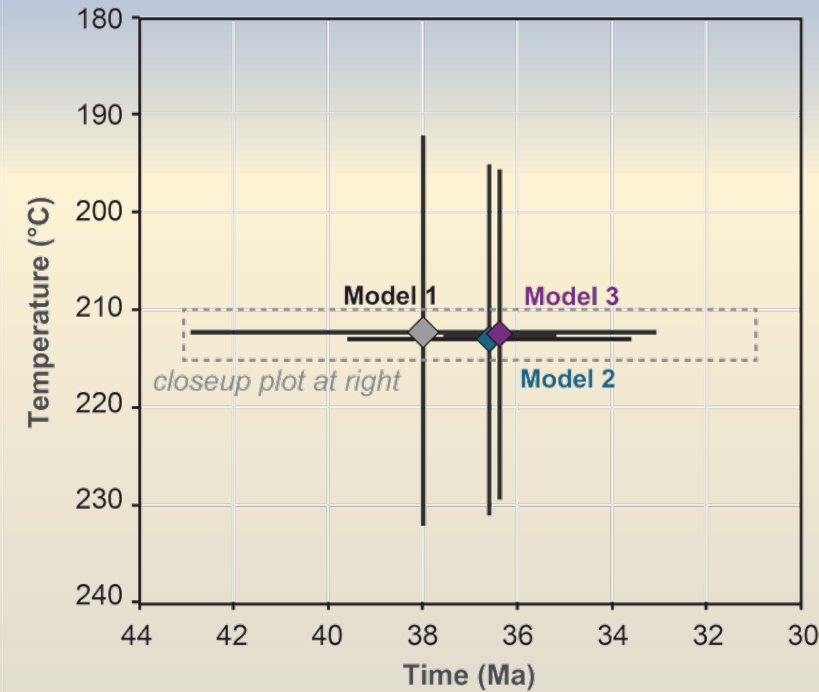


**Model 3:** constrained by basin subsidence record *and* requires compatible basin t-T history that satisfies all DZHe modes

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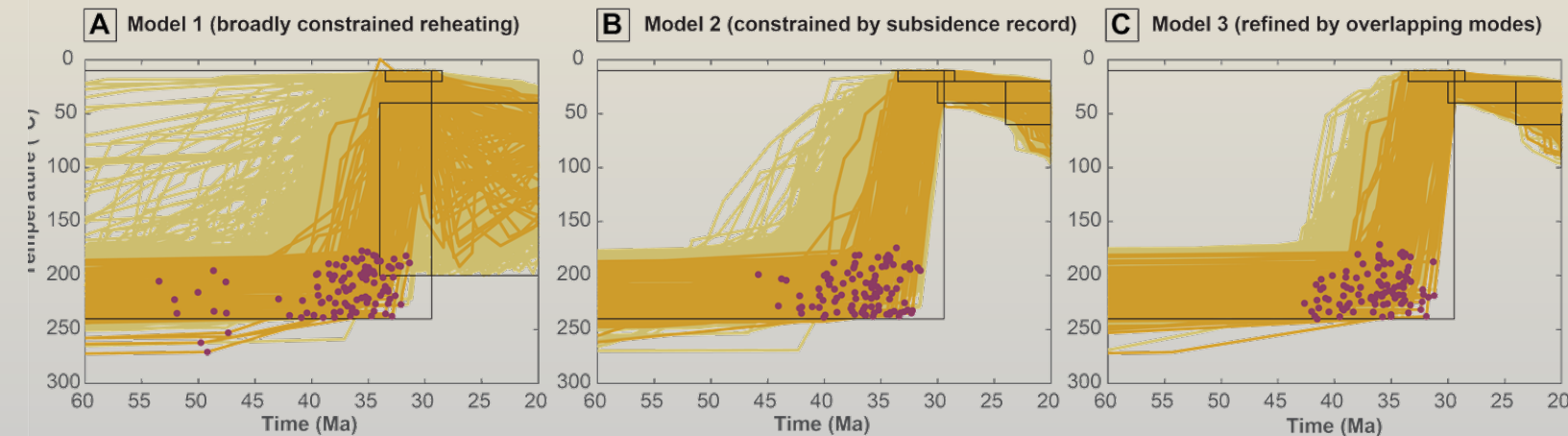


# Model comparison and sensitivity evaluation



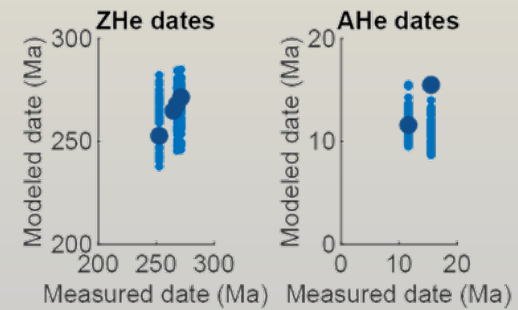
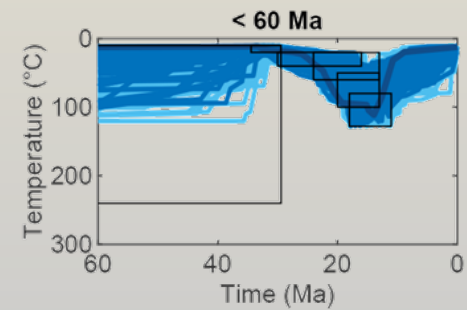
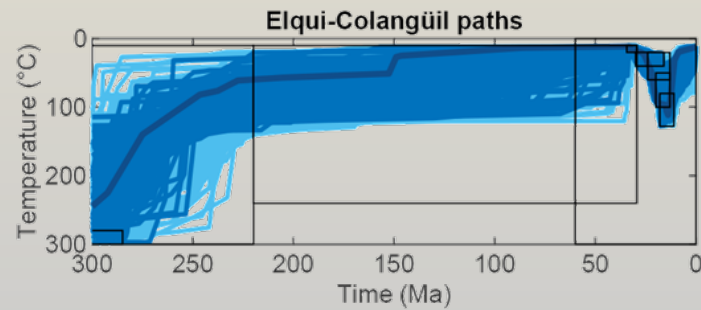
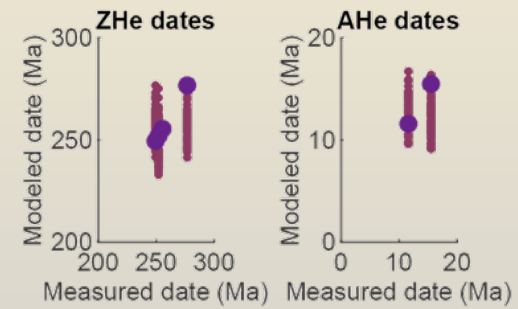
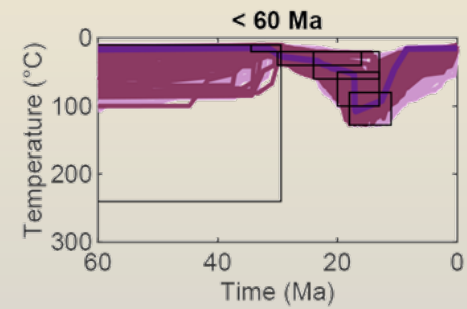
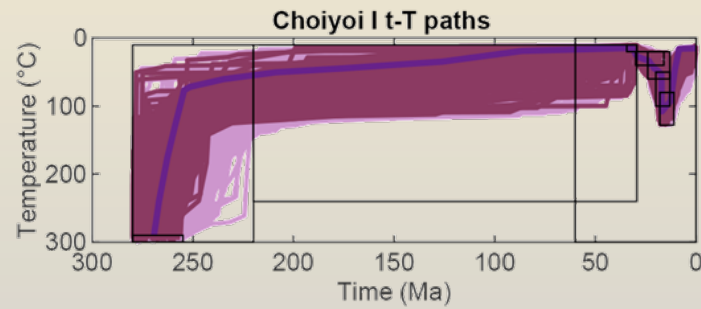
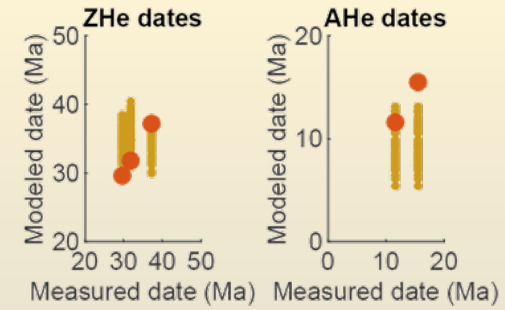
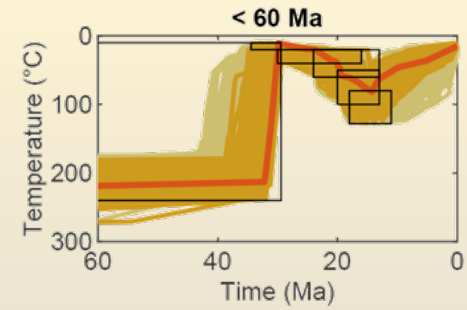
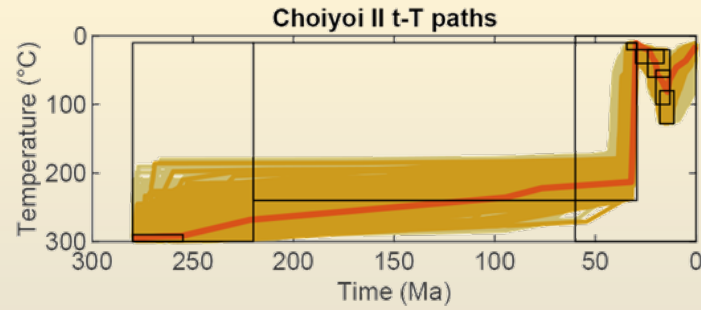
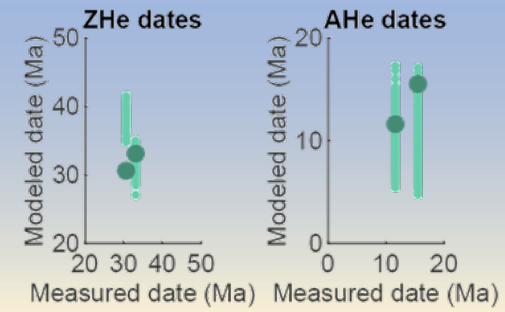
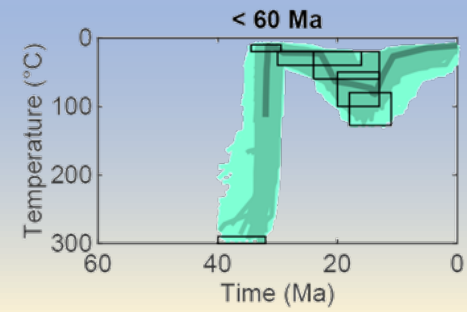
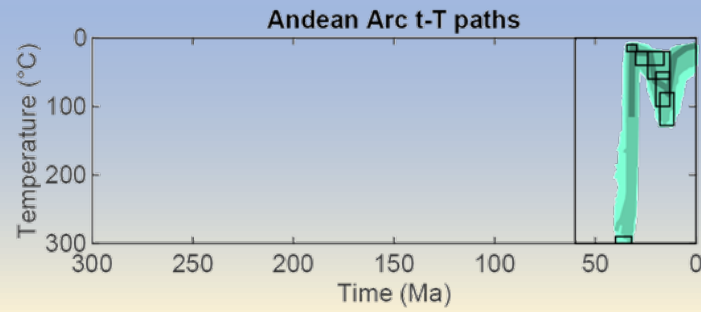
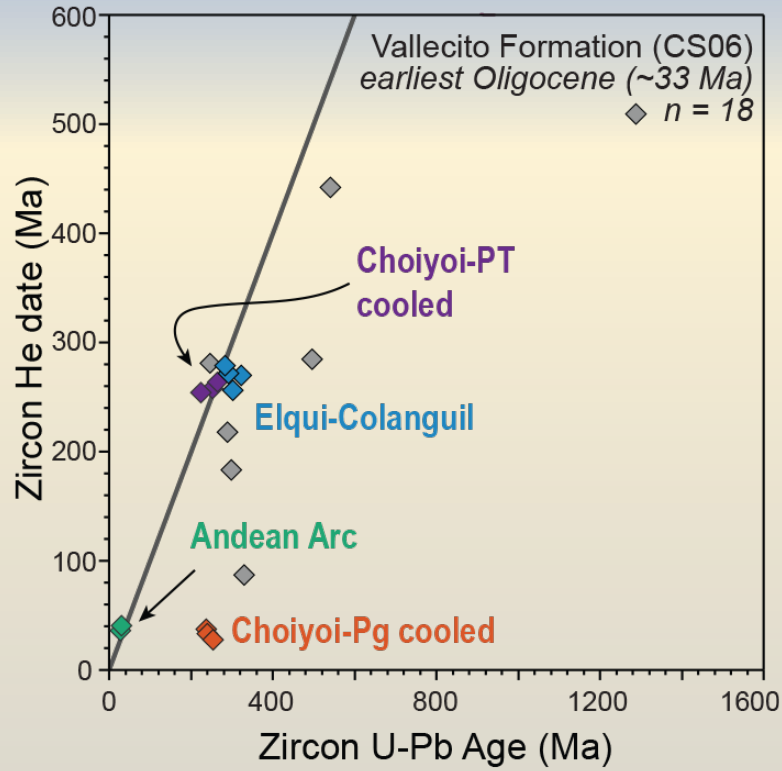
Models that leverage shared basin t-T history and satisfy all DZHe modes refine estimates in **onset of cooling** and maximum post-depositional **burial heating**

For Choiyoi II DZHe mode, lag time ( $t_c - t_d$ ) is refined from up to 10 Myr (Model 1) versus 2-5 My (Model 3)

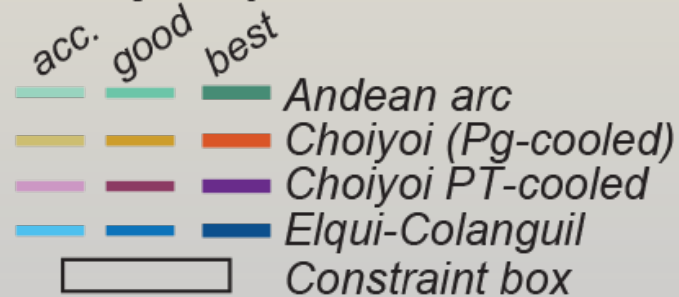


Reheating 100-120°C at 23 Ma (Model 1) versus 90-110°C at 15 Ma (Model 3) prior to cooling and inferred basin inversion

# All DZHe Modes



## HeFTy t-T paths

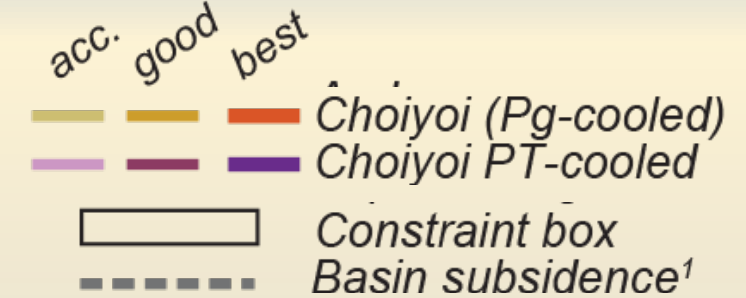


# Implications for Eocene hinterland unroofing

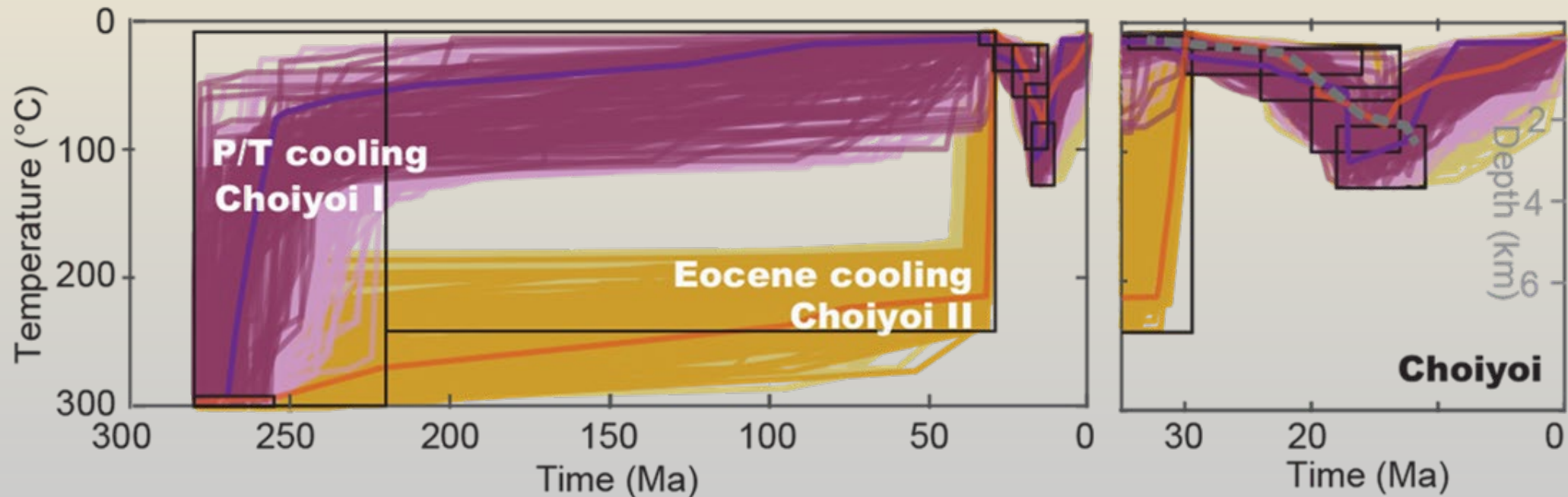
Two distinct t-T histories of Choiyoi-derived detritus requires multiple sources:

- *pre-orogenic* signature (Choiyoi I)
- *synorogenic* unroofing signature (Choiyoi II)

## HeFTy t-T paths

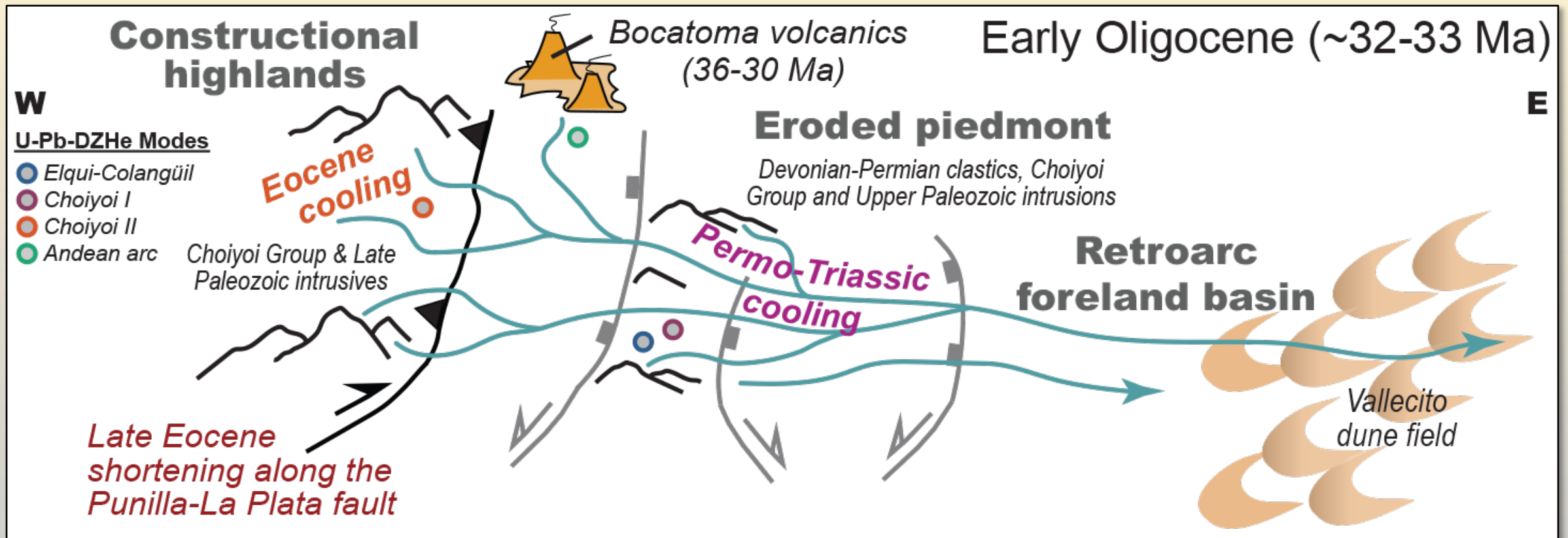


Onset of cooling ~38–35 Ma; 2-5 Myr lag time between cooling and deposition

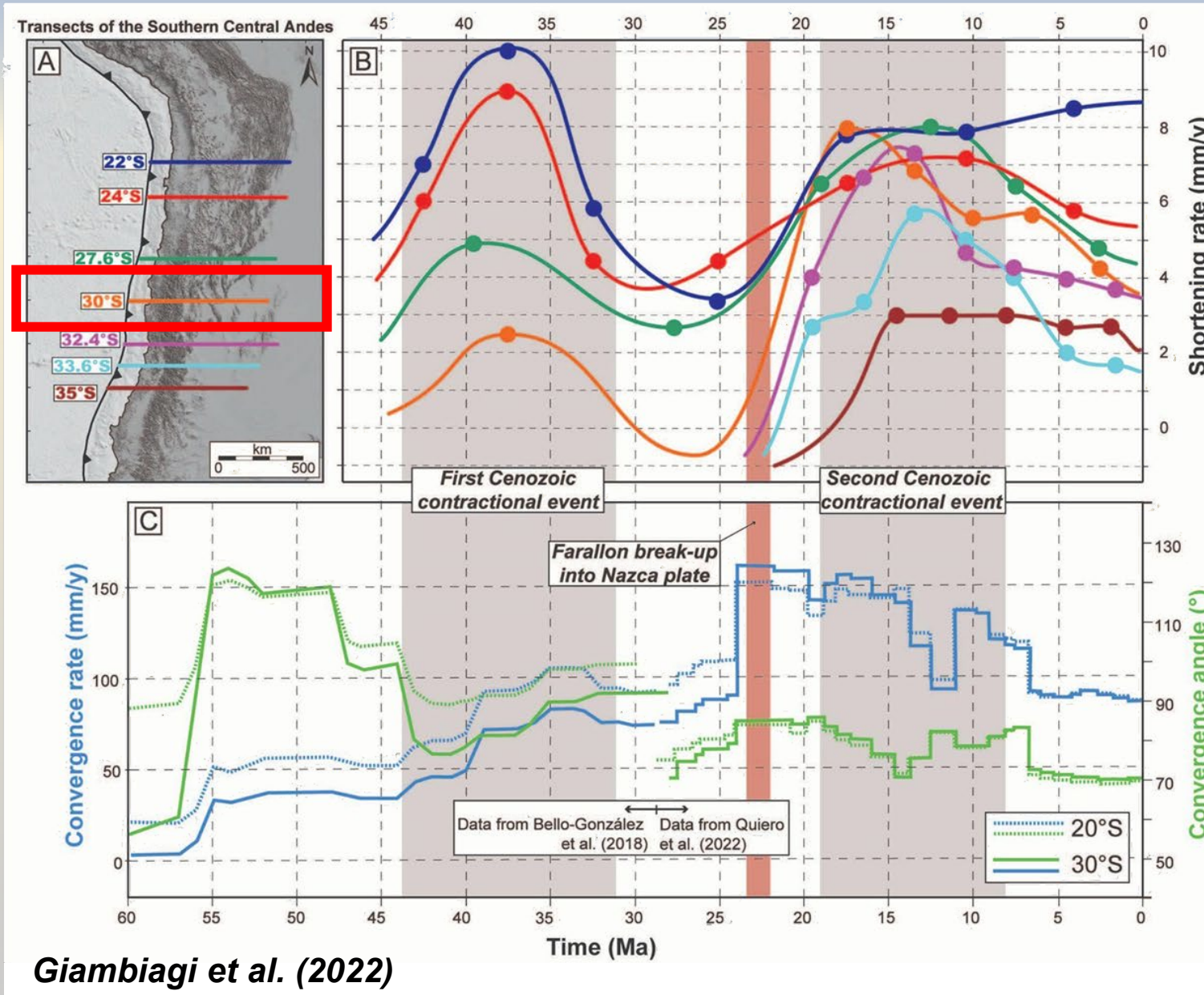


# Implications for Eocene hinterland unroofing

Rejuvenated phase of **shortening, exhumation, and sediment routing** to the distal foreland at 30°S in late Eocene–early Oligocene time



# Implications for Eocene hinterland unroofing



Giambiagi et al. (2022)

## Along-strike development of Andean shortening

Two major phases of Cenozoic orogenesis across the Central Andes

Integrated foreland basin system across changes in tectonic subduction mode, basement structures, and climate gradients

Southward decrease in Eocene shortening rate → define the southernmost extent of this contractional phase?

# Concluding points and take-aways

- Detrital U-Pb-ZHe thermochronology **differentiates ambiguous** hinterland sediment sources with unique t-T histories
- **Modeling approach for discrete detrital U-Pb-ZHe modes** can be powerful tool to resolve source-to-basin t-T histories
  - opportunities for interpreting detrital datasets as double-dating thermochronology capabilities expand (e.g., LADD)
  - test specific hypotheses for a given sediment source
  - refine lag time and peak reheating estimates
- In the southern Central Andes  $\sim 30^\circ\text{S}$ , **rapid latest Eocene hinterland cooling** and inferred exhumation
  - Difficult to reconcile with neutral-state stress model for the Paleogene Andes → **rejuvenated phase** of shortening, exhumation, and sediment routing to distal foreland



**Thank you!!**

